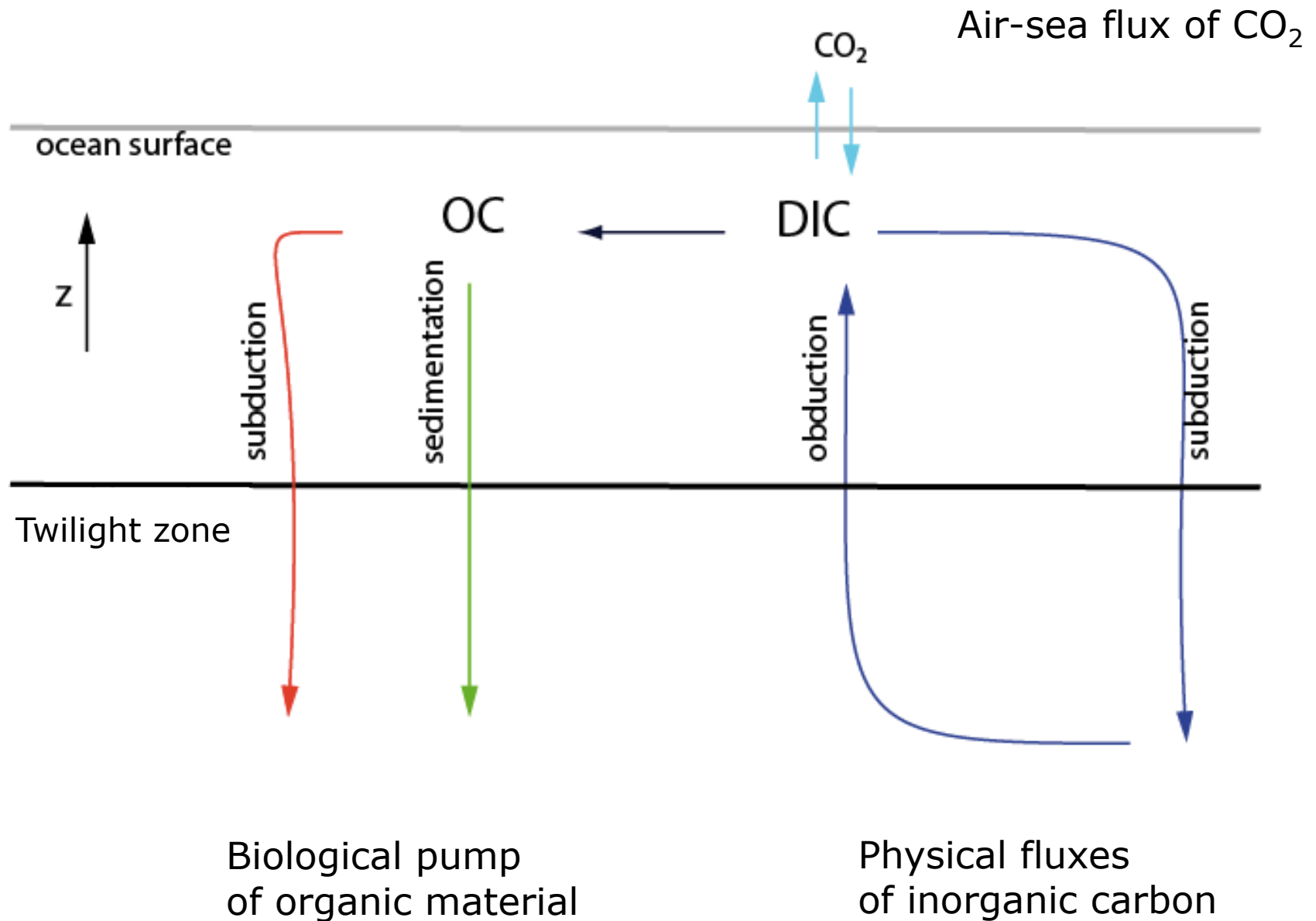


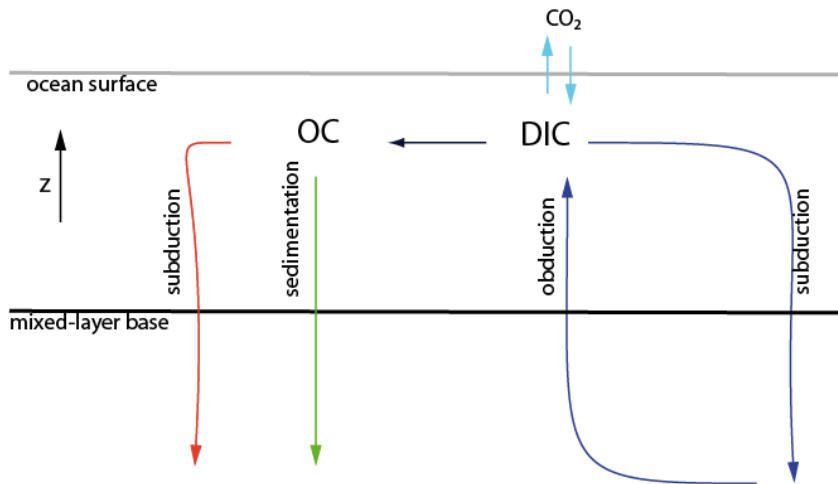


Physical processes and linkages to carbon uptake

Marina Lévy

LOCEAN-IPSL, Paris, France

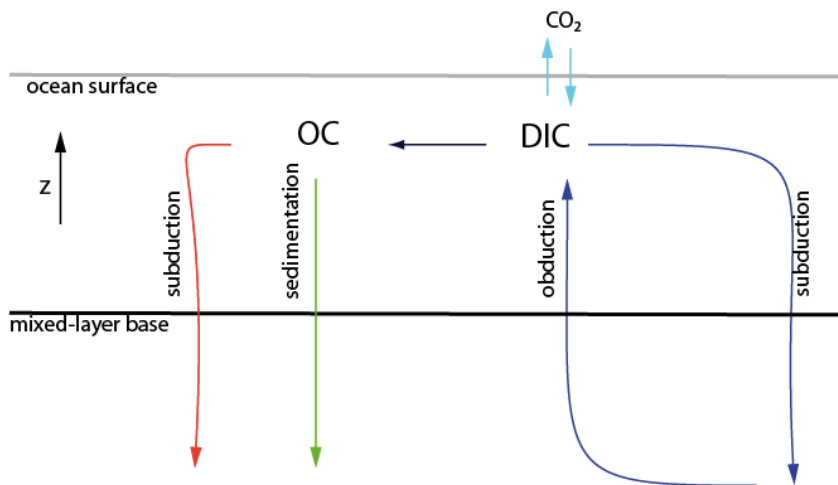




Equilibrium :

Net physical fluxes of DIC = biological pump

Net air-sea flux is close to zero



Equilibrium :
 Net physical fluxes of DIC = biological pump
 Net air-sea flux is close to zero

Are the gross fluxes of DIC much larger than the biological pump ?

How large is subduction of OC compared to sedimentation ?

What are the regional distribution of these carbon fluxes ?



CO₂ concentrations in the atmosphere have increased due to anthropogenic emissions

Ocean has absorbed between 1/3 and 1/4 of this excess carbon emissions

Ocean uptake is saturating

Ocean circulation and productivity have not drastically changed

Oceanic uptake of anthropogenic CO₂ is achieved by physical fluxes:
subduction of DIC



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Oceanic uptake of anthropogenic CO₂ is achieved by physical fluxes:
subduction of DIC

How and where is subduction operating ?

What is the contribution of meso/submesoscale features ?

What will be the effect of a change in circulation ?

What will be the effect of a change in physiology / ecosystem / C:N ratio ?



1) Change of CO₂ in atmosphere

Unperturbed circulation, unperturbed biological pump

2) Change in ocean circulation

-> change in N supply

-> change in biological pump

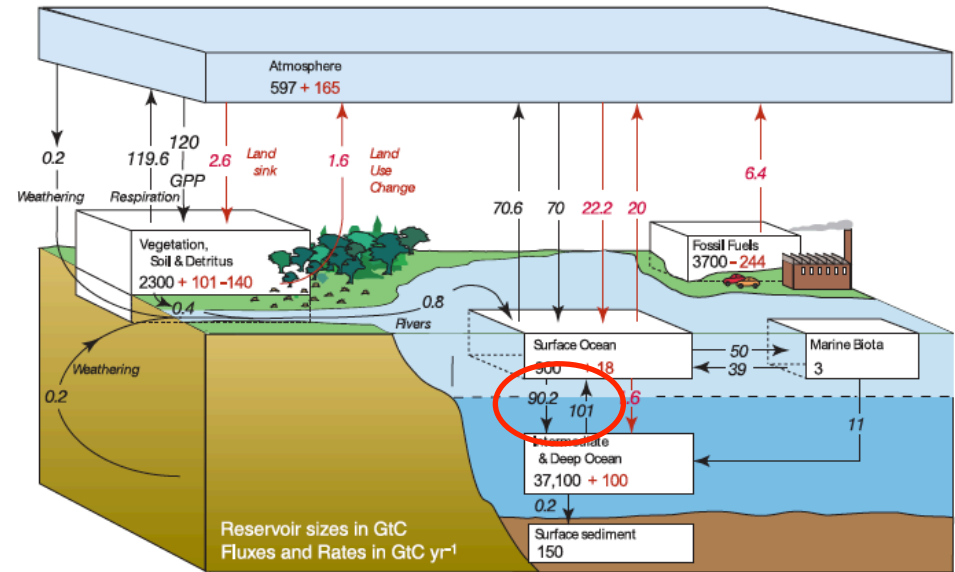
Unperturbed circulation, perturbed biological pump

3) Change in physiology, ecosystem structure, C/N ratio

-> change in biological pump

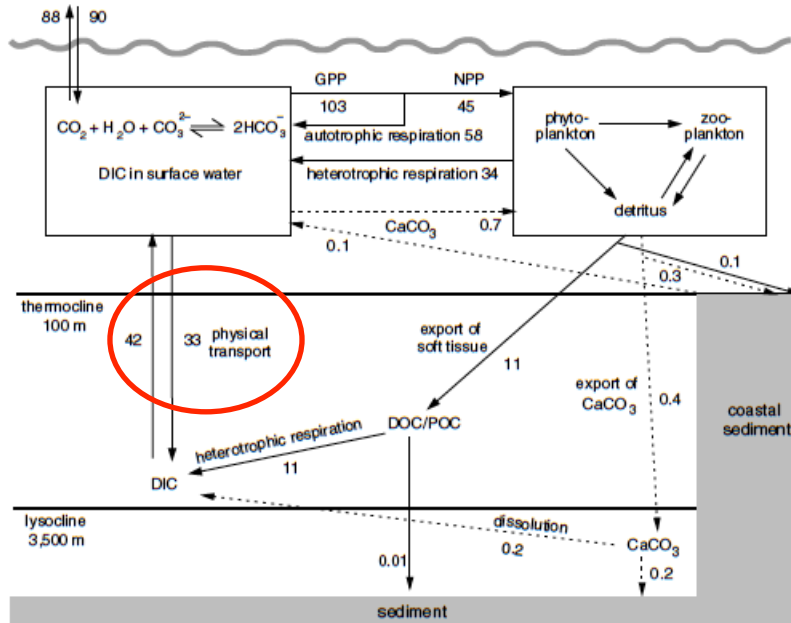
Perturbed circulation, perturbed biological pump

How will these different changes modify the uptake in the future ?

