

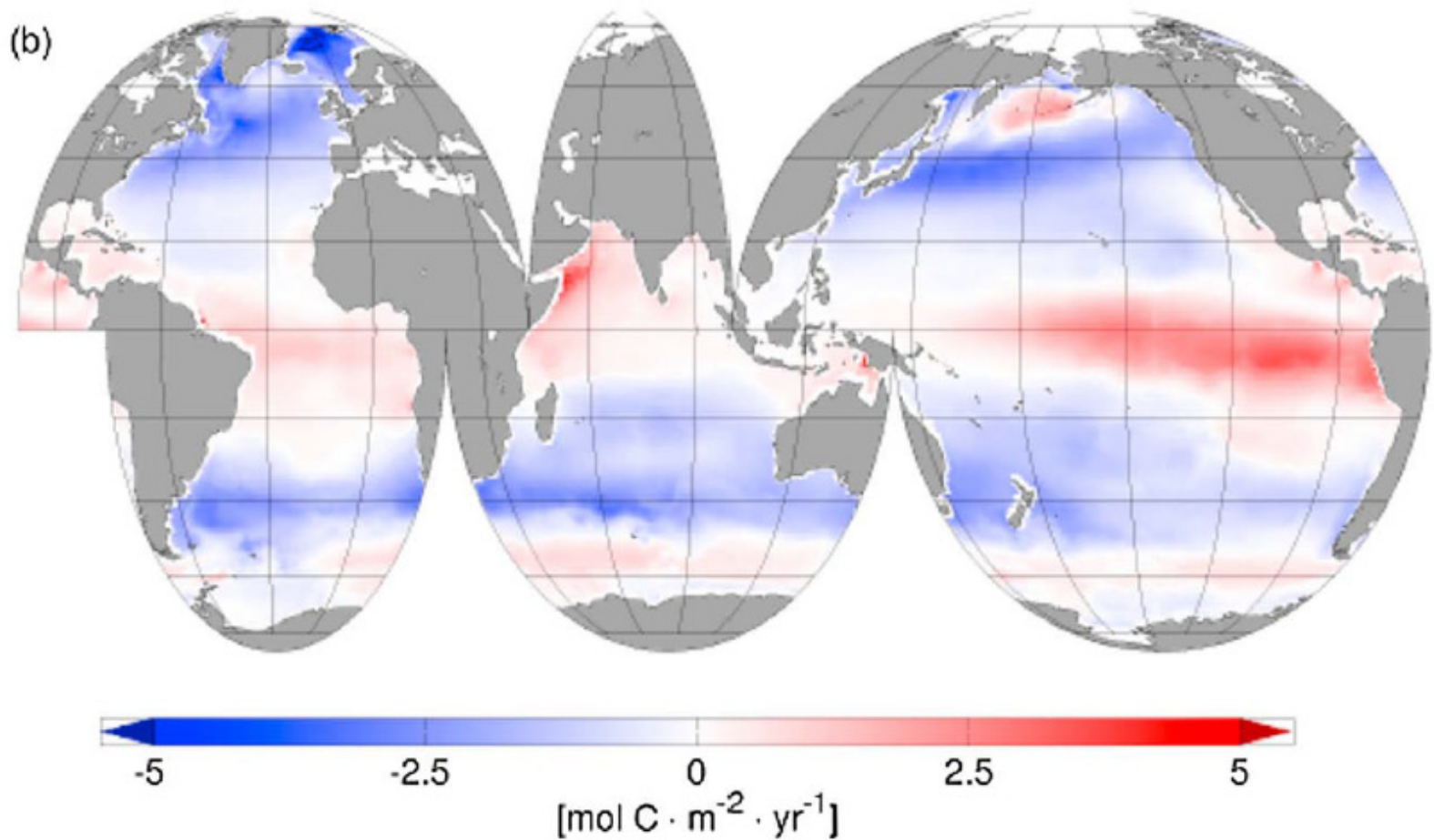
# A large Southern Ocean CO<sub>2</sub> source detected by biogeochemical profiling floats

A. Gray (postdoc, Princeton), Ken Johnson (MBARI), J. L. Sarmiento (Princeton), et al.

- I. Introduction
- II. Methods & results
  - A.  $p\text{CO}_2$
  - B.  $\Delta p\text{CO}_2$
  - C. CO<sub>2</sub> flux
- III. Hypotheses
  - A. Methodology is flawed?
  - B. Previous studies underestimate winter degassing?
  - C. Recent years are anomalous?

Conclusion: CO<sub>2</sub> flux to atmosphere  
>> than previous estimates

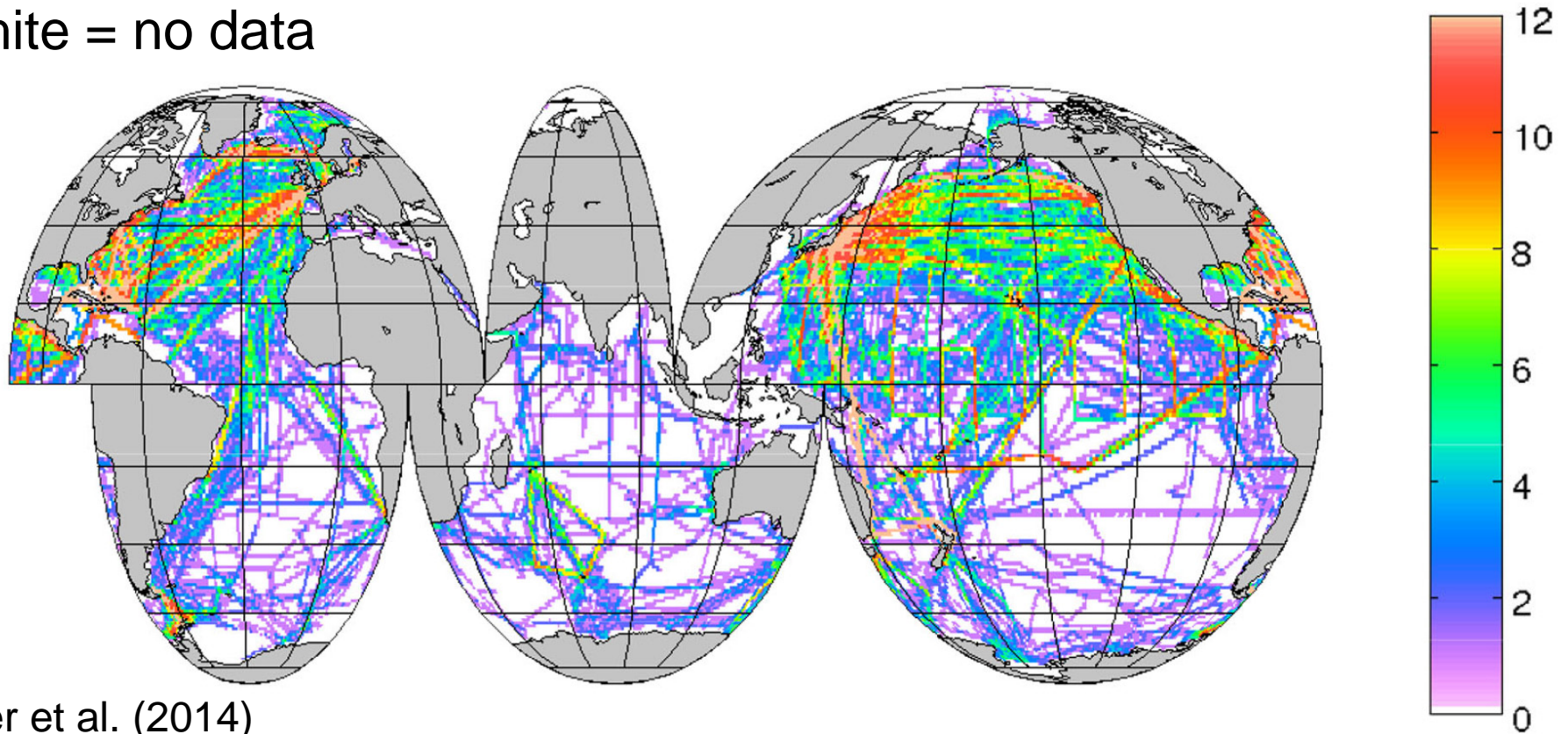
# Landschützer et al. (2014) climatological air-sea flux estimate for 1998-2011



