Estimating the Influences of Flexible POC:POP Stoichiometry

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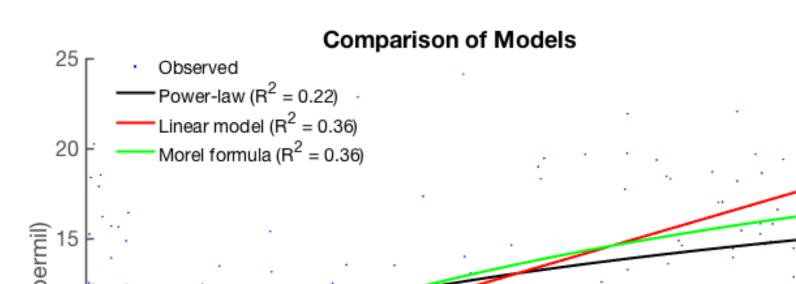


Introduction

Export production of particulate organic matter (POM) from the surface to the deep ocean is a key driver of global carbon cycle. The amount of carbon (C) removed from the surface ocean by this export depends critically on the elemental ratios in POM of C to nitrogen (N) and phosphorus (P), two essential nutrients that limit productivity. Here we developed a simple power law model with a stoichiometry sensitivity factor, which is able to relate a fractional increase in C:P of POM to a fractional phosphate ambient decrease in concentration