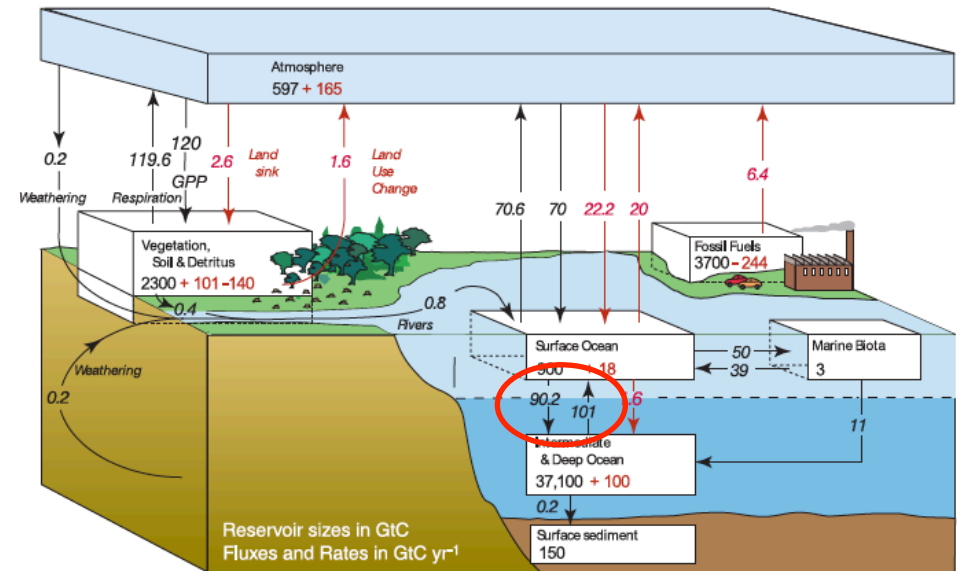


IPCC, 2007, 2013

Carbon cycling in the ocean

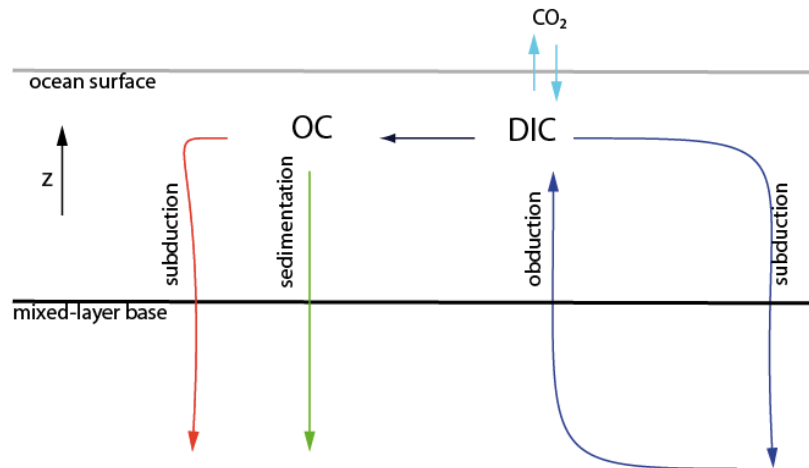


IPCC, 2001, 2011



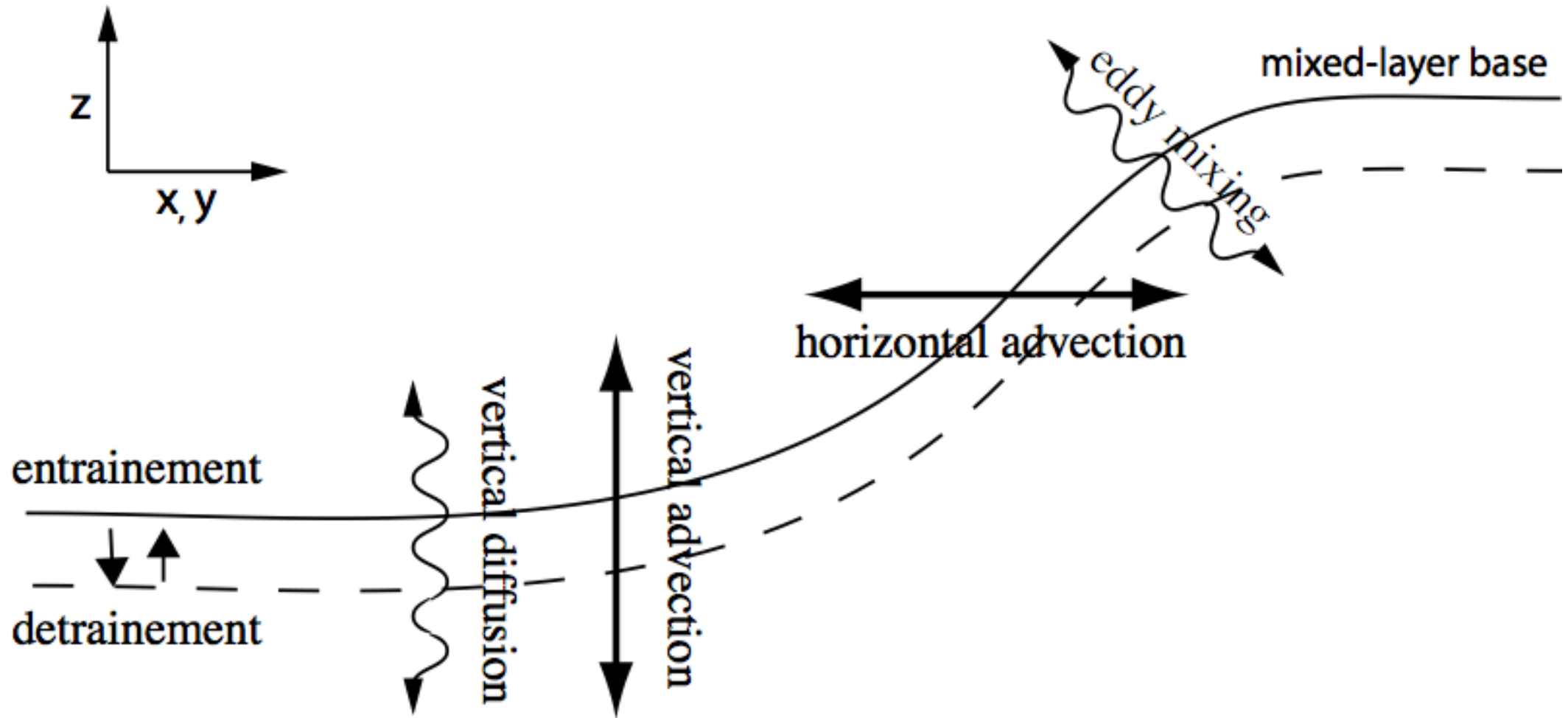
IPCC, 2007, 2013

Besides for a few regional studies, very little is known on the **physical transport** between surface and deep ocean



Meaningful boundary for biological pump: [euphotic depth](#)

Meaningful boundary for physical fluxes: surface boundary layer
Or [mixed-layer depth](#)





Instantaneous flux across
the time varying mixed-layer depth

$$S_{\text{ann}}^{\text{TR}} = \int_0^{365} \left(-c_h \cdot w_h - c_h \cdot \vec{u}_h \cdot \vec{\nabla} h - c_h \cdot \frac{\partial h}{\partial t} - A_z^h \cdot \frac{\partial c_h}{\partial z} \right) dt,$$

Annual Subduction / Obduction rate



Instantaneous flux across
the time varying mixed-layer depth

$$S_{\text{ann}}^{\text{TR}} = \int_0^{365} \left(-c_h \cdot w_h - c_h \cdot \vec{u}_h \nabla h - c_h \cdot \frac{\partial h}{\partial t} - A_z^h \cdot \frac{\partial c_h}{\partial z} \right) dt,$$

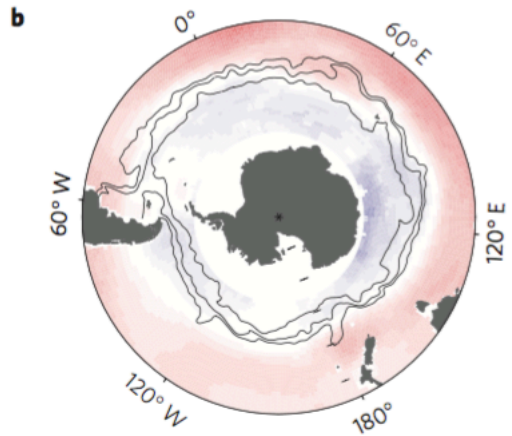
Annual Subduction / Obduction rate

Extremely difficult to estimate from observations

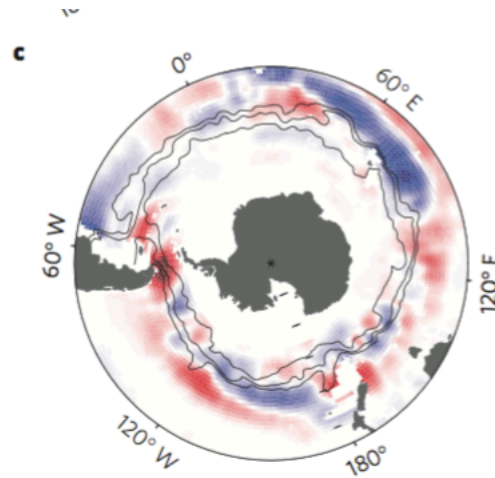
Models offer a framework where all the fields are consistent
and the fluxes can be estimated

Flux across the winter mixed-layer depth: estimates based on observations

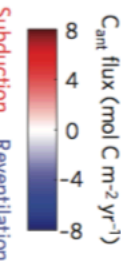
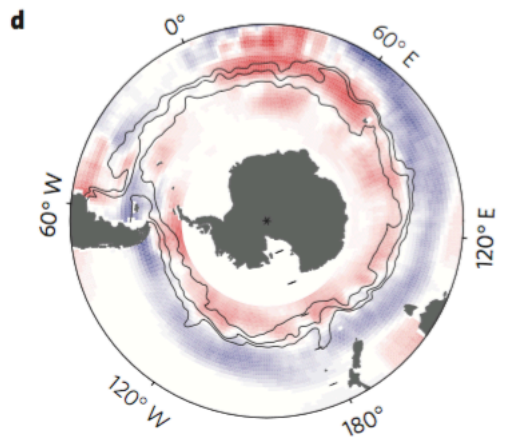
Vertical



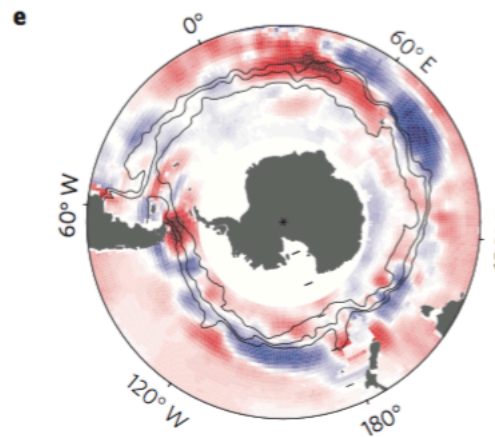
Horizontal



Eddy



Total





1. Pre-industrial carbon transfers in the global ocean (Levy et al., GBC, 2013)
2. Anthropogenic carbon penetration in the global ocean (Bopp et al., GRL, 2015)
3. Role of (sub-)mesoscales in the North East Atlantic (Karleskind et al., JGR 2011a, Oce. Mod. 2011b)

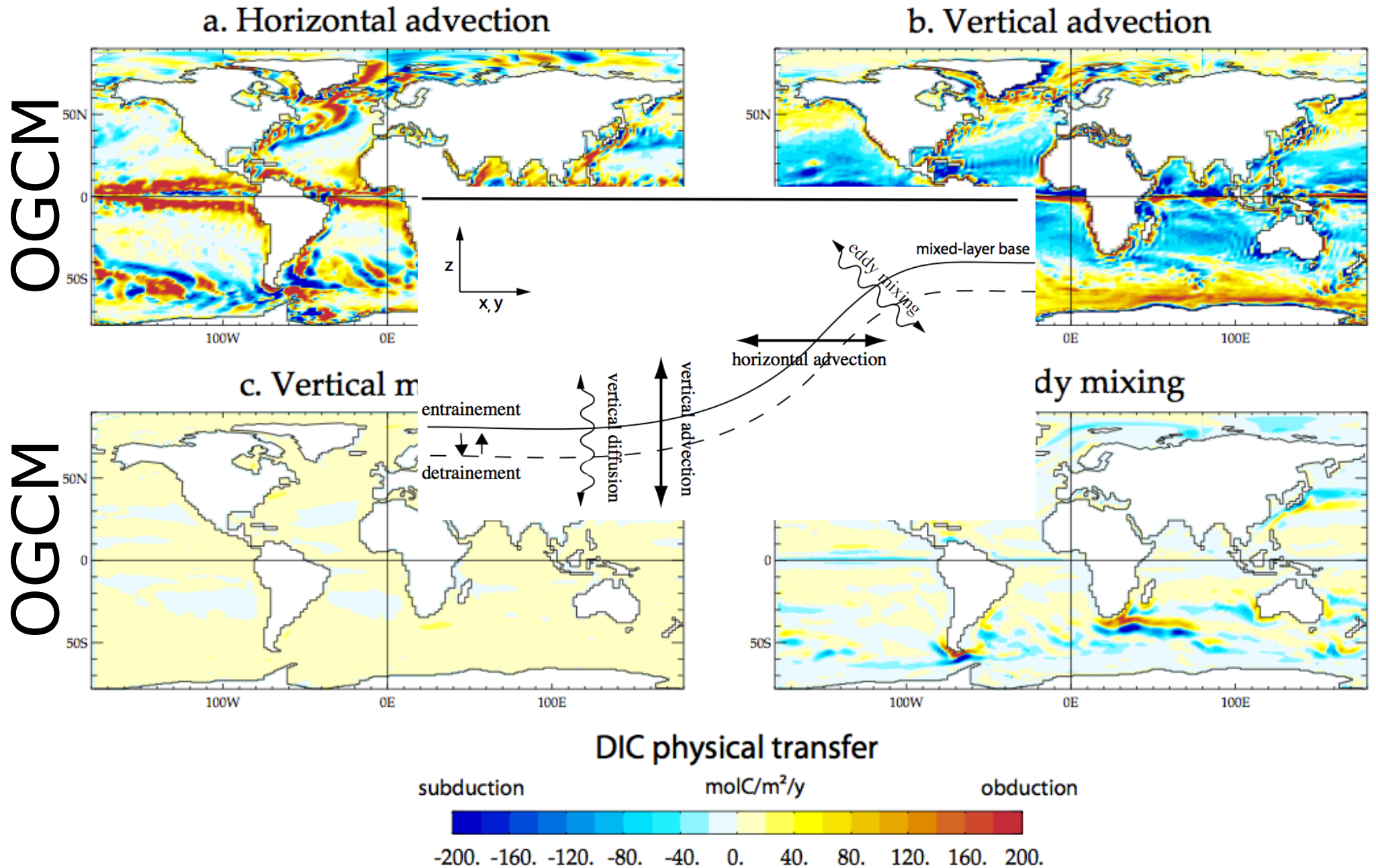


1. OGCM: ORCA2-PISCES

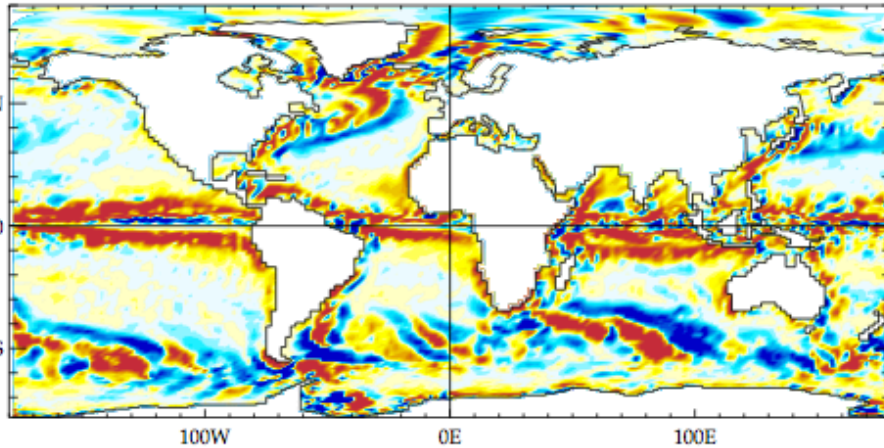
climatological simulation **for the pre-industrial ocean**

2. Global climatological data

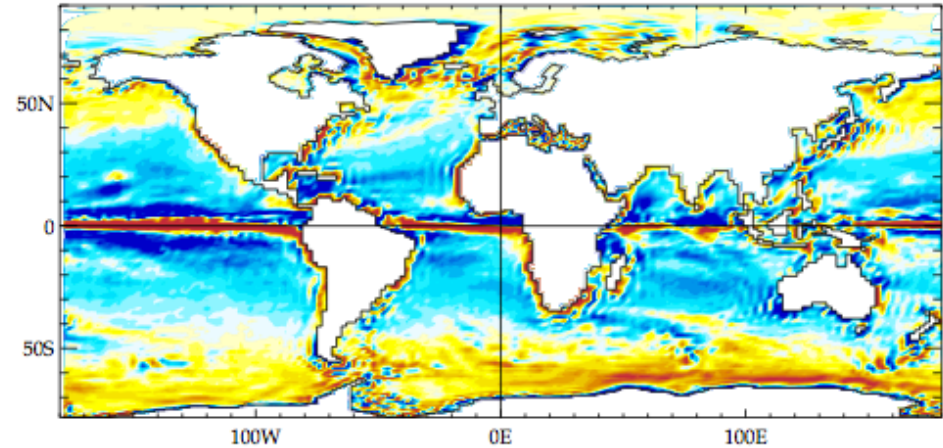
DIC (GLODAP), Wind (CORE2), U (AVISO), MLD (De Boyer Montegut)