# Undersampling of $p\text{CO}_2$

Months of year with surface  $p\text{CO}_2$  measurements based on all measurements between 1970 to 2011 binned in  $1^\circ$  squares.

White = no data



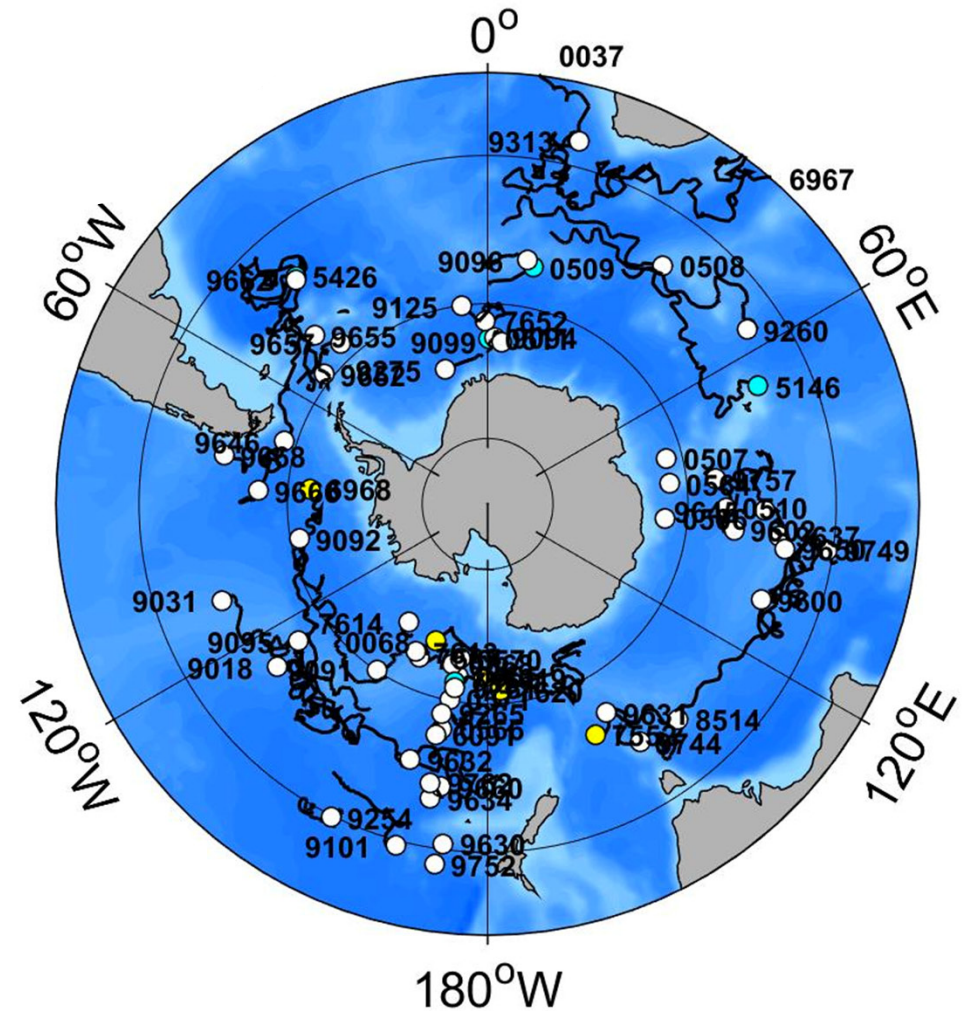
Bakker et al. (2014)



# SOCCOM Floats (so far...)



Hannah  
Zanowski



As of July 22, 2016

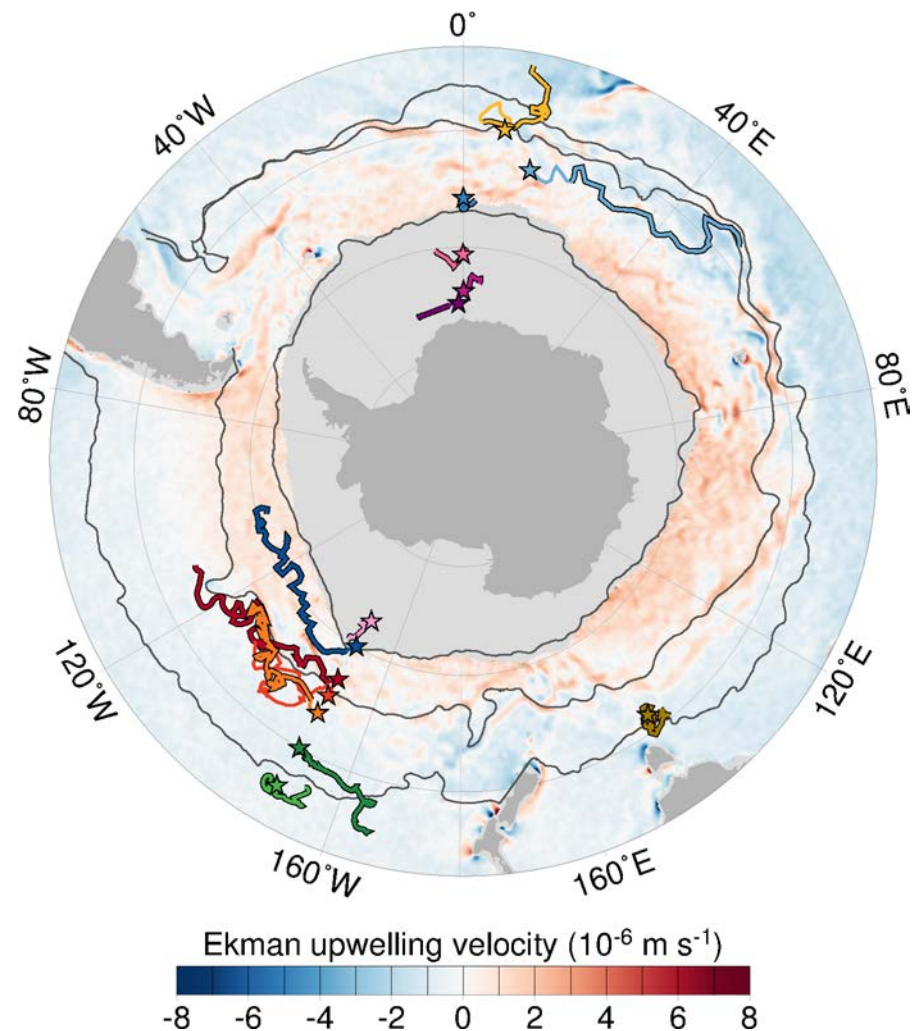
available at: <http://socom.princeton.edu>

## SOCCOM floats we will discuss (14)

Four zones are defined by the fronts:

- Subtropical Zone (STZ)
- Subantarctic Zone (SAZ)
- Antarctic Circumpolar Current Zone (ACCZ)
- Seasonal Ice Zone (SIZ)

Ekman velocity is from QuikSCAT 1999-2009 winds (Risien and Chelton 2011).