Estimating stoichiometry sensitivity factor, s

1. Against Global Compiled Data (Galbraith and Martiny, 2015)

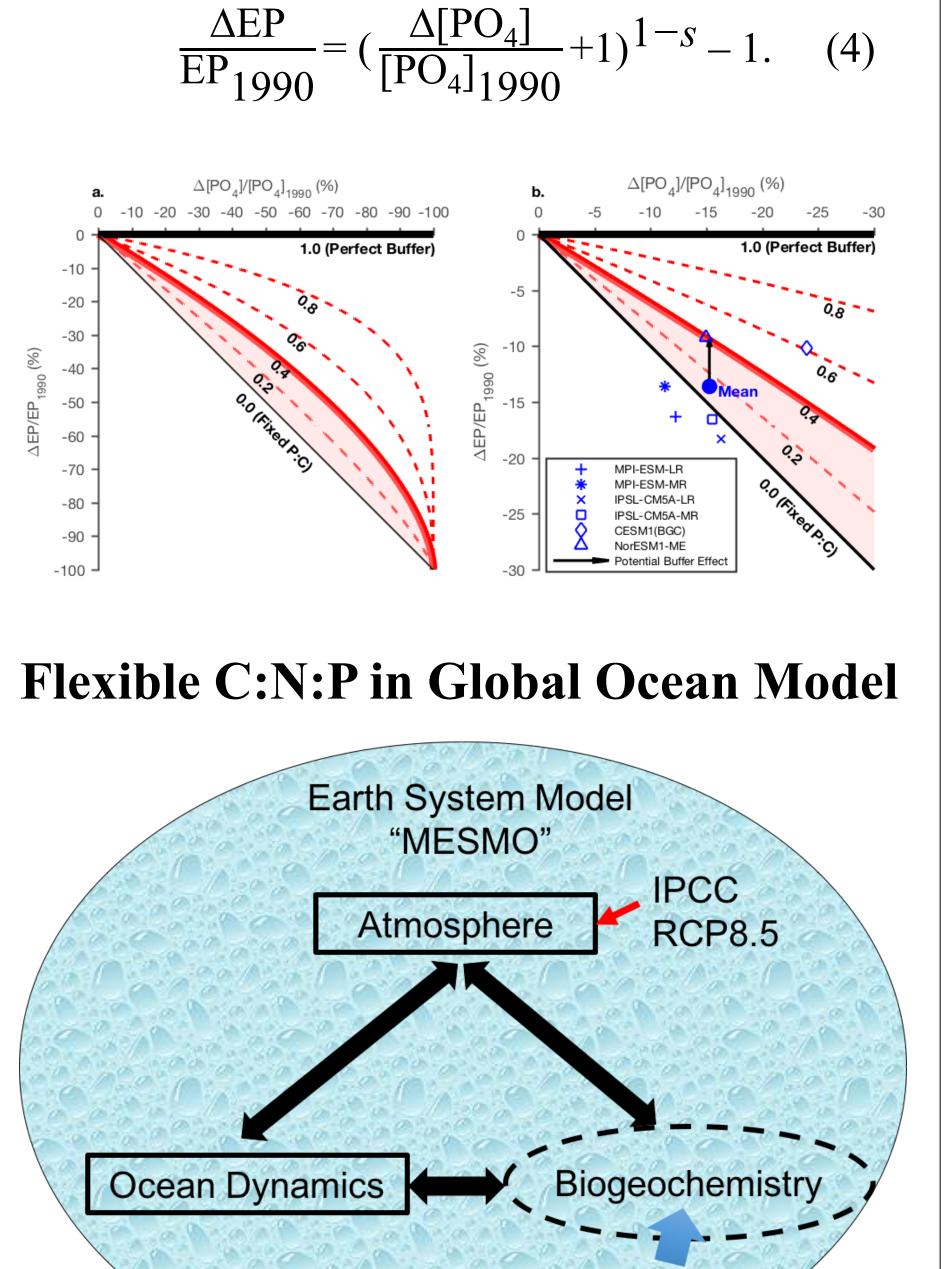


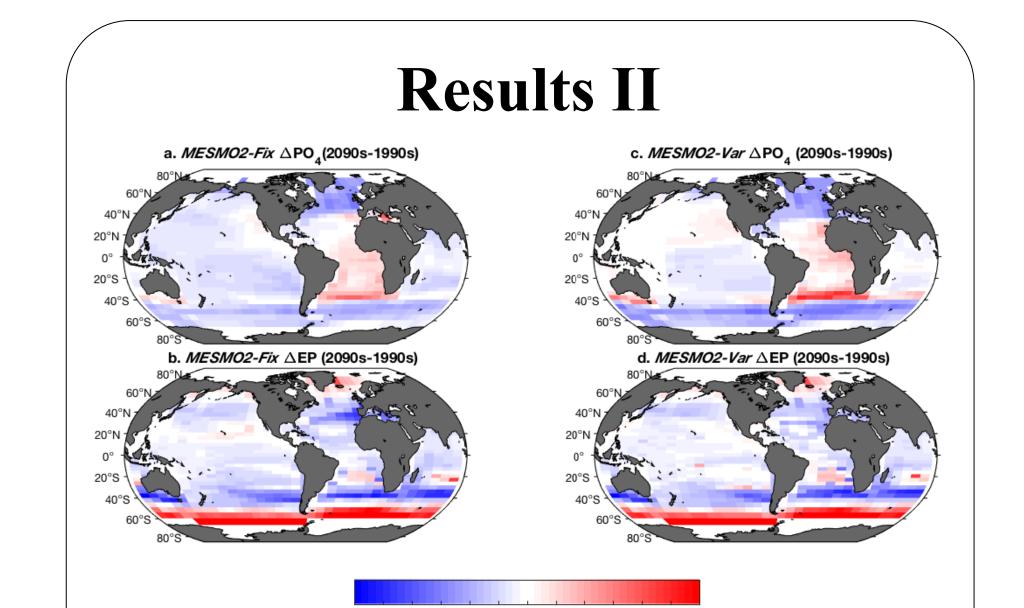
First-order estimation

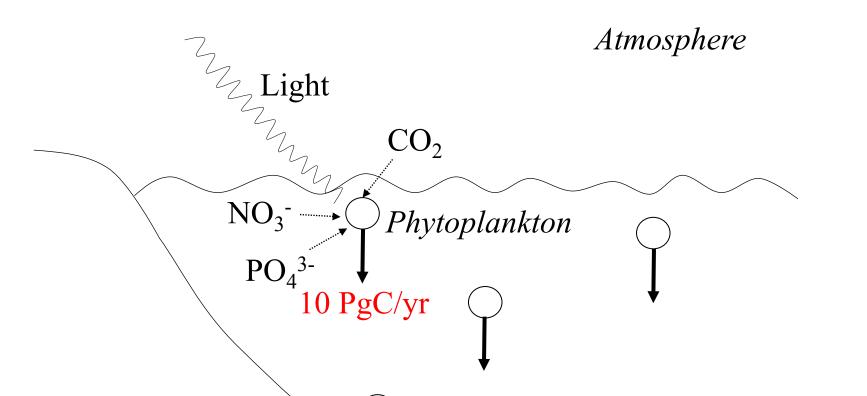
In the limiting case where export production (EP) of carbon is a function of PO_4 and P:C