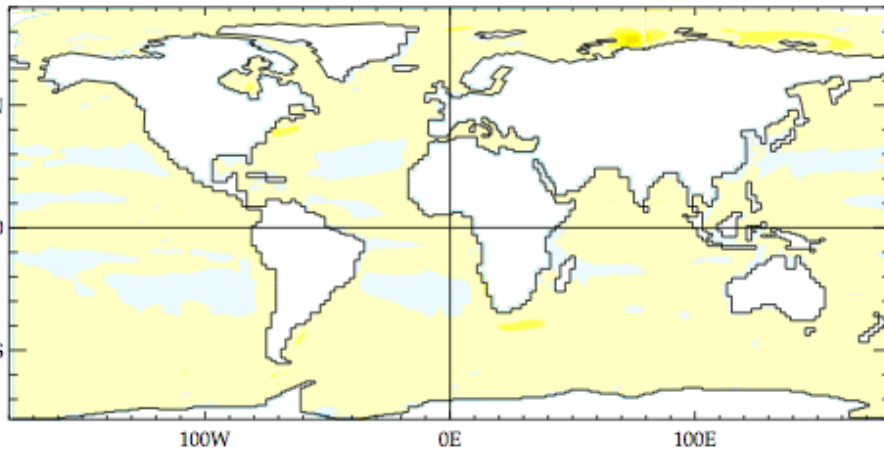
a. Horizontal advection



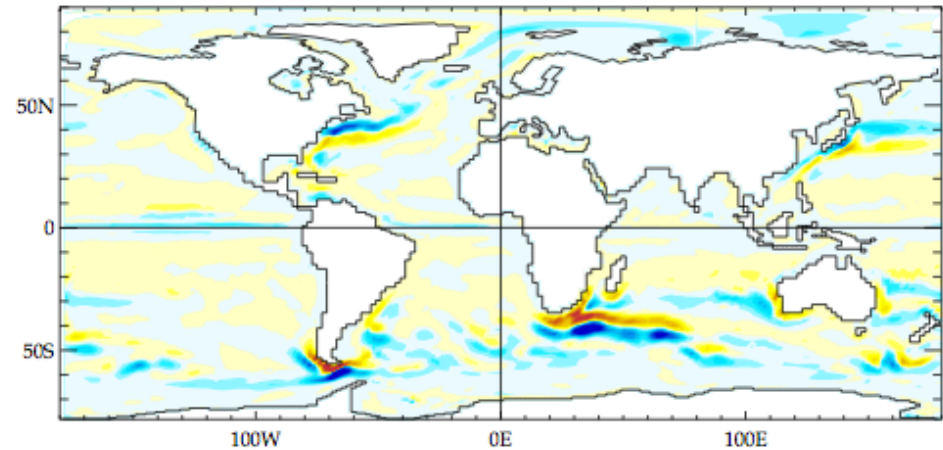
b. Vertical advection



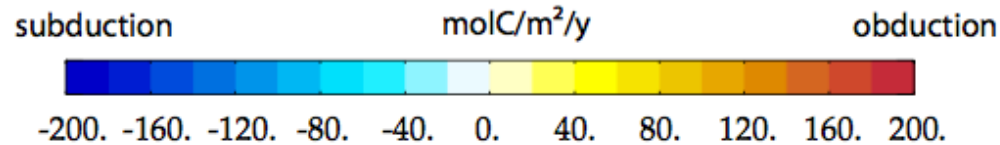
c. Vertical mixing



d. Eddy mixing



DIC physical transfer

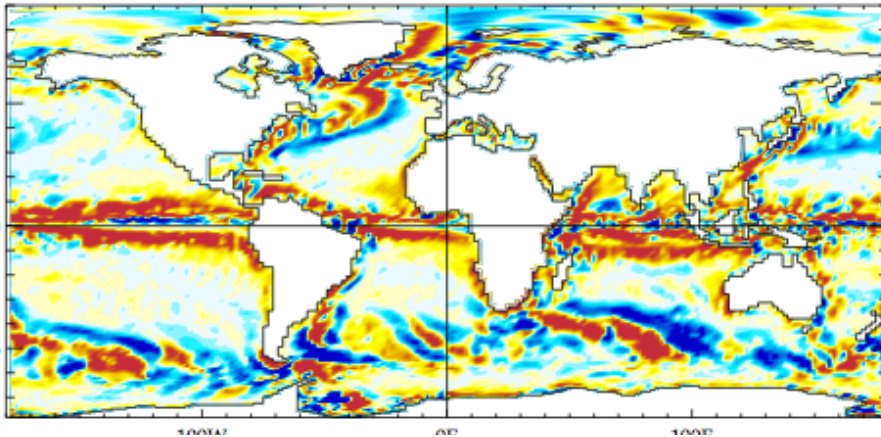


OGCM

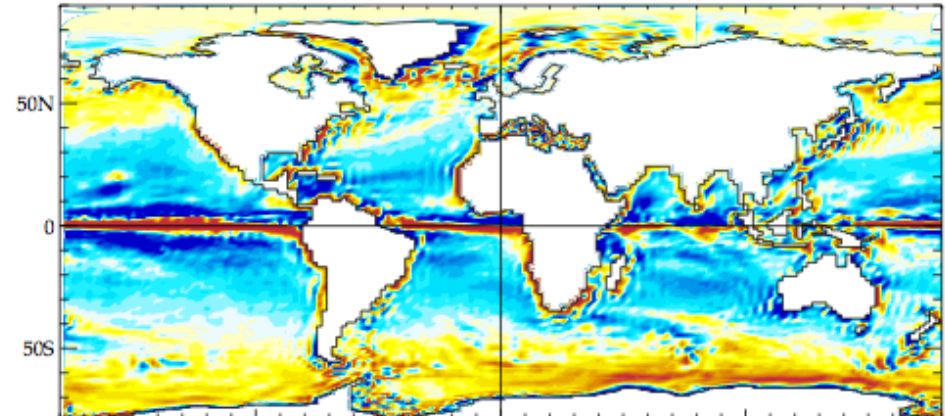
OGCM

OGCM

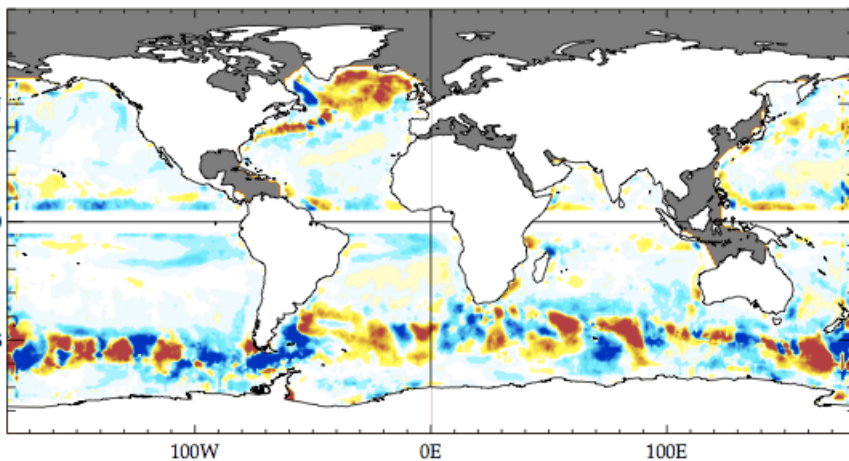
a. Horizontal advection



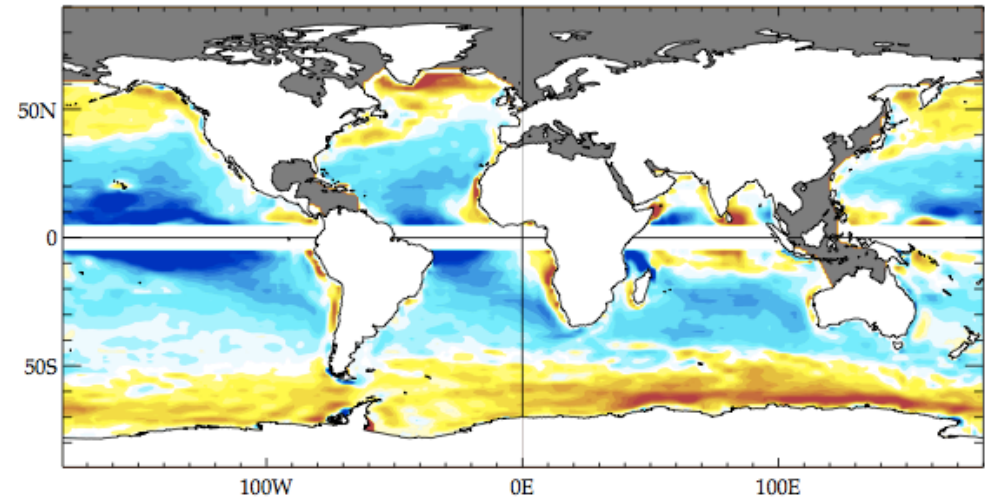
b. Vertical advection



b1. Horizontal Advection (from data)



c1. Vertical Advection (from data)