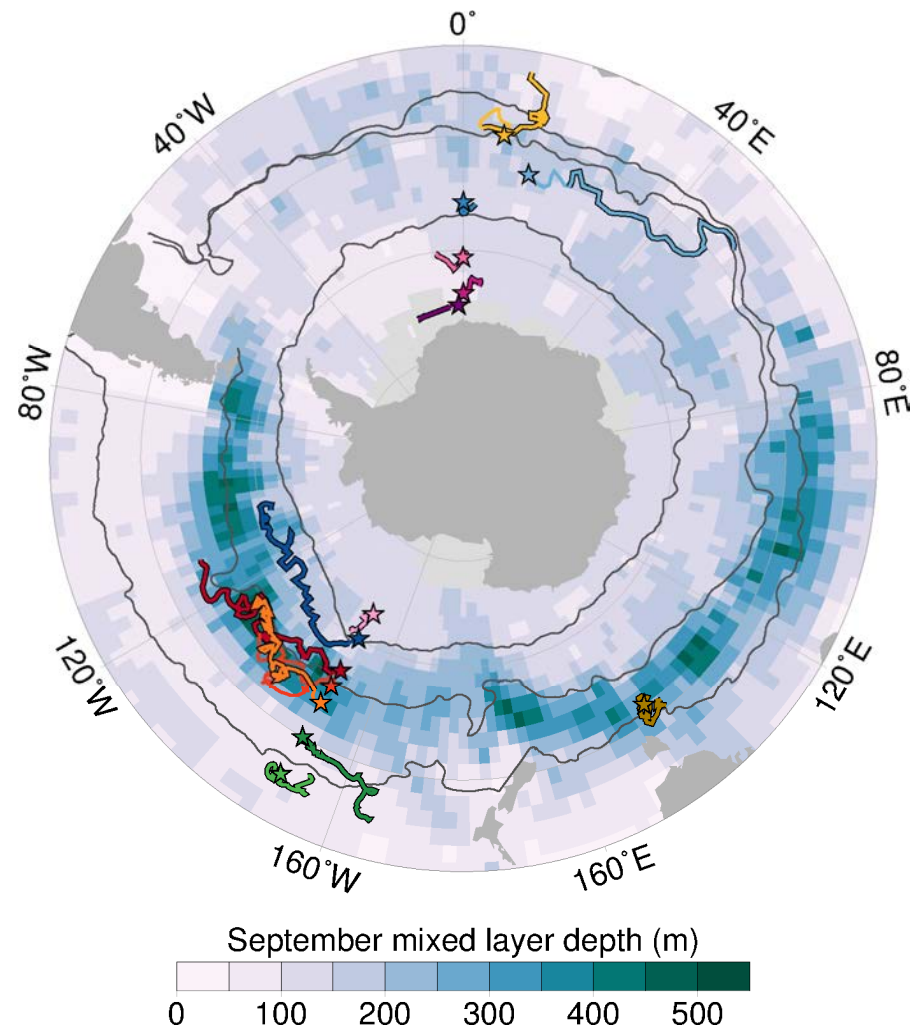


## SOCCOM floats we will discuss (14)

Four zones are defined by the fronts:

- Subtropical Zone (STZ)
- Subantarctic Zone (SAZ)
- Antarctic Circumpolar Current Zone (ACCZ)
- Seasonal Ice Zone (SIZ)

MLDs from the de Boyer-Montégut et al. (2004) climatology.



# A large Southern Ocean CO<sub>2</sub> source detected by biogeochemical profiling floats

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→ II. Methods & results

- A.  $p\text{CO}_2$
- B.  $\Delta p\text{CO}_2$
- C. CO<sub>2</sub> flux

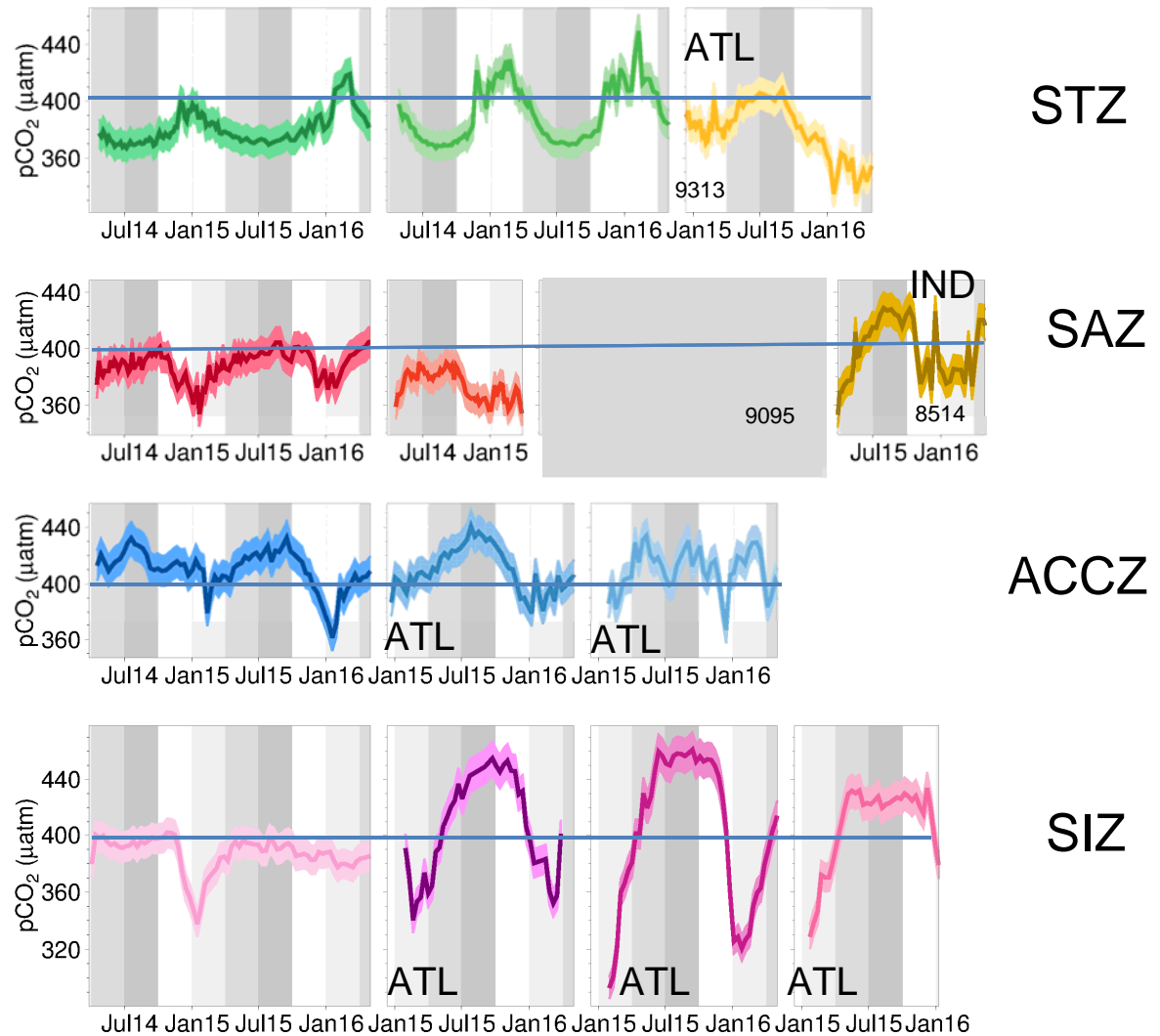
III. Hypotheses

- A. Methodology is wrong
- B. Previous estimates are wrong
- C. Recent years are anomalous