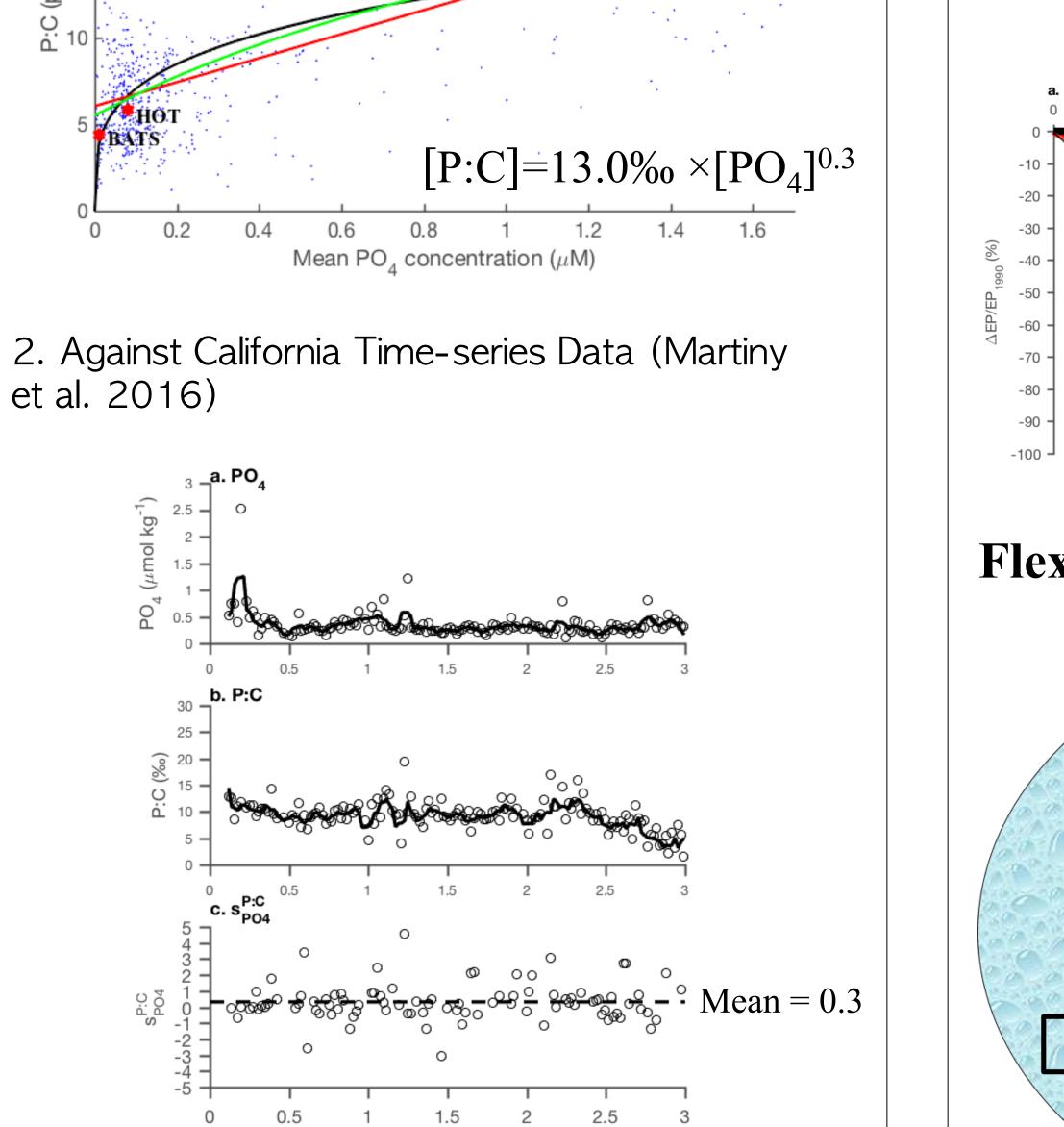
stoichiometric compensation effect referenced to

a recent time (e.g. decade of the 1990s) as,









Simulated fractional changes in surface PO_4 (0-100 m) and total EP in 2090s relative to 1990s with fixed C:N:P (a, b) and variable C:N:P (c, d) under RCP8.5 scenario. **Results III** (a) Stoichiometric buffer effect in 2090s.