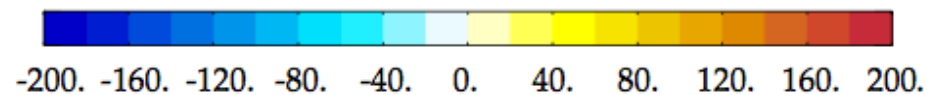


DIC physical transfer

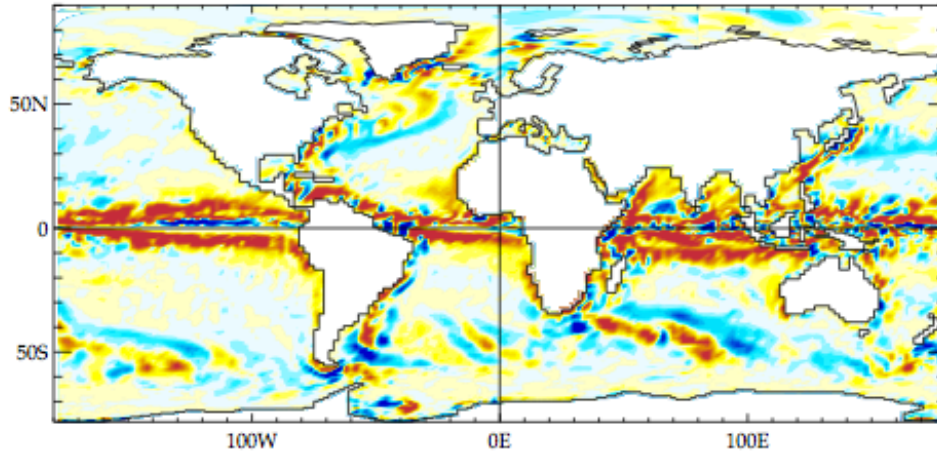
subduction

molC/m²/y

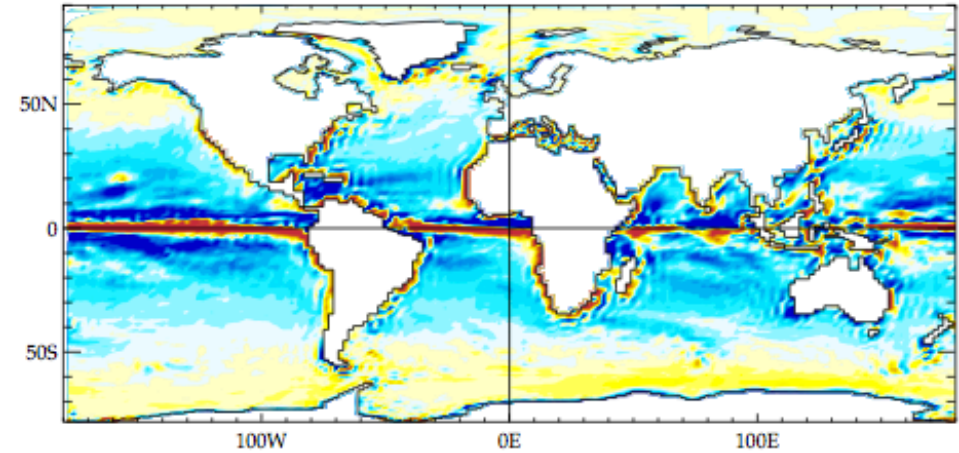
obduction



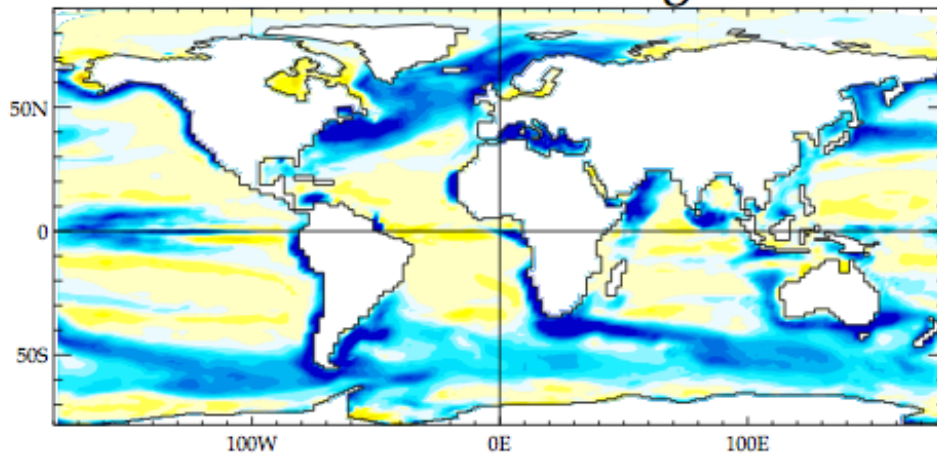
a. Horizontal Advection



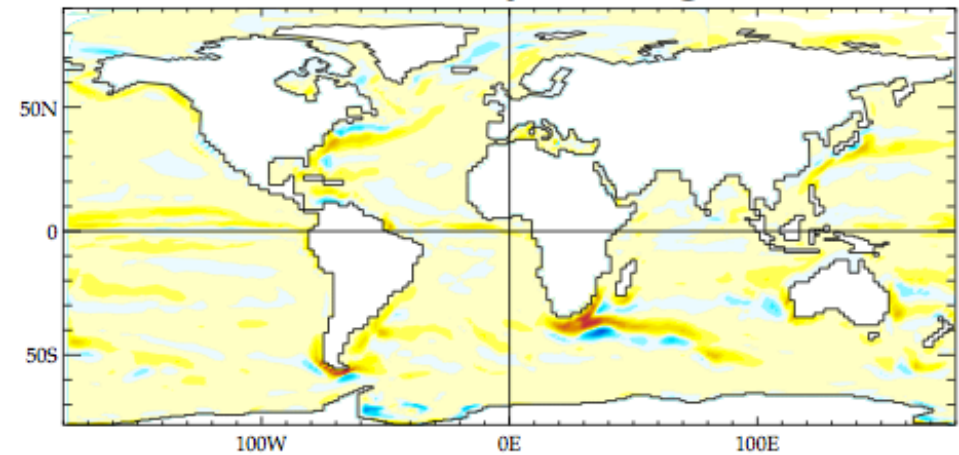
b. Vertical Advection



c. Vertical mixing

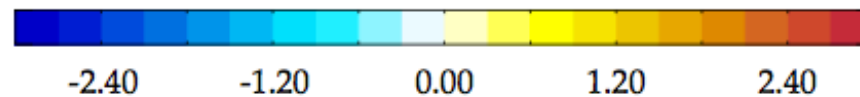


d. Eddy mixing



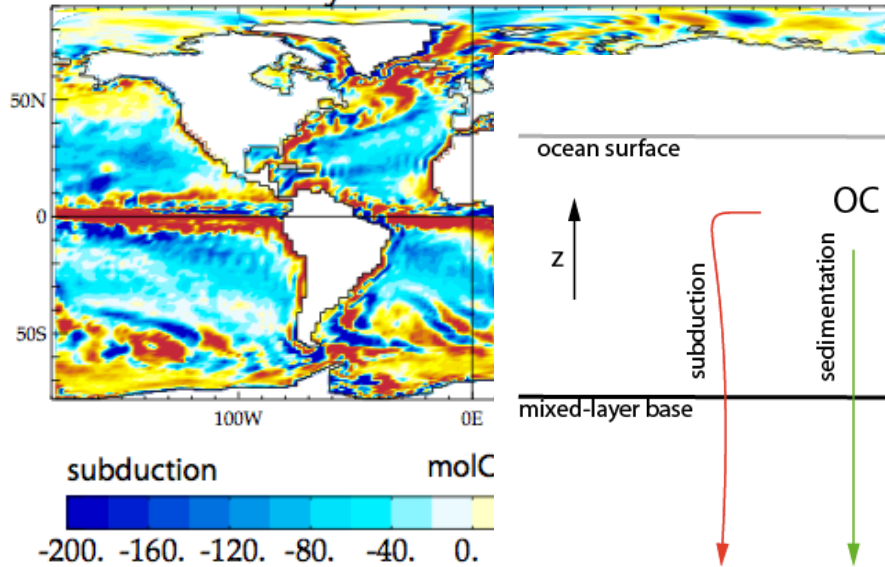
Organic Carbon physical transfer ($\text{molC}/\text{m}^2/\text{y}$)