Conclusion: CO<sub>2</sub> flux to atmosphere  
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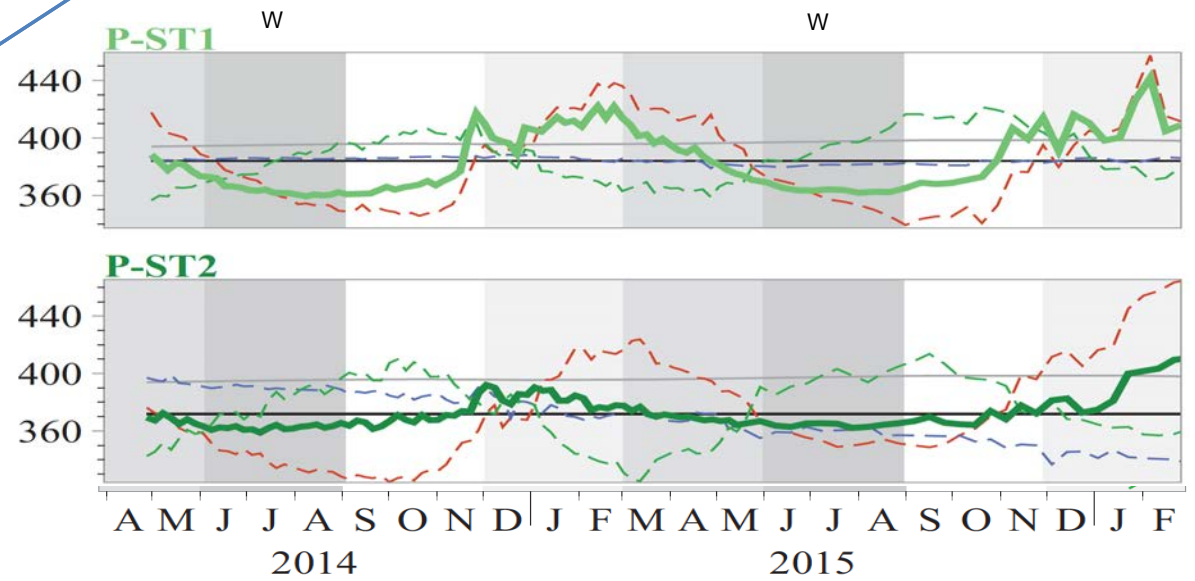
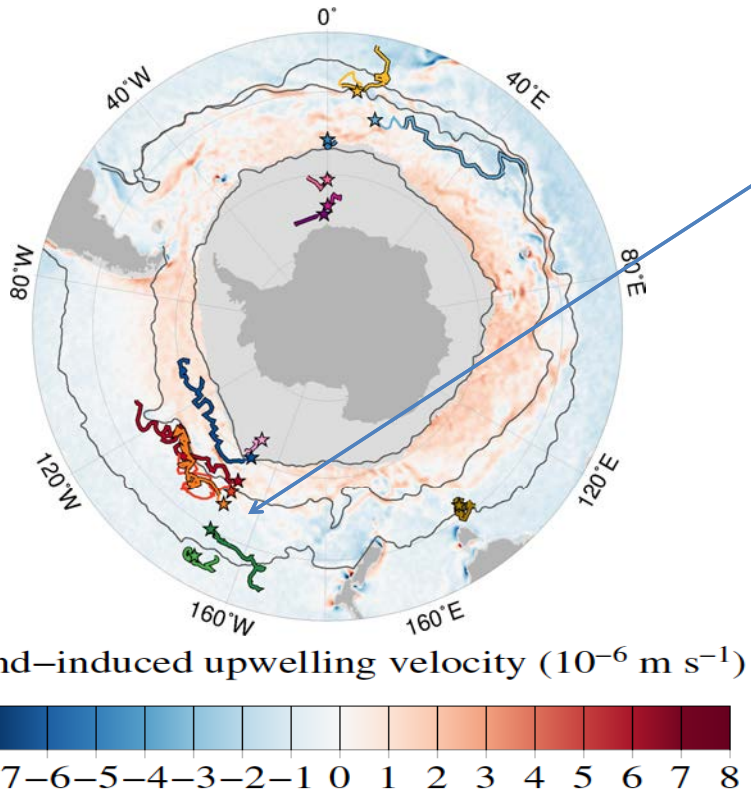
## pCO<sub>2</sub> from float pH & MLR alkalinity

- All floats are in the Pacific unless labeled otherwise
- Dark gray bar is winter
- Unusual floats:
  - **Float 9313** (yellow) from Atlantic STZ crosses front
  - **Float 9095** (grayed out) from Pacific SAZ: pH sensor had uncharacteristically large drift. Left out of subsequent analyses.
  - **Float 8514** (gold) from Indian Ocean is in an active eddy region