Simple Stoichiometry Models

1) Power-law model (this study): Sensitivity of P:C with respect to surface phosphate, s:

 $s = \frac{\partial [P:C]/[P:C]}{\partial [PO_{A}]/[PO_{A}]} = \frac{\partial \ln[P:C]}{\partial \ln[PO_{A}]}.$

(s: stoichiometry sensitivity factor) This is analogous to the famous Revelle factor, R:

 $R = \partial \ln[\mathrm{CO}_2] / \partial \ln[\mathrm{DIC}]$

Solve (1) to express P:C as a function of PO_4 :

 $[P:C] = [P:C]_0 (\frac{[PO_4]}{[PO_4]_0})^S$

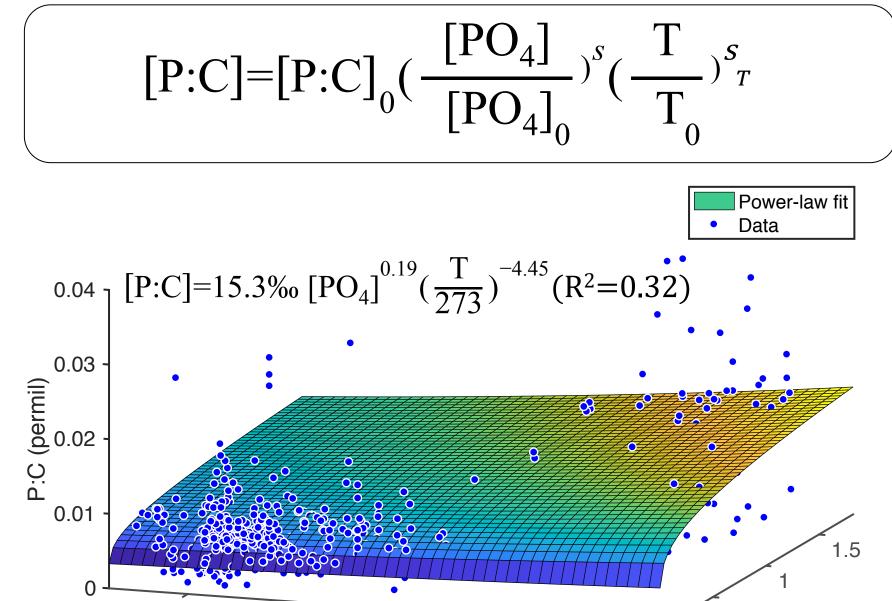
Our best estimate of global mean *s* is 0.3-0.4,

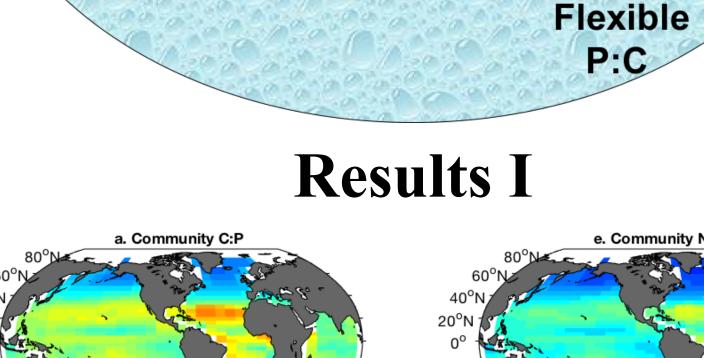
i.e. 1% change in PO_4 translates to a 0.3-0.4%

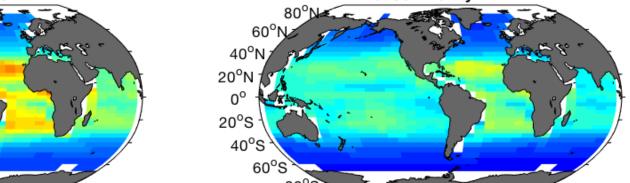
change in P:C of POM.