Subduction

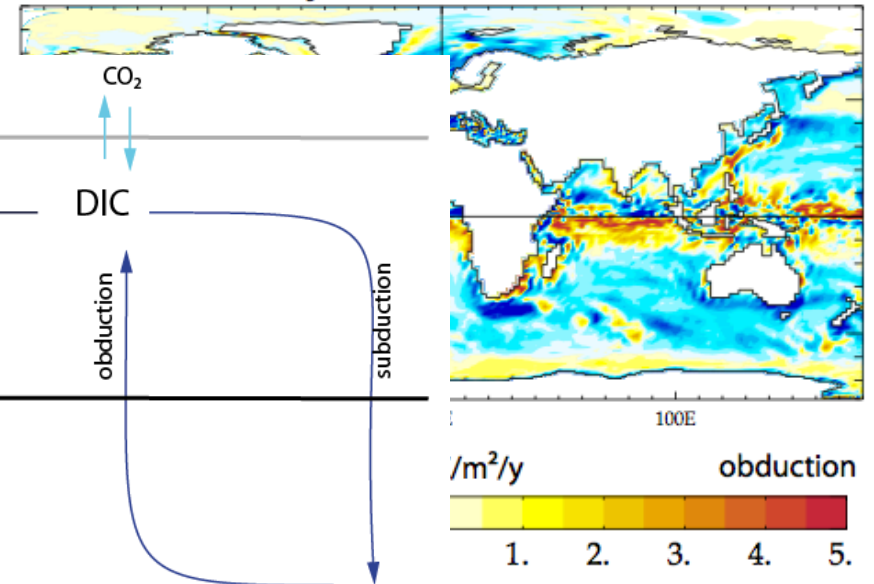


Obduction

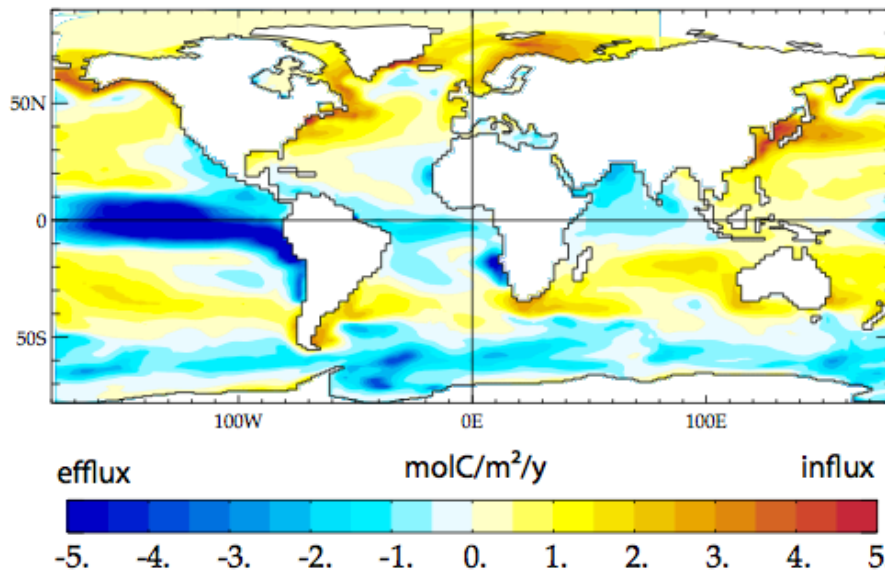
a. Physical Transfer of DIC



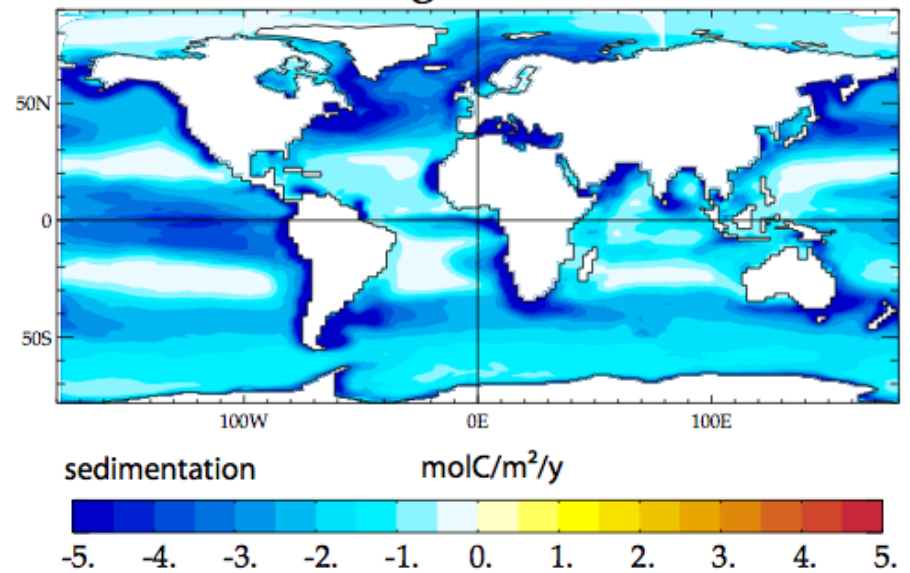
b. Physical Transfer of OC



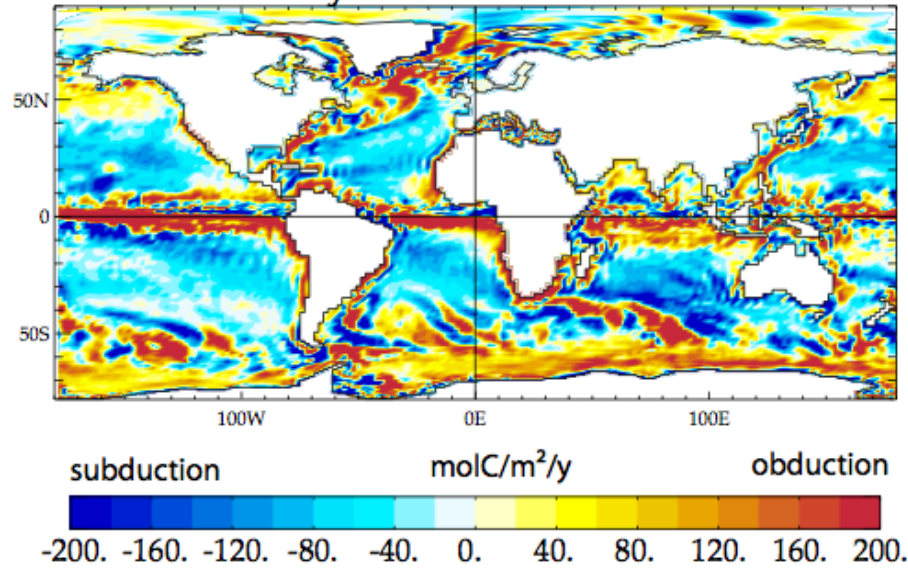
c. air-sea CO₂ flux



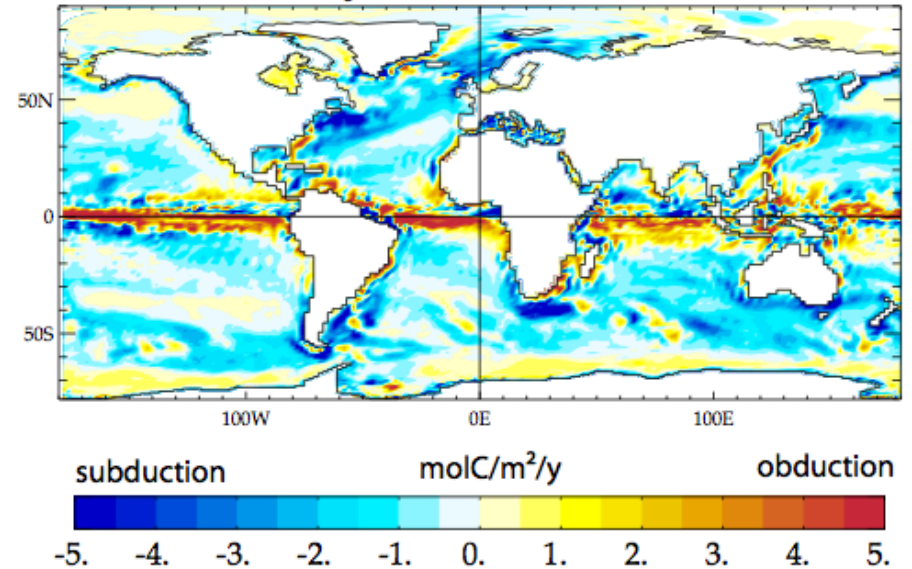
d. Biological Transfer



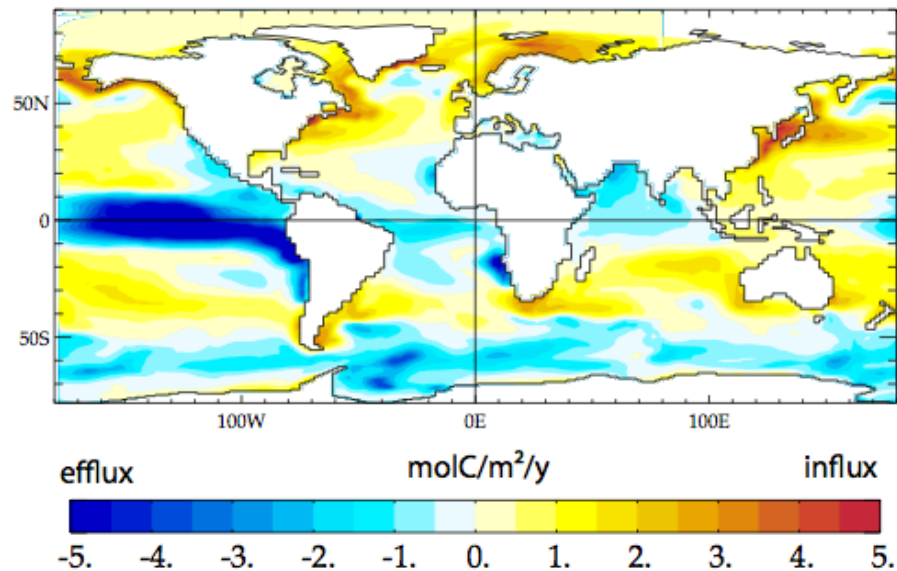
a. Physical Transfer of DIC



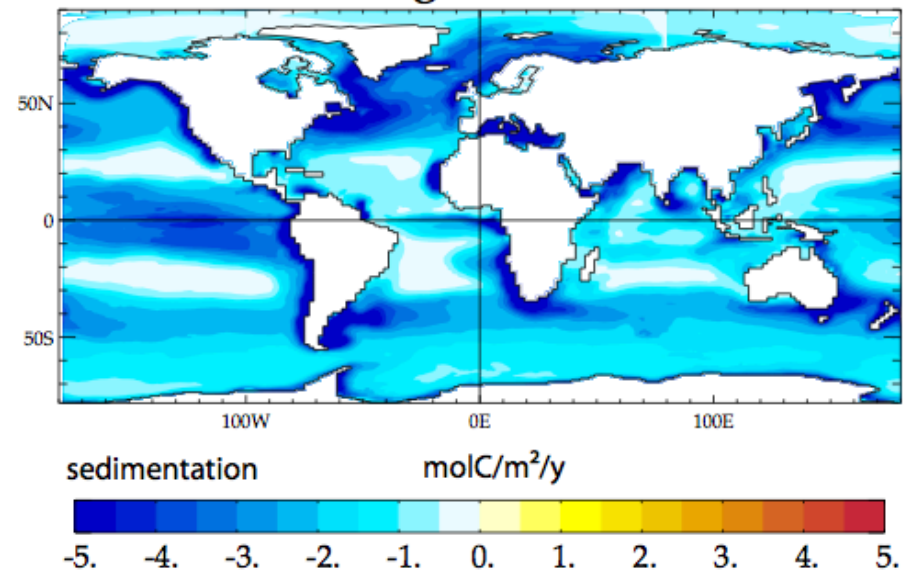
b. Physical Transfer of OC



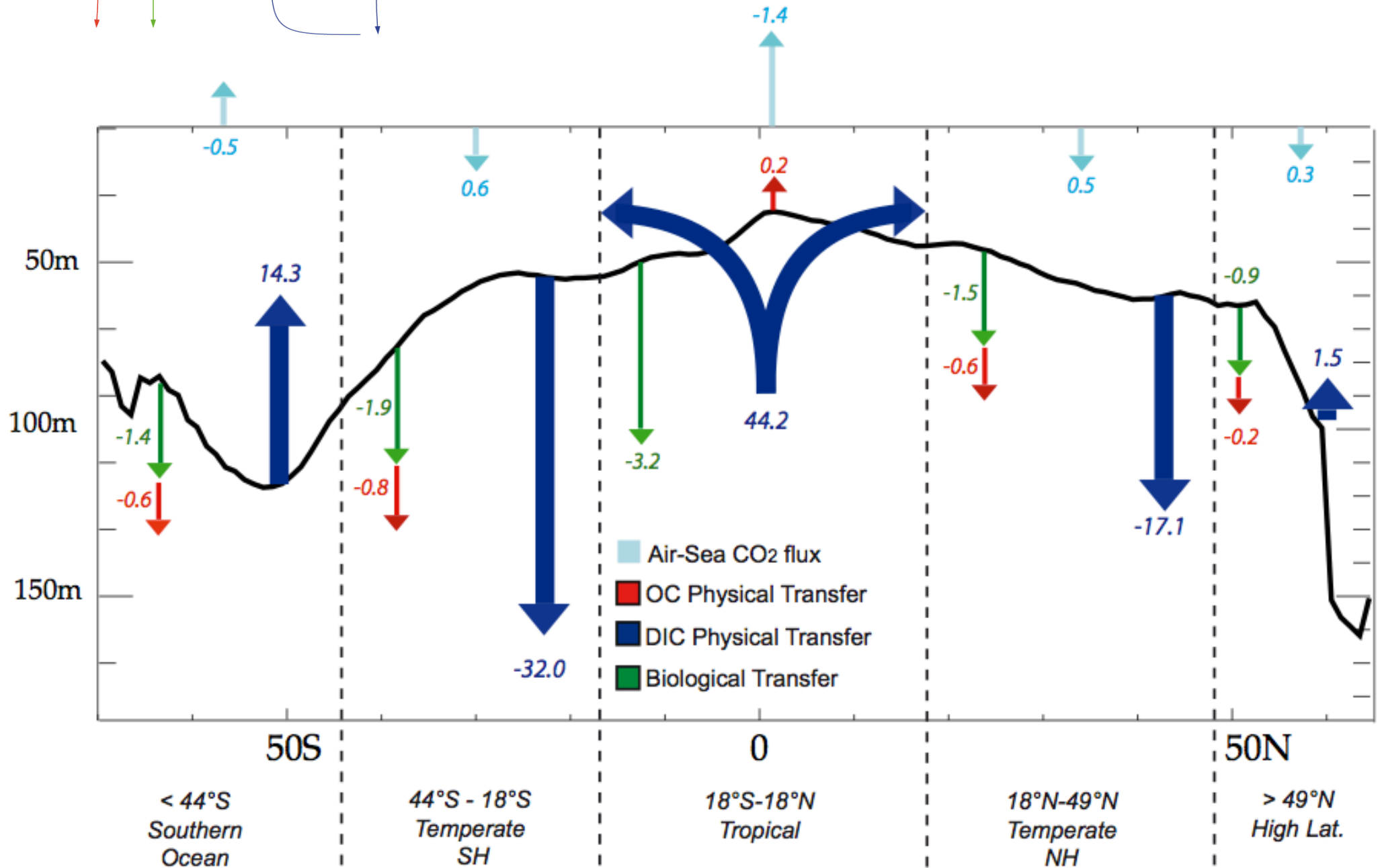
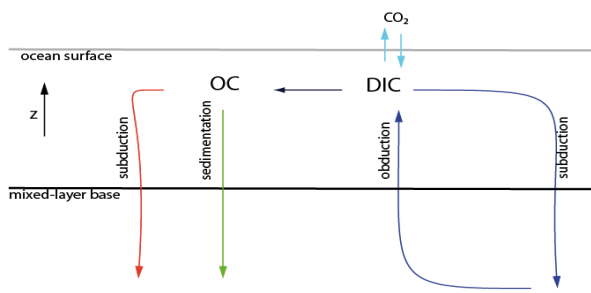
c. air-sea CO₂ flux

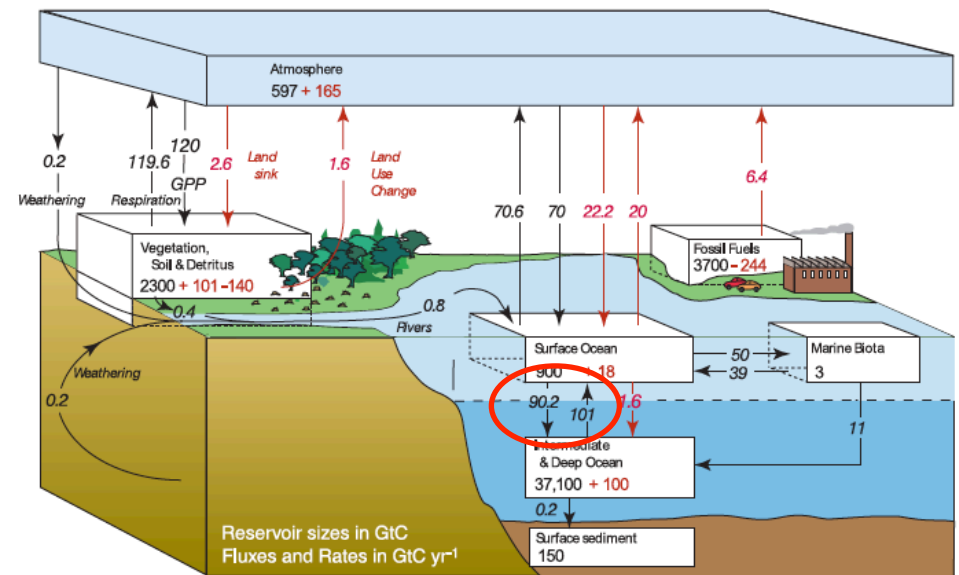
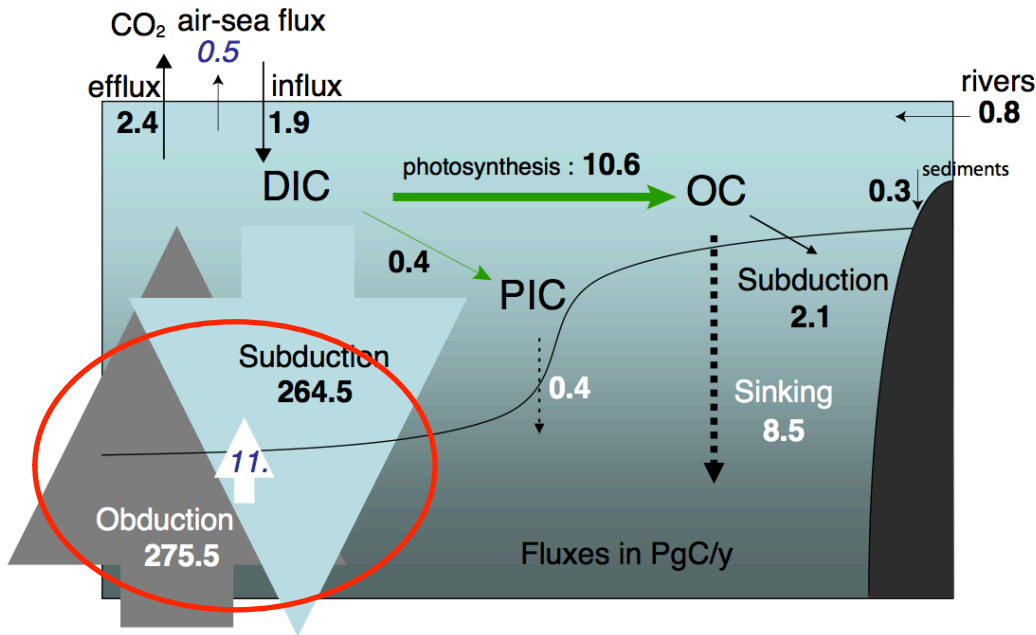