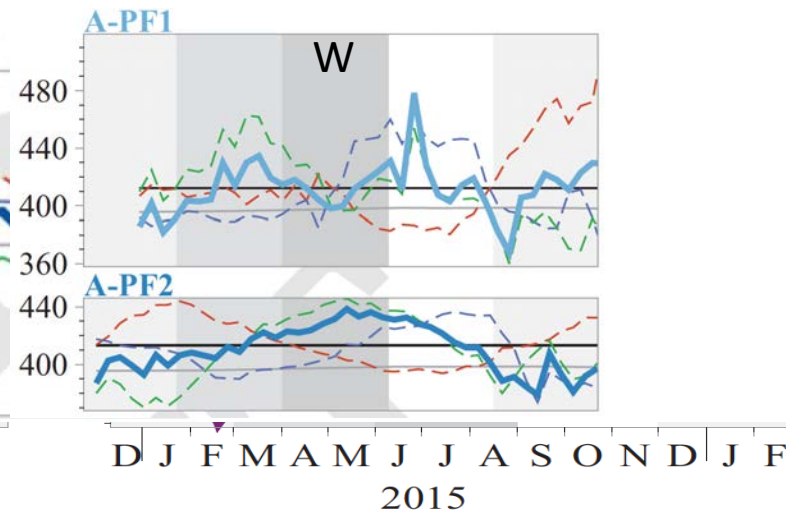
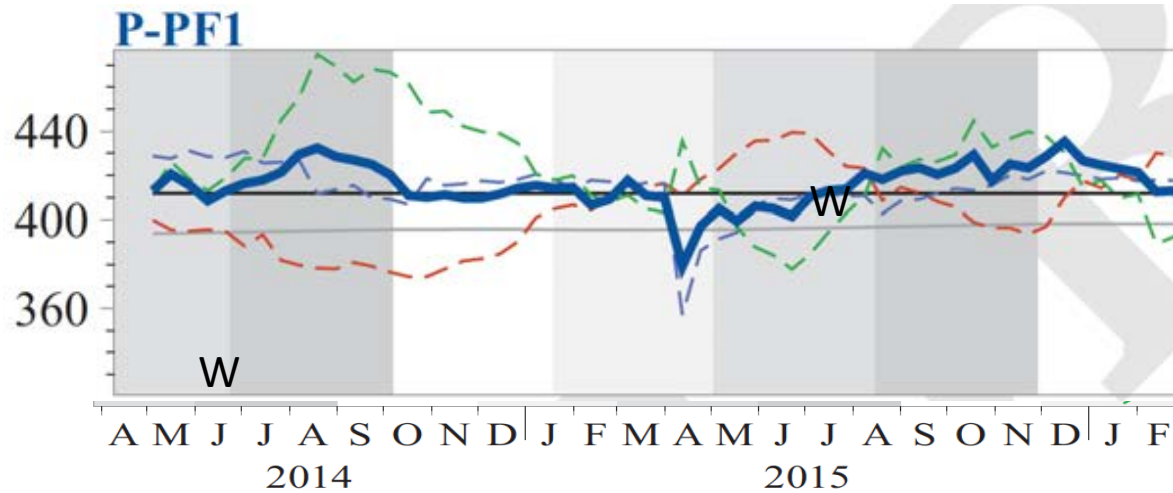
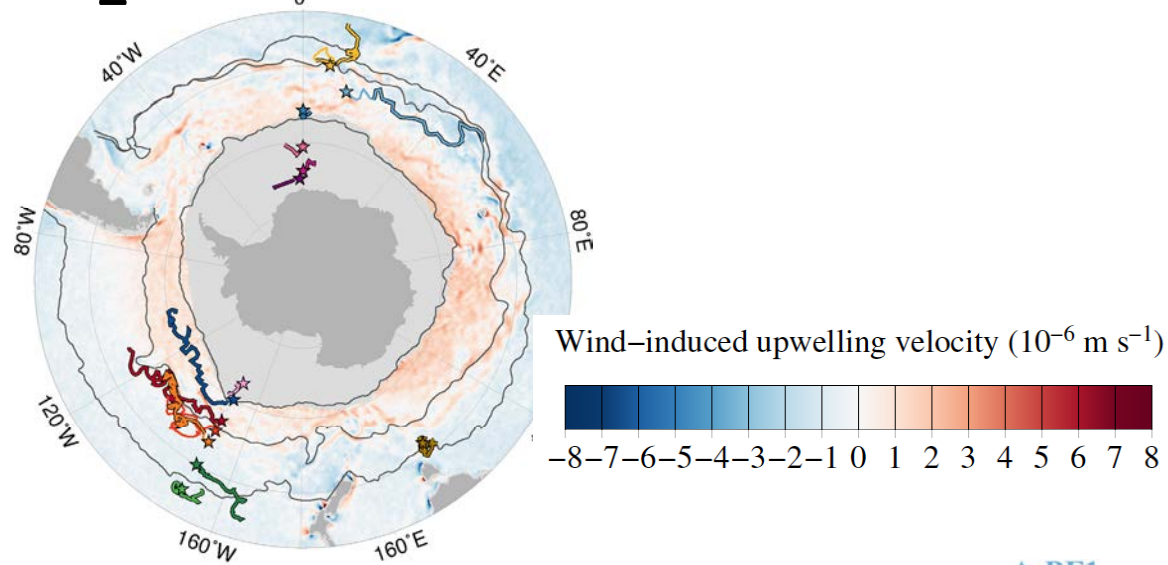




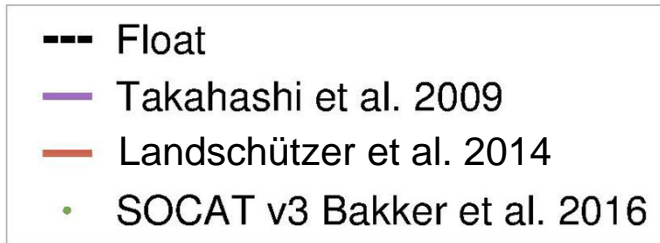
# pCO<sub>2</sub> in subtropical gyre (Pacific only)