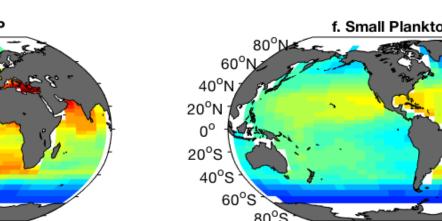
Extension of stoichiometry model: adding temperature dependence

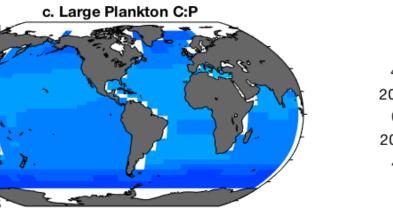
 $s_{\rm T} = \frac{\partial [{\rm P:C}] / [{\rm P:C}]}{\partial {\rm T} / {\rm T}} = \frac{\partial \ln [{\rm P:C}]}{\partial \ln {\rm T}}$

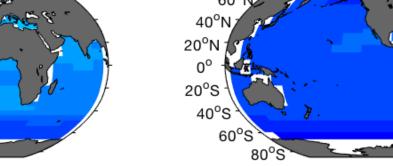




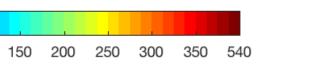


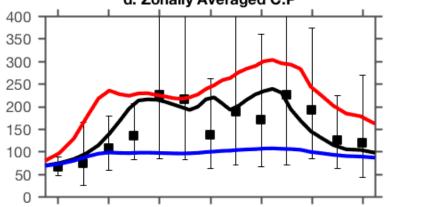






. Zonally Averaged N:F







Positive value indicates positive buffer effect (i.e. smaller reduction in EP for variable C:N:P model). (b-d) Percent change of community, small plankton, and large plankton C:P in 2090s relative to 1990s.

Conclusions

- A new, alternative method for predicting C:P of POM as a function of PO_4 is presented.
- New stoichiometry model can be implemented successfully and easily in a global model to reproduce the C:N:P variability in the ocean. • Flexible C:P can buffer changes in export production under the business as usual global warming scenario by up to 1%.

(1) $([P:C]_0 = reference P:C)$ 2) Linear model (Galbraith and Martiny, 2015): $[P:C] = 6.9 \% \mu M^{-1} \times [PO_4] + 6.0\% (2)$ 3) Morel's formula (Morel, 1987): $[P:C] = [P:C]_{max} \cdot \frac{K_1 + [PO_4]}{K_2 + [PO_4]}$ (3) $(K_1, K_2: constants)$

0	* • •	•				I	
1.1	1.08	1.06	1.04	1.02	0.5	PO ₄ (μM)	
	Tem	perature	(K/273 K	.)			
Stoichio	met	ric (Con	pens	sation	Effect	
					C:N:P =	106:16:1	
C:N:P = 7	106:16 16 N	:1	No Pla	sticity	53	.5	
		P su	ipply ha	alves 50	•	ort than initia	
Cell	P		Plas	ticity @	106	212:16:1	
Exports 10	6 mole	s C).5	
					No reductio	on in export	
							/

Measured — Community Small Plankton Large Plan	ton

Prediction of POM stoichiometry by a

global ocean model enable with the new

stoichiometry model (Equation 3) under

steady state.(a-d) Modeled C:P ratio of

aggregate POM, small plankton, large

plankton, and zonal mean. (f-h) Modeled

N:P calculated by dividing C:P with a

fixed C:N value of 7.06.

Reference

- Galbraith, E. D., and A. C. Martiny (2015), A simple nutrientdependence mechanism for predicting the stoichiometry of marine ecosystems., Proc. Natl. Acad. Sci. U. S. A., 112(27), 8199-204.
- Martiny, A. C., A. Talarmin, C. Mouginot, J. A. Lee, J. S. Huang, A. G. Gellene, and D. A. Caron (2016), Biogeochemical interactions control a temporal succession in the elemental composition of marine communities, Limnol. Oceanogr., 61(2), 531–542, doi:10.1002/lno.10233.
- Morel, F. M. M. (1987), Kinetics of Nutrient Uptake and Growth in Phytoplankton, J. Phycol., 23(1), 137–150, doi:10.1111/j.0022-3646.1987.00137.x.
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