d. Biological Transfer



Carbon fluxes through the ML







For antropogenic carbon flux estimates

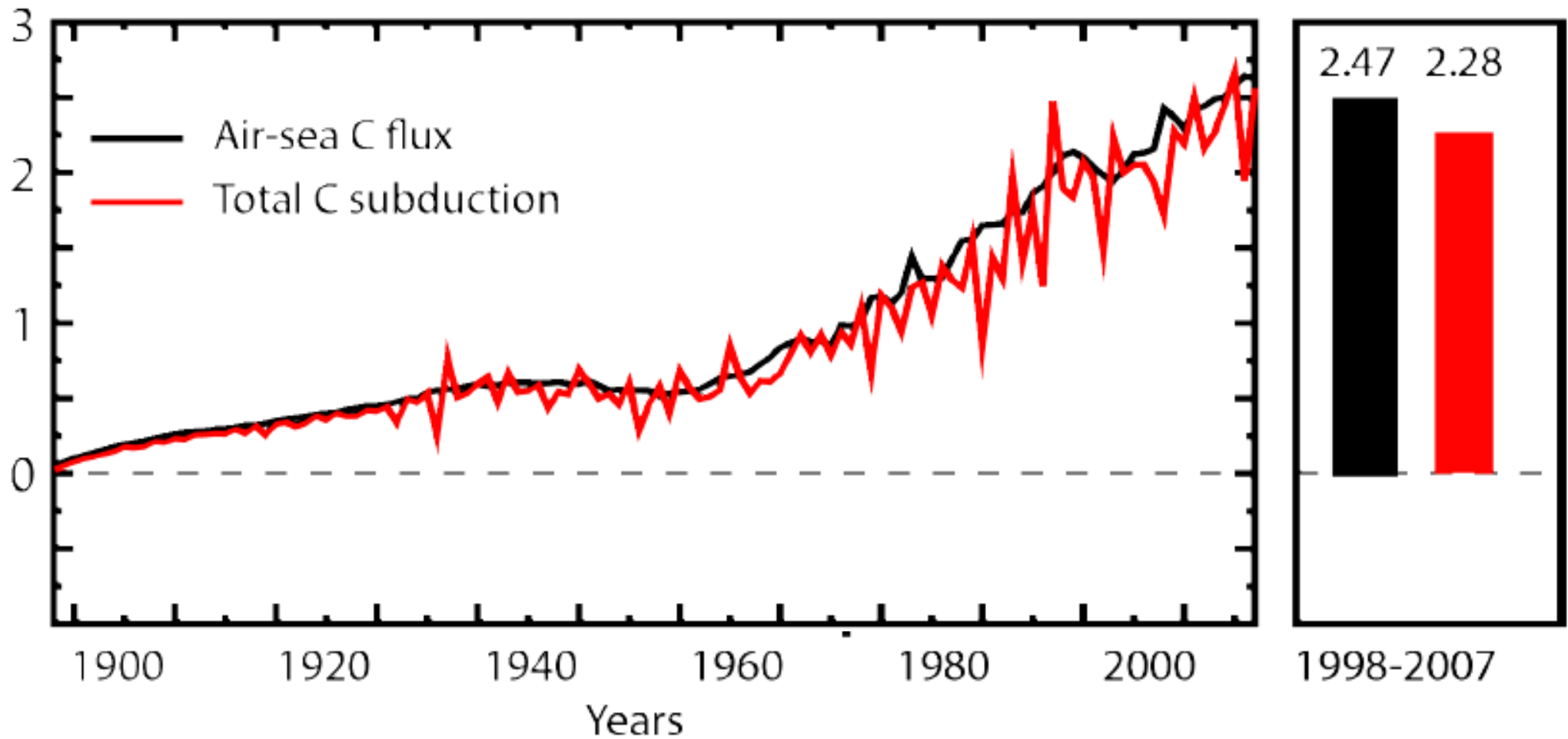
OGCM: ORCA2-PISCES

Historical simulation (1890-2007) with / without trend in atm pCO₂

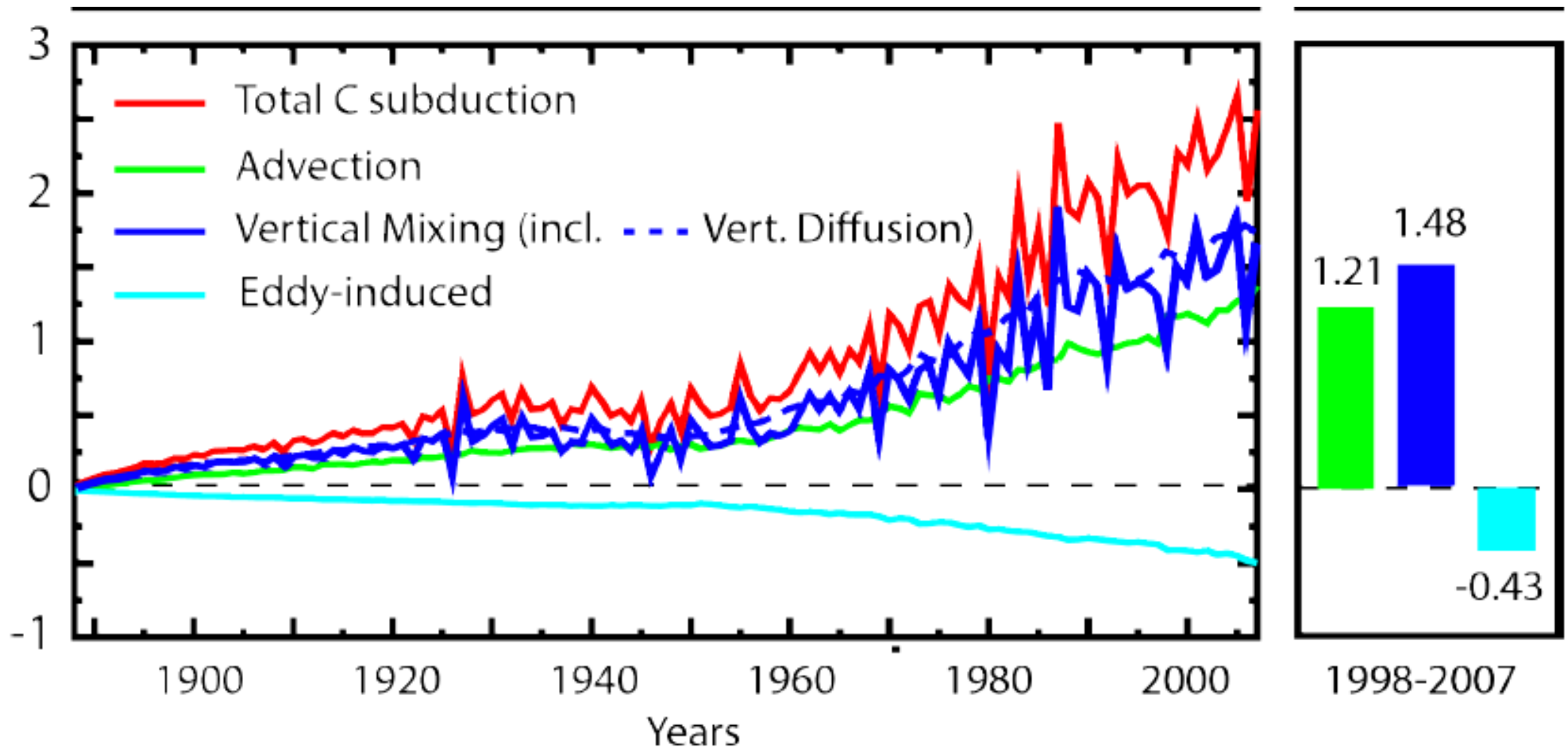
The circulation and biological pump are identical in the two simulations

We look at the difference between the two to access anthropogenic carbon

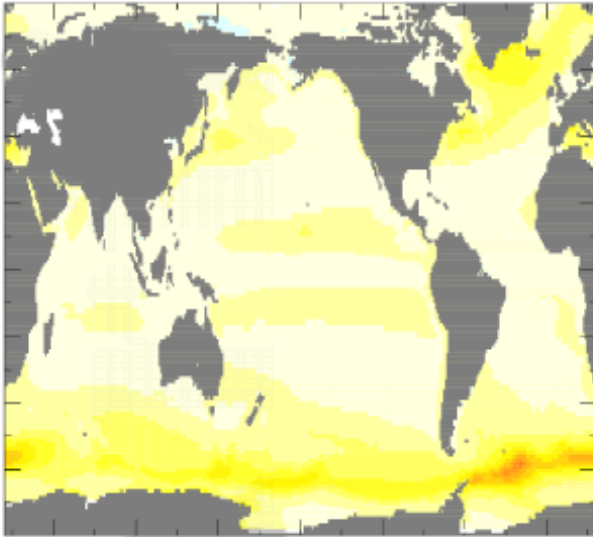
PgC / y



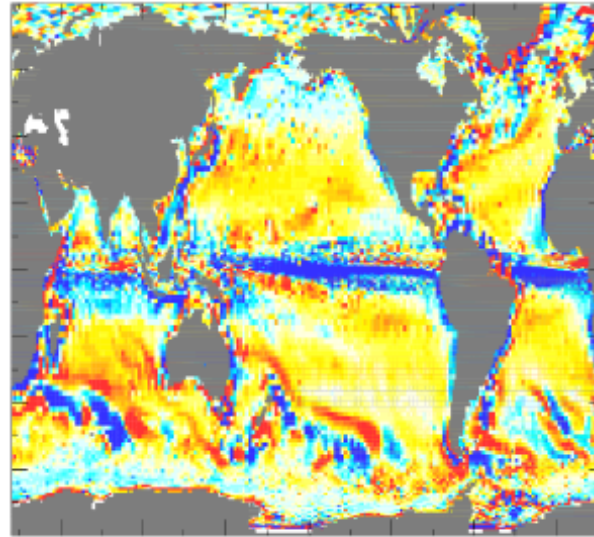
larger interannual variability in carbon subduction than in air-sea carbon flux



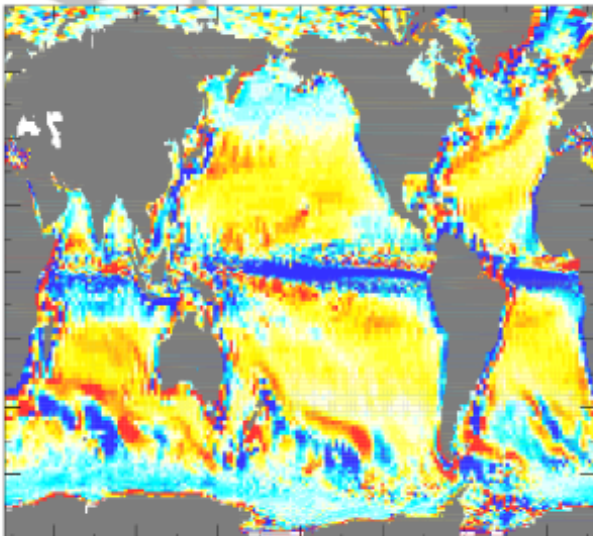
a. Air -Sea C Flux



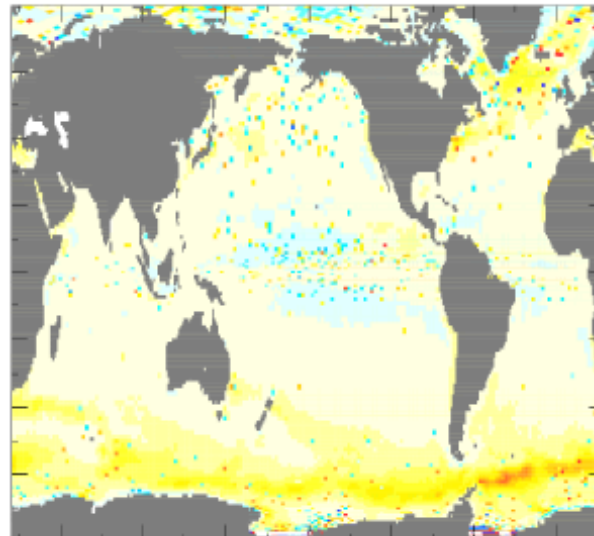
b. Total C Subduction



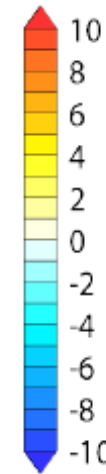
c. Advection



d. Vertical Mixing



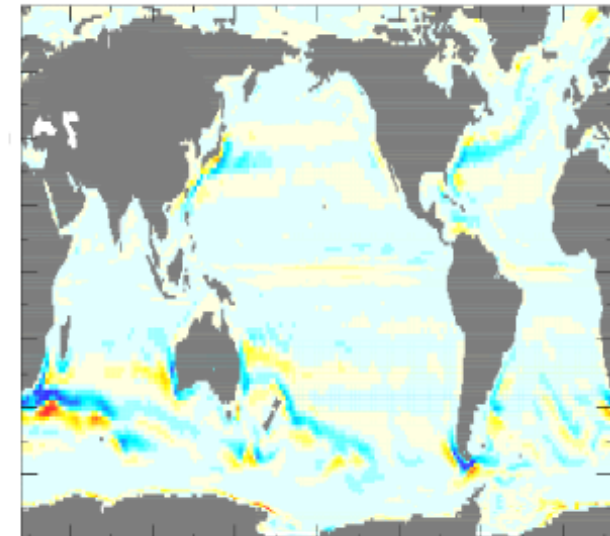
$\text{gC m}^{-2} \text{y}^{-1}$



↑
From air to sea From surface to deep

Canth air-sea flux and Canth penetration do not occur at the same location

e. Eddy-induced





For antropogenic carbon flux estimates

OGCM: ORCA2-PISCES

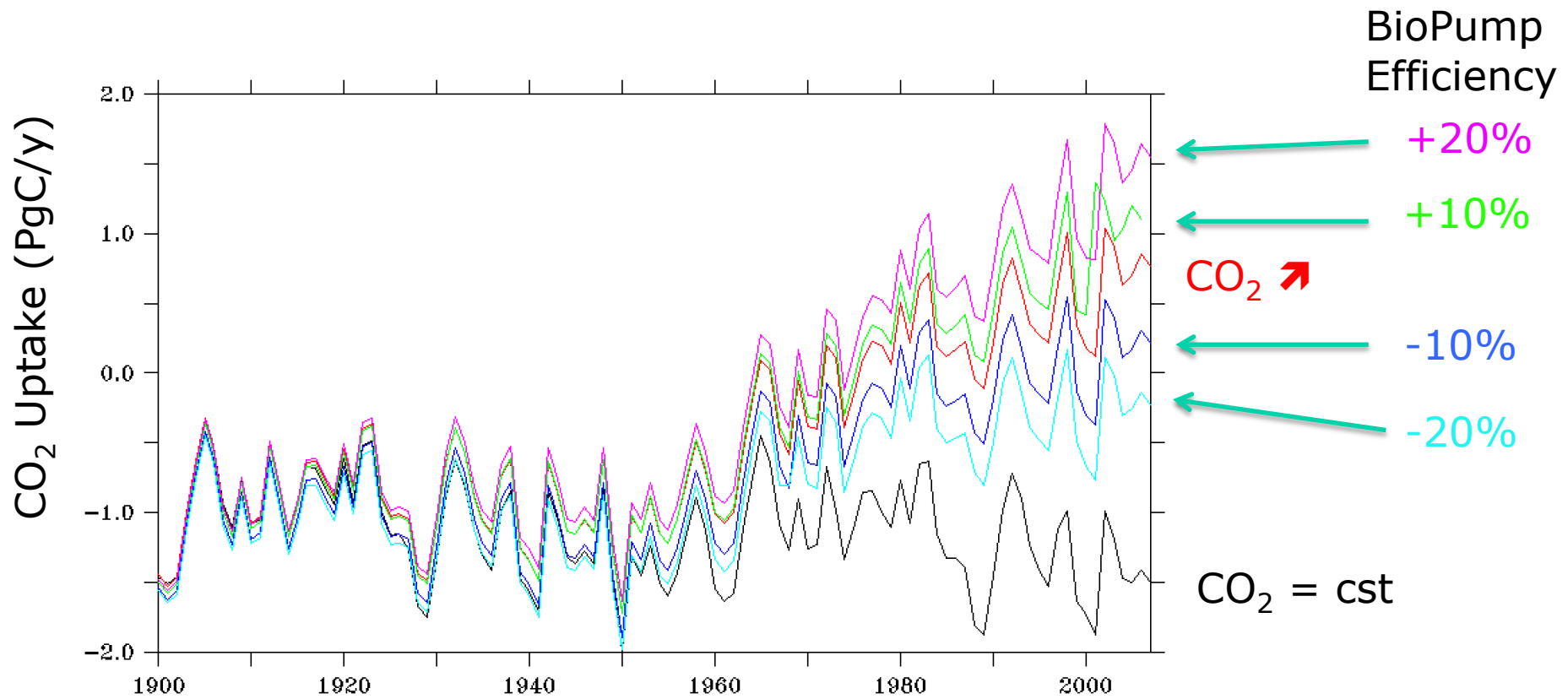
Historical simulation (1890-2007) with trend in atm pCO₂

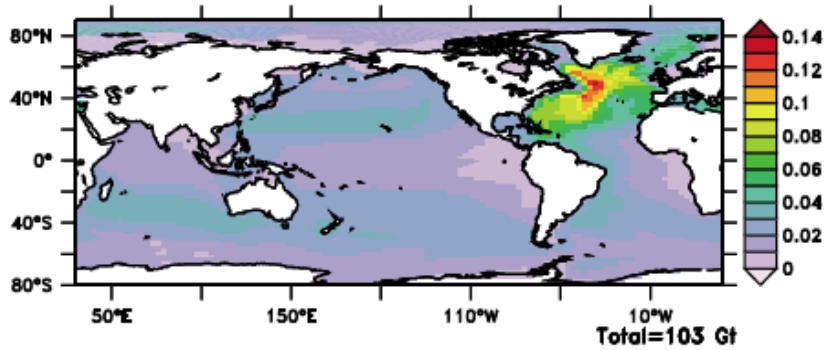
Efficiency of the biological pump is controlled by changing the C:N ratio of PP

$$C/N = 7.625 (1 + x \cdot (CO_2^t - CO_2^{1888}) / (CO_2^{2007} - CO_2^{1888}))$$