# pCO<sub>2</sub> in ACC/upwelling zone

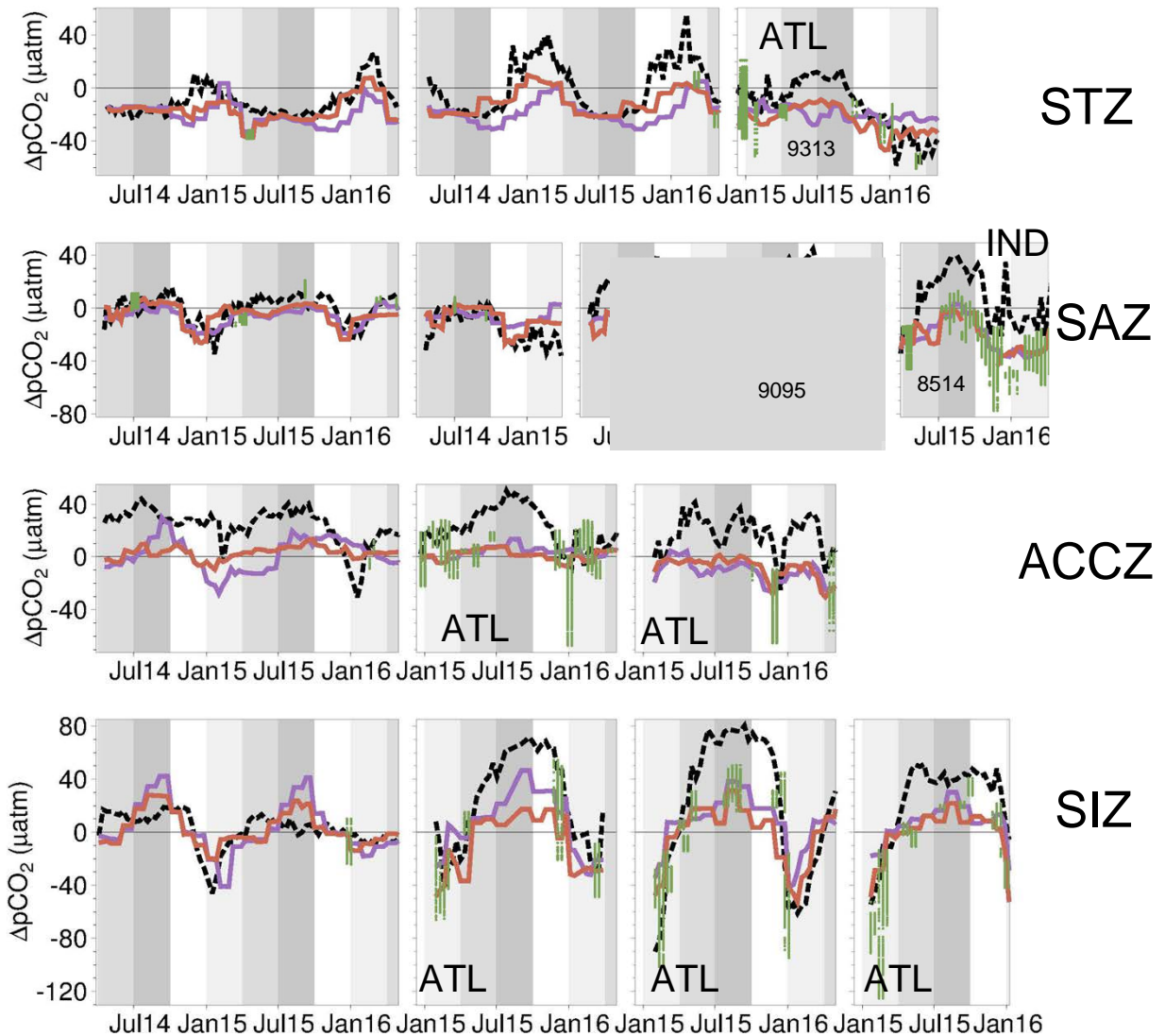


$$\Delta p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$$



Recall:

- **Float 9313** from Atlantic STZ crosses front
- **Float 9095** (grayed out) from Pacific SAZ: *pH* sensor had uncharacteristically large drift. Left out of subsequent analyses.
- **Float 8514** from Indian Ocean is in an active eddy region



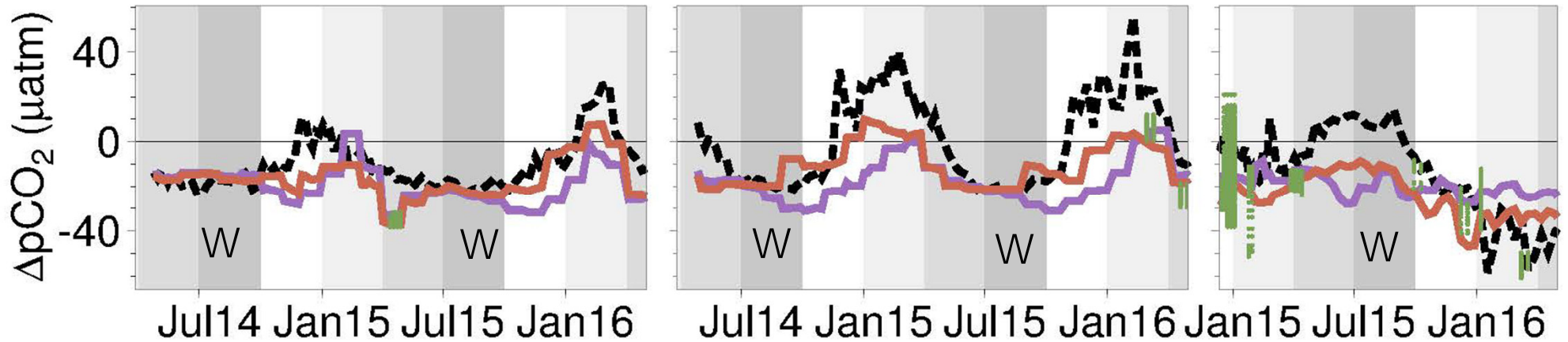
# $\Delta p\text{CO}_2$ in STZ

$$\Delta p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$$

- Float
- Takahashi et al. 2009
- Landschützer et al. 2014
- SOCAT v3 Bakker et al. 2016

Pacific

Atlantic



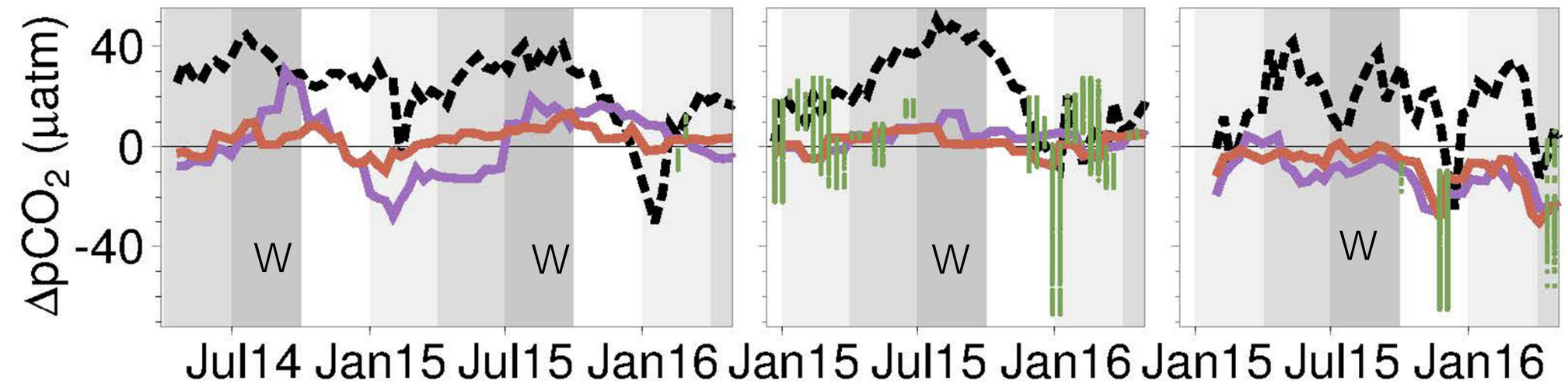
# $\Delta p\text{CO}_2$ in ACC zone

$$\otimes p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$$

- Float
- Takahashi et al. 2009
- Landschützer et al. 2014
- SOCAT v3 Bakker et al. 2016