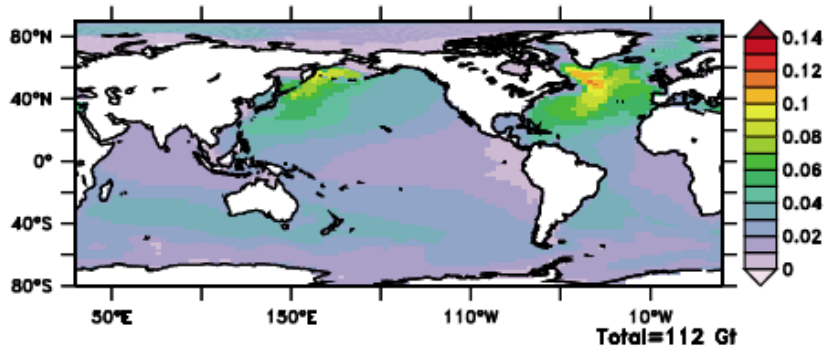
With $x = 0.2, 0.1, -0.1, -0.2$

Circulation is the same for all experiments

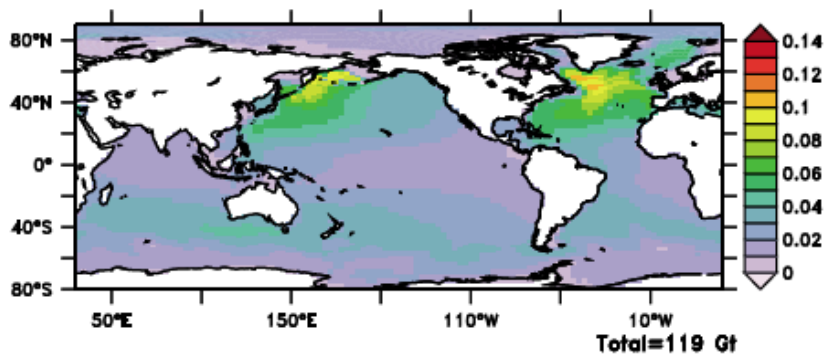




(a) AREDI400, Anthro DIC 1995

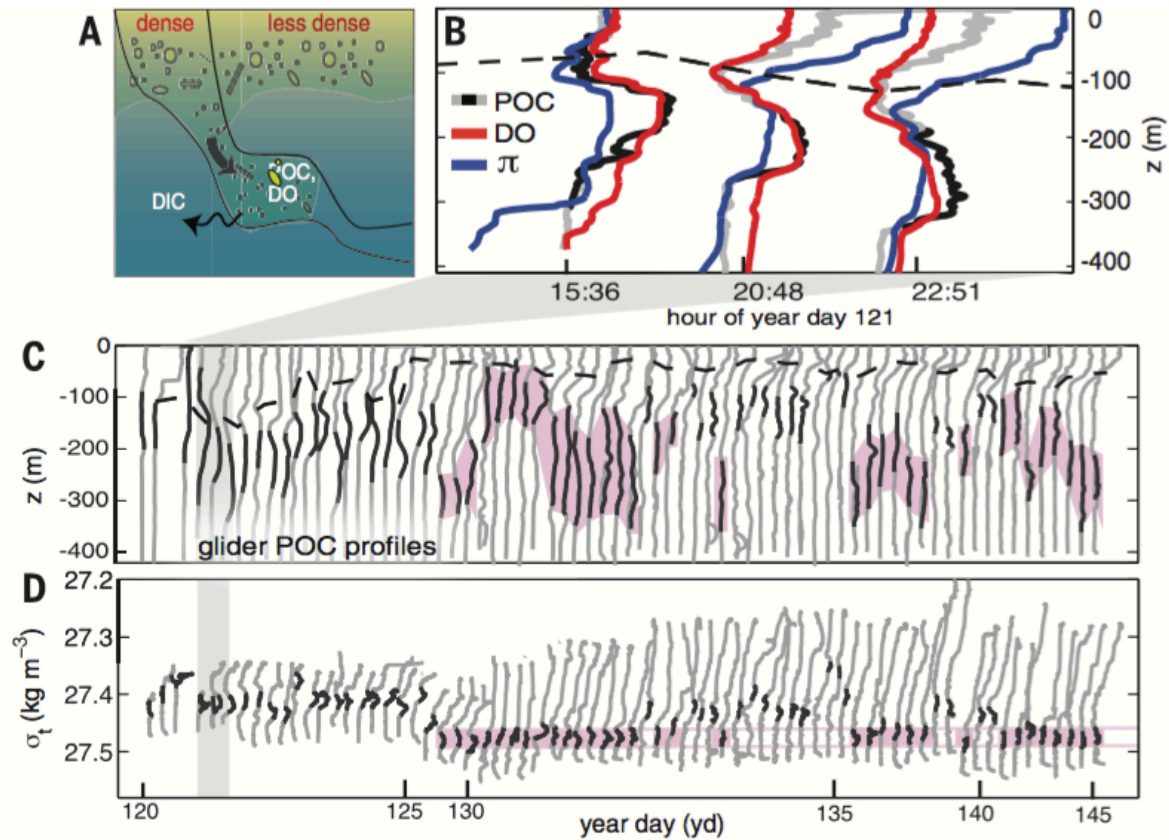


(c) ABER2D, Anthro DIC 1995



(e) AREDI2400, Anthro DIC 1995

Sensitivity of Canth upatke
to lateral eddy mixing



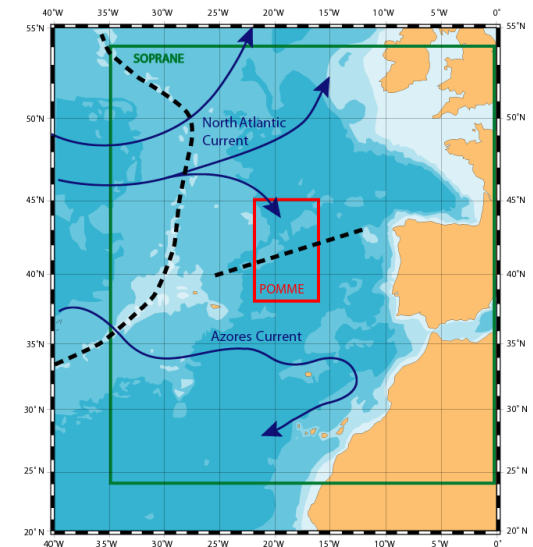
Subduction of POC and DOC

Program Ocean Multidisciplinary MEscale

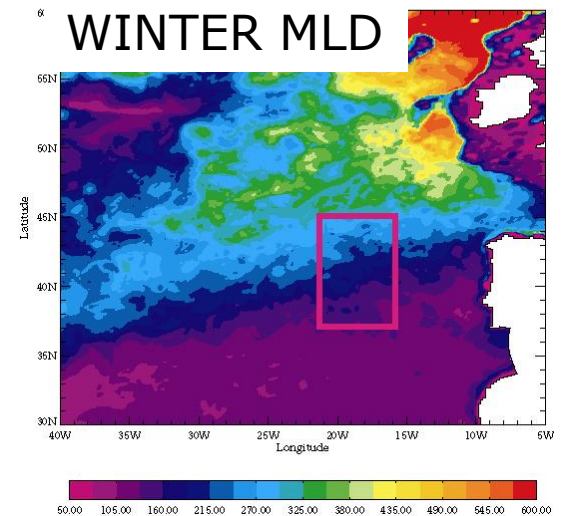
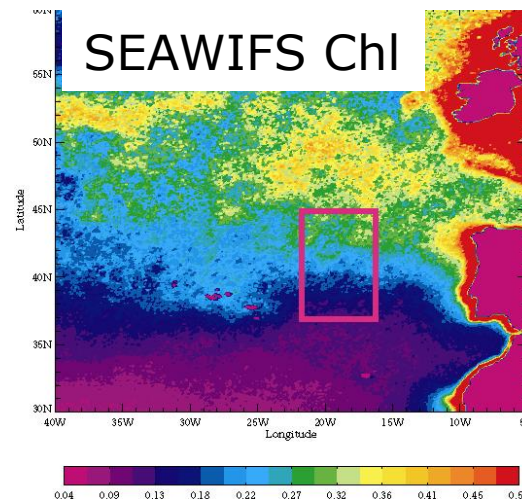
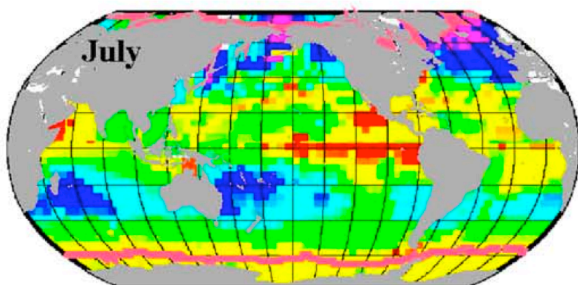
Objective : Role of eddies on subduction of NEAMW and on their biogeochemical properties

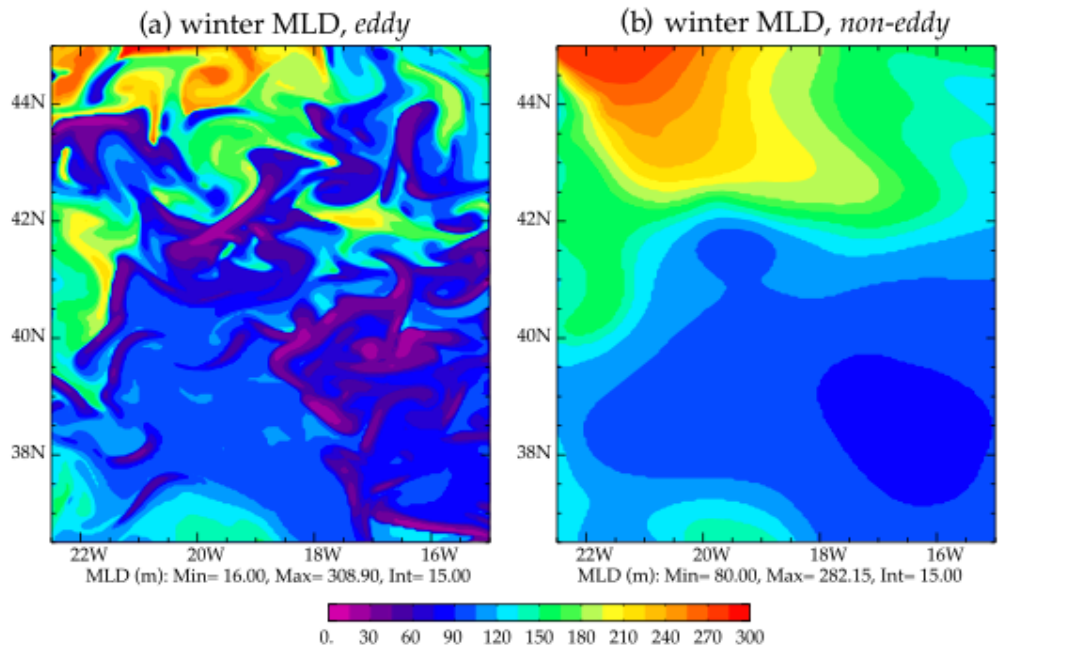
Pomme area

- Strong CO₂ sink
- Mixed-layer depth gradient
- Spring phytoplankton bloom



Delta pCO₂

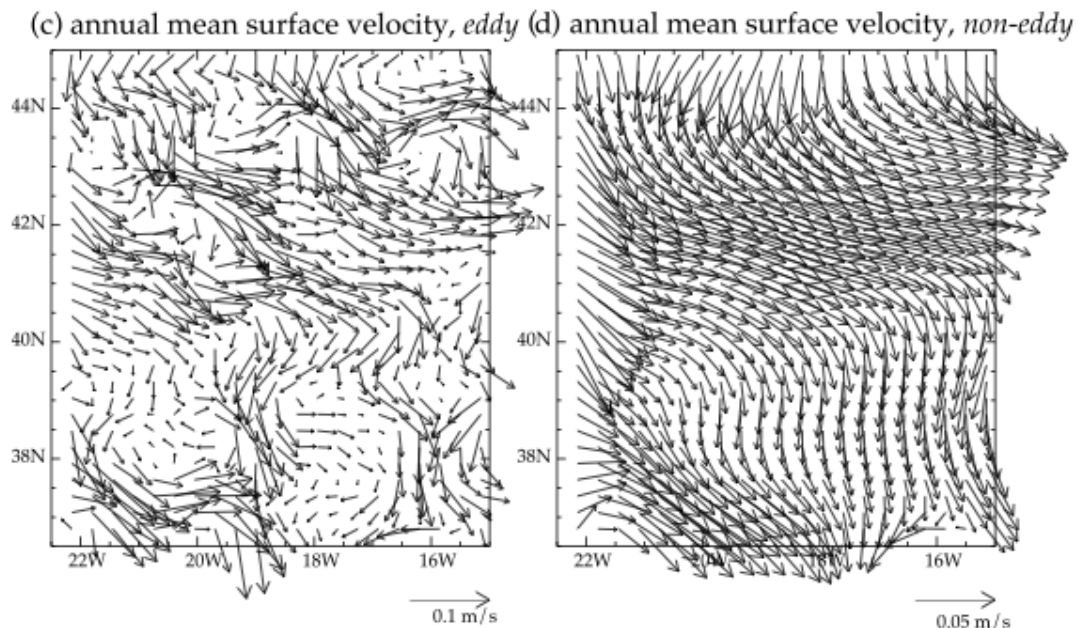




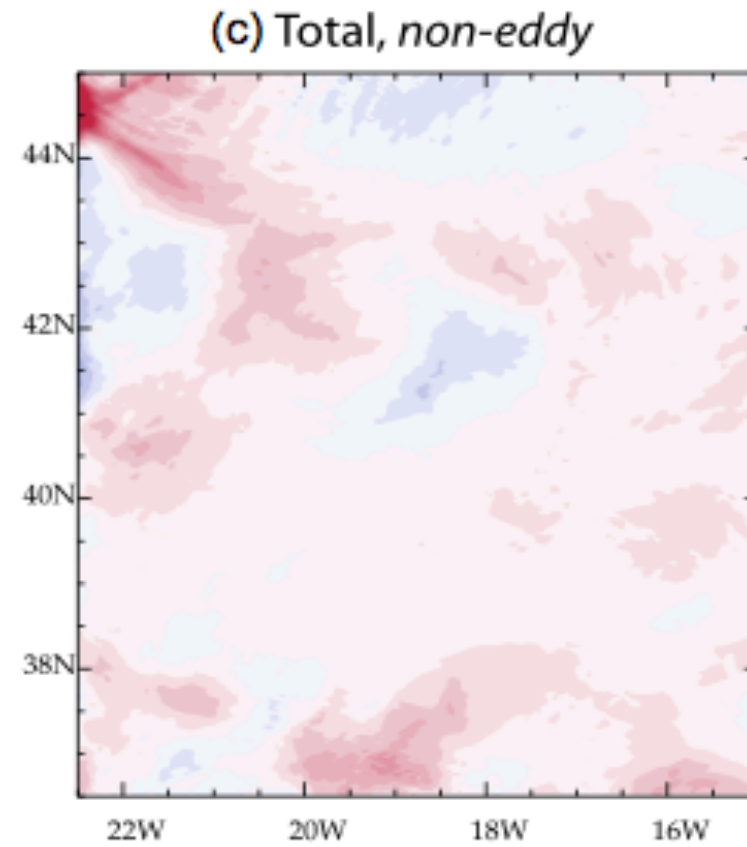
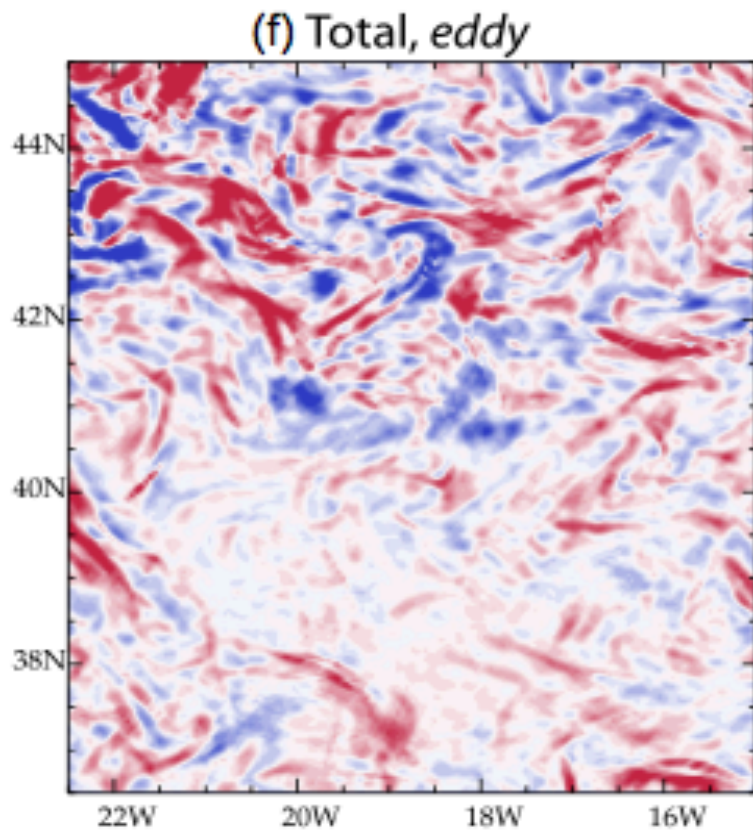
2 simulations

Eddy

Non-eddy

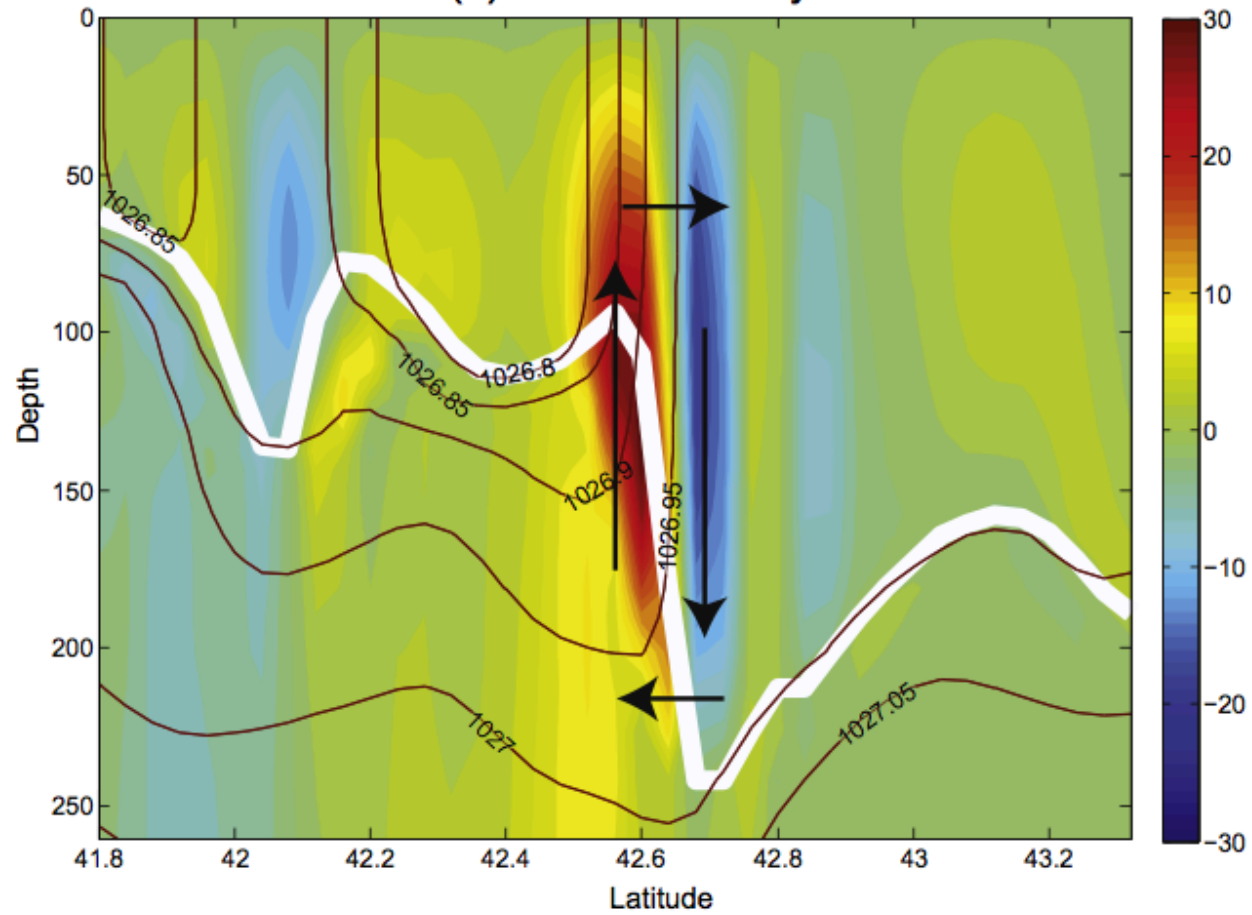


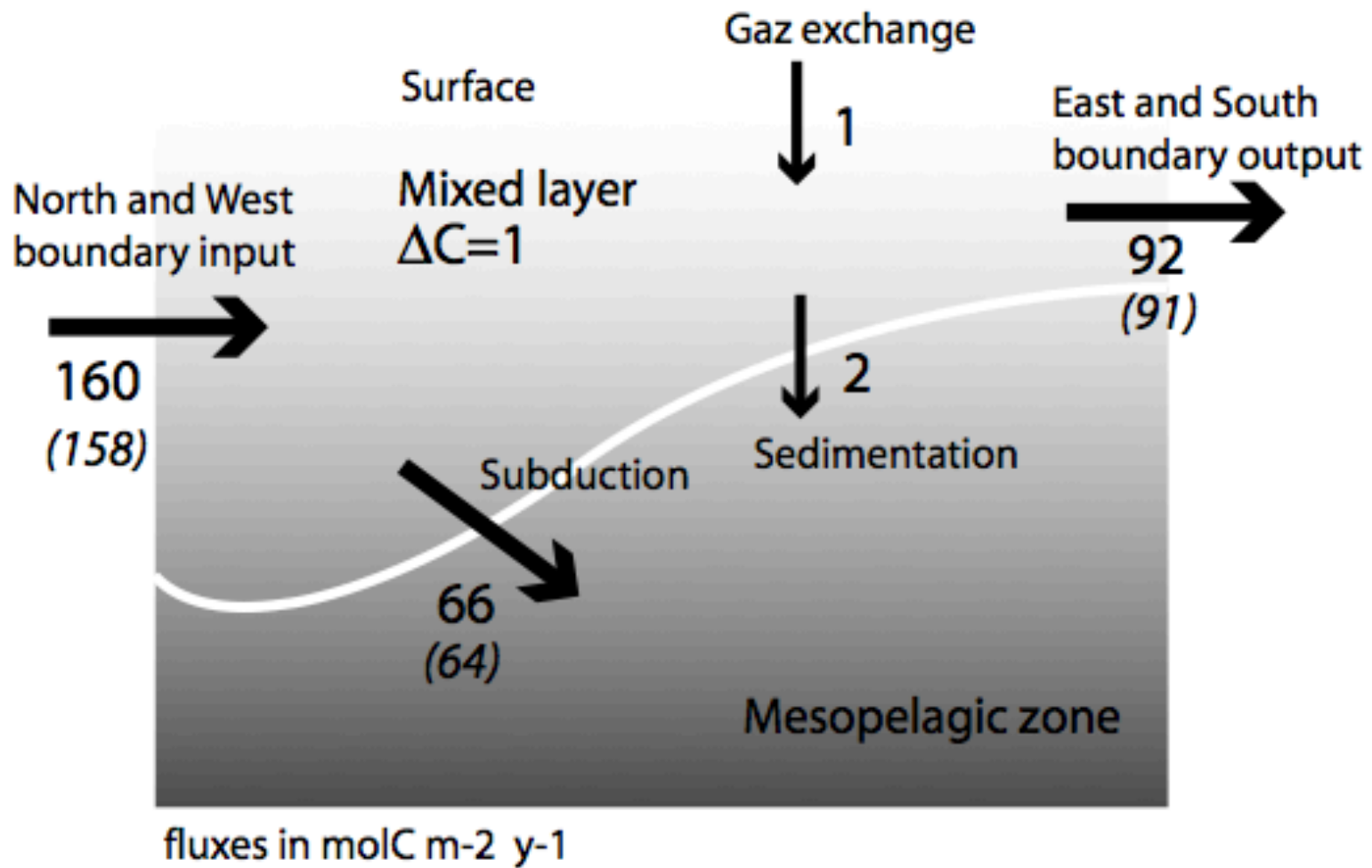
Karleskind et al, 2011



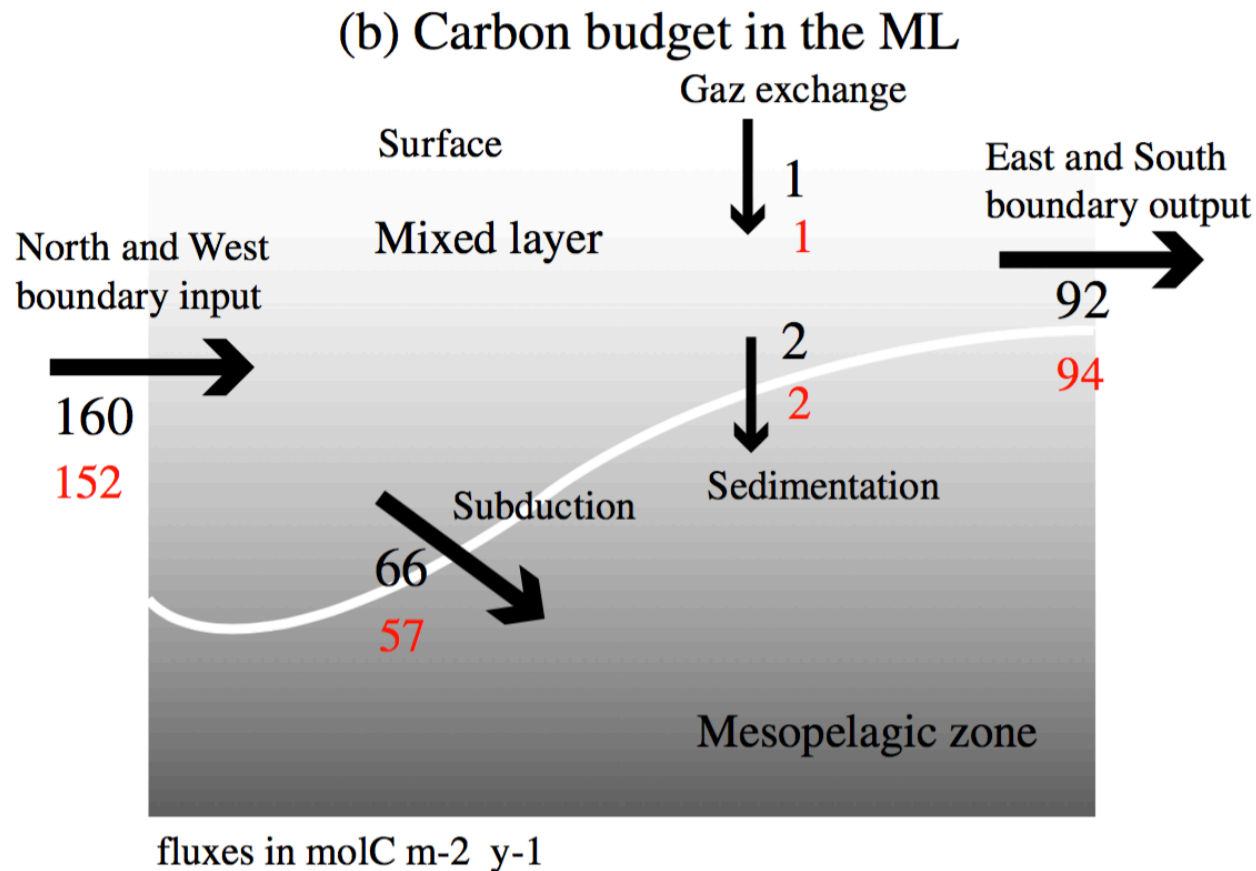
Subduction rate

(e) Vertical velocity





Carbon penetration in the eddy simulation:
 96 % subduction of DIC
 2 % subduction of Organic carbon
 2 % sedimentation of Organic carbon



Carbon penetration in the non-eddy simulation (red):

- 10 % less subduction of DIC
- 10 % less subduction of Organic carbon
- No change sedimentation of Organic carbon



- At temperate latitudes, more than 90% of the carbon flux to the deep ocean due to subduction of DIC, the rest by sedimentation
- 1/3 organic carbon export is through subduction, 2/3 by sedimentation (without contribution of submesoscales)
- Canth air-sea flux and Canth penetration are close in magnitude but do not occur at the same location

- At temperate latitudes, more than 90% of the carbon flux to the deep ocean due to subduction of DIC, the rest by sedimentation
- 1/3 organic carbon export is through subduction, 2/3 by sedimentation (without contribution of submesoscales)
- Canth air-sea flux and Canth penetration are close in magnitude but do not occur at the same location
- Contribution of submesoscales to the export of DIC, DOC, POC and Canth ?
- Vertical structure of the physical export ?
- Interplay between the changes in physical fluxes and of the biological pump efficiency ?