Pacific

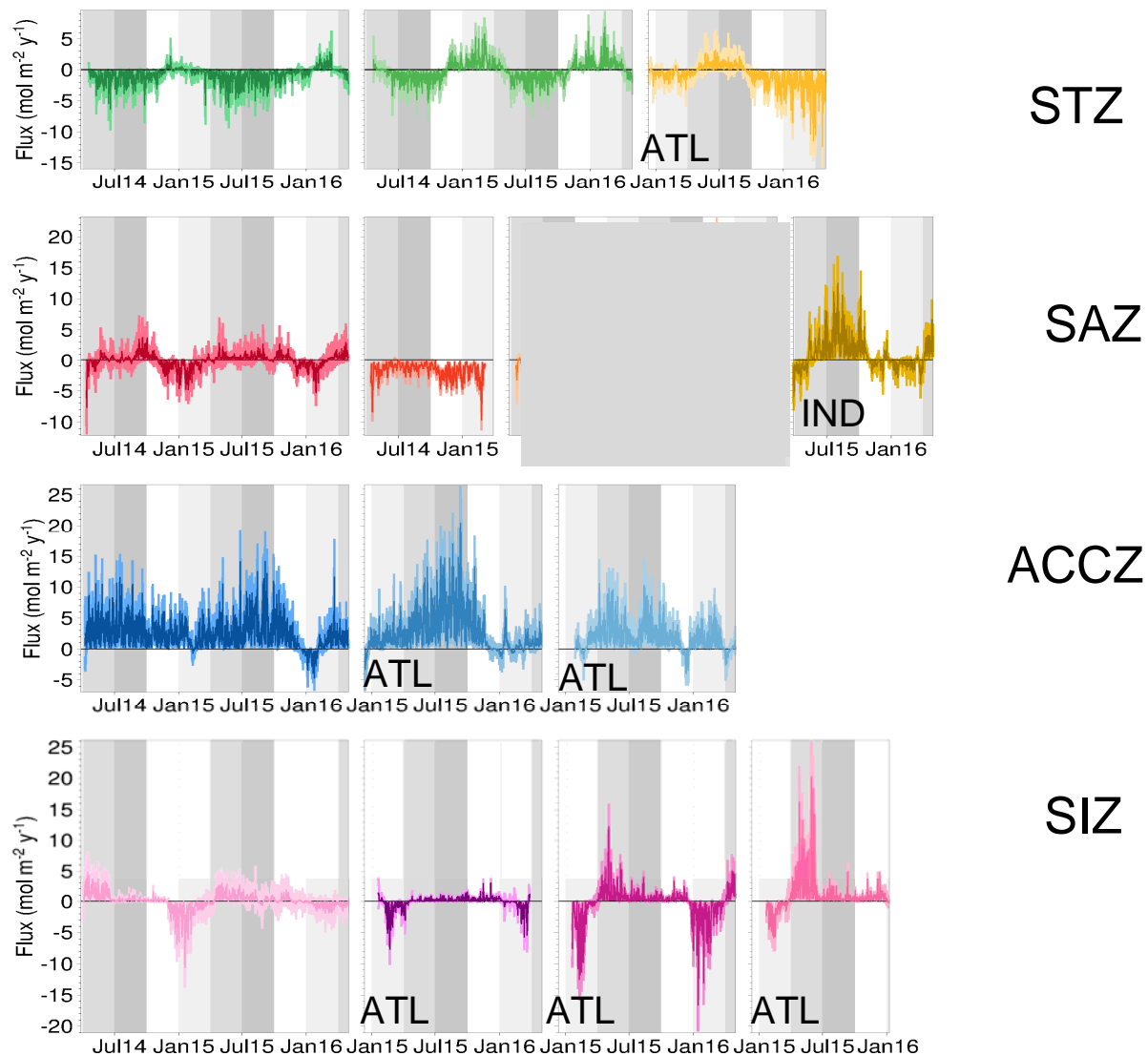
Atlantic





## Sea-air CO<sub>2</sub> flux

- Flux equation:  $F = k_g * K_0 * \Delta pCO_2$
- $k_g$  = wind speed squared dependent gas transfer velocity from Wanninkhof (2014)
- Winds are 6-hourly ERA-interim
- $K_0$  = solubility constant
- $\Delta pCO_2 = pCO_2^{ocn} - pCO_2^{atm}$   
with  $pCO_2^{atm}$  from Cape Grim monthly data



## Annual net CO<sub>2</sub> flux to atmosphere

- Units: mol m<sup>-2</sup> y<sup>-1</sup>
- Covers region south of 35°S
- Positive indicates net outgassing, negative is uptake.
- Bold numbers are statistically different from zero
- .
- For scale, the average anthropogenic CO<sub>2</sub> uptake is 0.46 ± 0.14 mol m<sup>-2</sup> y<sup>-1</sup>
- **Note:** problematic float 9095 (in parentheses) is not included in mean

Zone	Float	Flux May14-May15	Flux May15-May16	Mean mol m <sup>-2</sup> y <sup>-1</sup>
STZ	P-ST1	-0.09 ± 0.5	-0.2 ± 0.5	<b>-0.6 ± 0.3</b>
	P-ST2	<b>-0.8 ± 0.5</b>	<b>-0.8 ± 0.5</b>	
	A-SA1	-	<b>-1.0 ± 0.7</b>	
SAZ	P-SA1	<b>(1.3 ± 0.7)</b>	<b>(2.9 ± 0.8)</b>	0.3 ± 0.5
	P-SA3	-0.2 ± 0.8	0.1 ± 0.7	
	I-SA1	-	<b>1.1 ± 0.7</b>	
ACCZ	P-ACC1	<b>2.6 ± 0.8</b>	<b>2.2 ± 0.8</b>	<b>2.4 ± 0.5</b>
	A-ACC1	-	<b>1.9 ± 0.8</b>	
	A-ACC2	-	<b>3.1 ± 1.0</b>	
SIZ	P-SI1	-0.2 ± 0.5	0.1 ± 0.7	0.2 ± 0.3
	A-SI1	-	<b>0.9 ± 0.4</b>	
	A-SI2	-	-0.006 ± 0.2	
	A-SI3	-	0.03 ± 0.16	

## Annual net oceanic CO<sub>2</sub> uptake

- Units Pg C y<sup>-1</sup>
- Covers region south of 35°S
- Positive indicates net outgassing, negative is uptake.
- Dashes indicate fluxes less than 0.03

Zone	SOCOM floats	Takahashi et. al. (2009)	Landschützer et al. (2014)
STZ	-0.2 ± 0.1	-0.4	-0.5
SAZ	0.1 ± 0.1	-0.2	-0.2
ACCZ	0.8 ± 0.2	-0.2	-0.1
SIZ	--	--	--
TOTAL	0.7 ± 0.2	-0.8	-0.8
L			

## Annual net oceanic CO<sub>2</sub> uptake

- Units Pg C y<sup>-1</sup>
- Covers region south of 35°S
- Positive indicates net outgassing, negative is uptake.
- Dashes indicate fluxes less than 0.03
- Final column is calculated from Landschützer et al. (2015) by subsampling where the float profiles are, then integrating, averaging, and multiplying by the same areas.

Zone	SOCCOM floats	Landschützer et al. (2014)	Subsampled as SOCCOM
STZ	-0.2 ± 0.1	-0.5	-0.3
SAZ	0.1 ± 0.1	-0.2	-0.2
ACCZ	0.8 ± 0.2	-0.1	--
SIZ	--	--	--
TOTA	0.7 ± 0.2	-0.8	-0.5

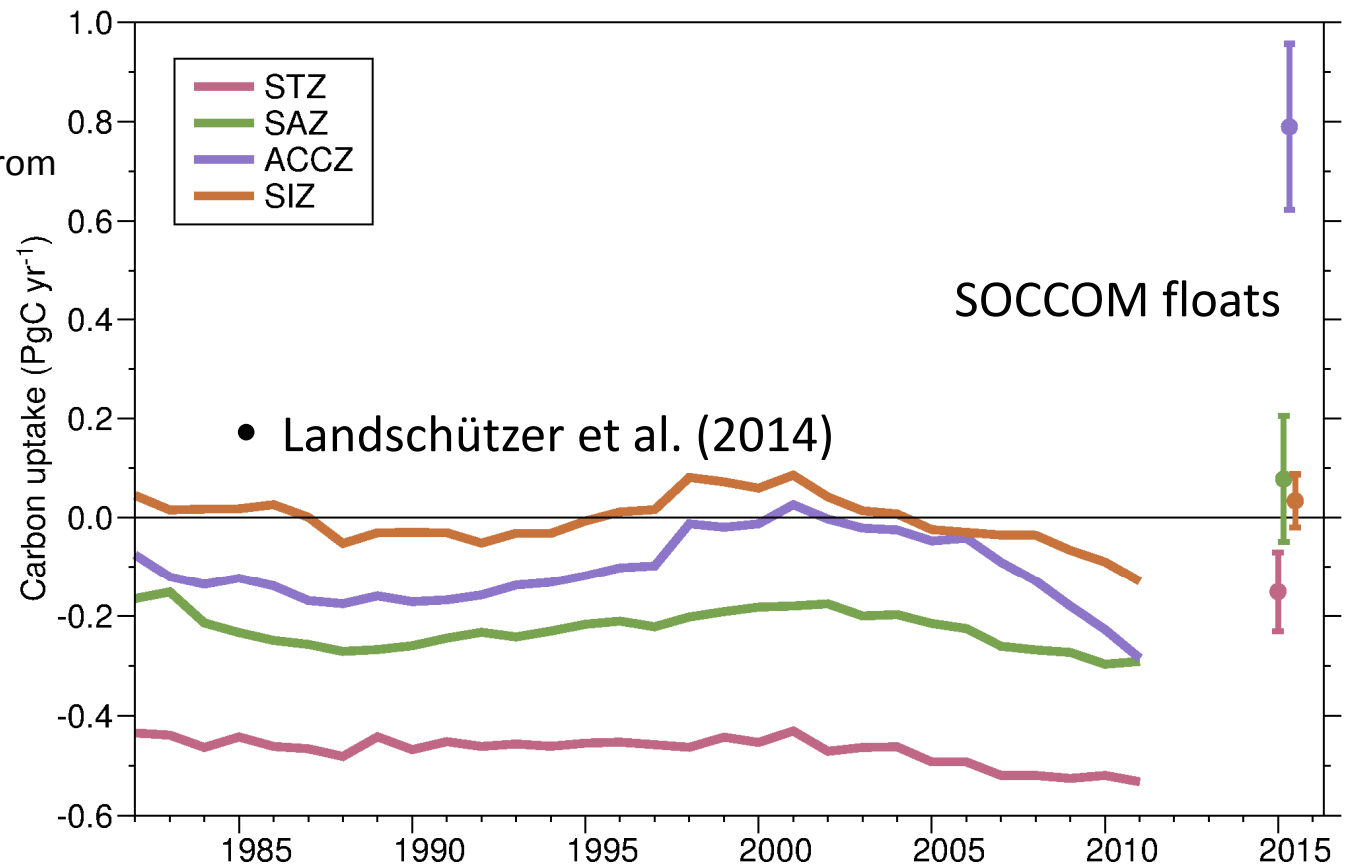
# Conclusion: Southern Ocean CO<sub>2</sub> flux to atmosphere is greater than previous estimates

- Hypothesis 1: flawed methodology
  - Small number of floats is worrisome, but
  - Good agreement when shipboard data is available is reassuring
  - Method for converting pH to pCO<sub>2</sub> looks good.
- Hypothesis 2: Climatological baseline should have a stronger Southern Ocean source
  - Our results support this.
  - However, there must be a larger sink elsewhere to compensate the Southern Ocean source and maintain large ocean carbon sink
- Hypothesis 3: 2014-present is anomalous
  - This appears to be the case, but past history suggest this can only explain ~0.5 Pg C y<sup>-1</sup> of 1.5 Pg C y<sup>-1</sup> anomaly we find



## Time history of CO<sub>2</sub> fluxes (PgC yr<sup>-1</sup>)

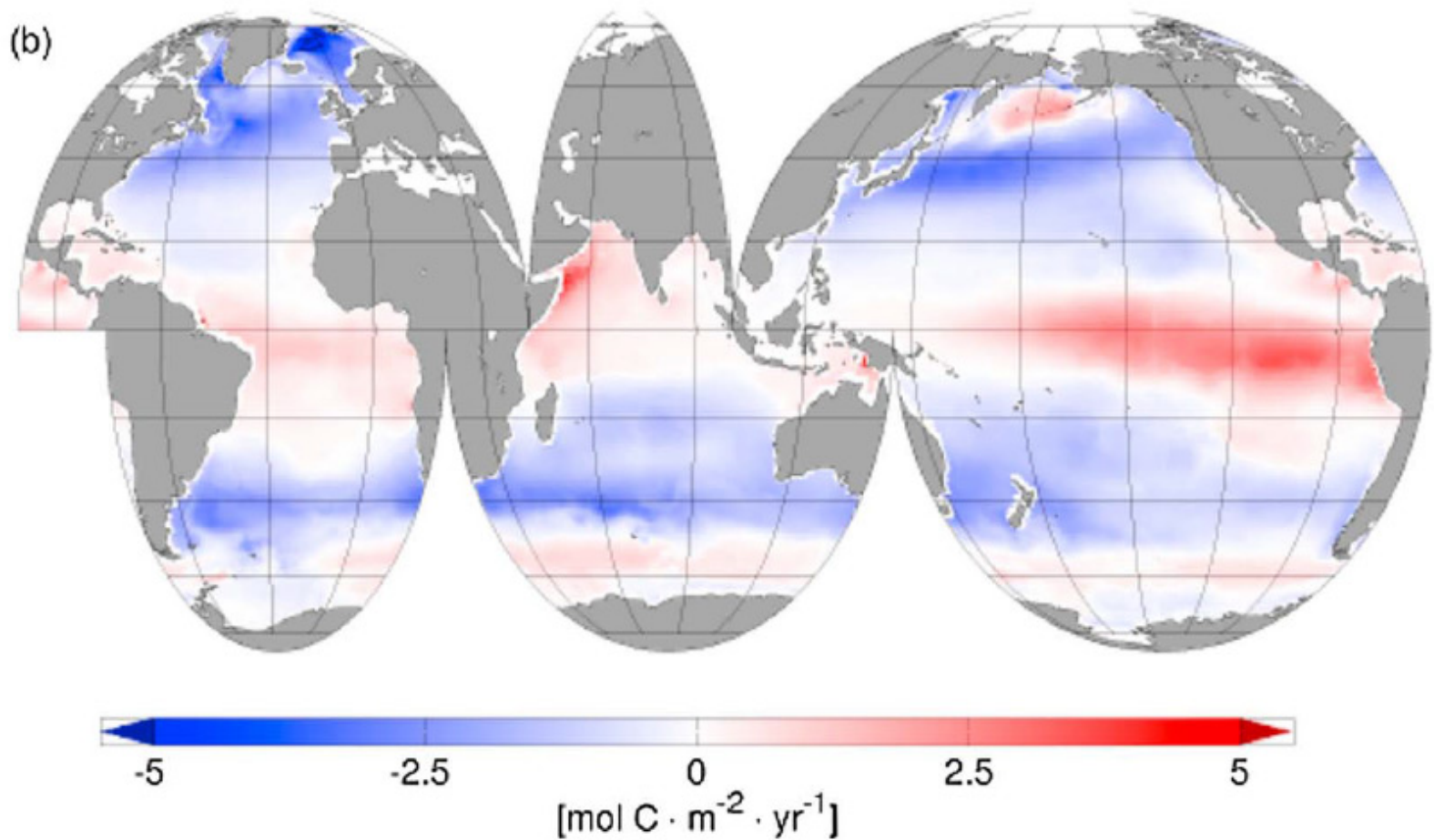
- The 1982 to 2011 time history is from Landschützer et al. (2014)
- Positive is degassing
- The individual points are from SOCCOM floats
- STZ = subtropical zone
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- SIZ = seasonal ice covered zone



# Conclusion: Southern Ocean CO<sub>2</sub> flux to atmosphere is greater than previous estimates

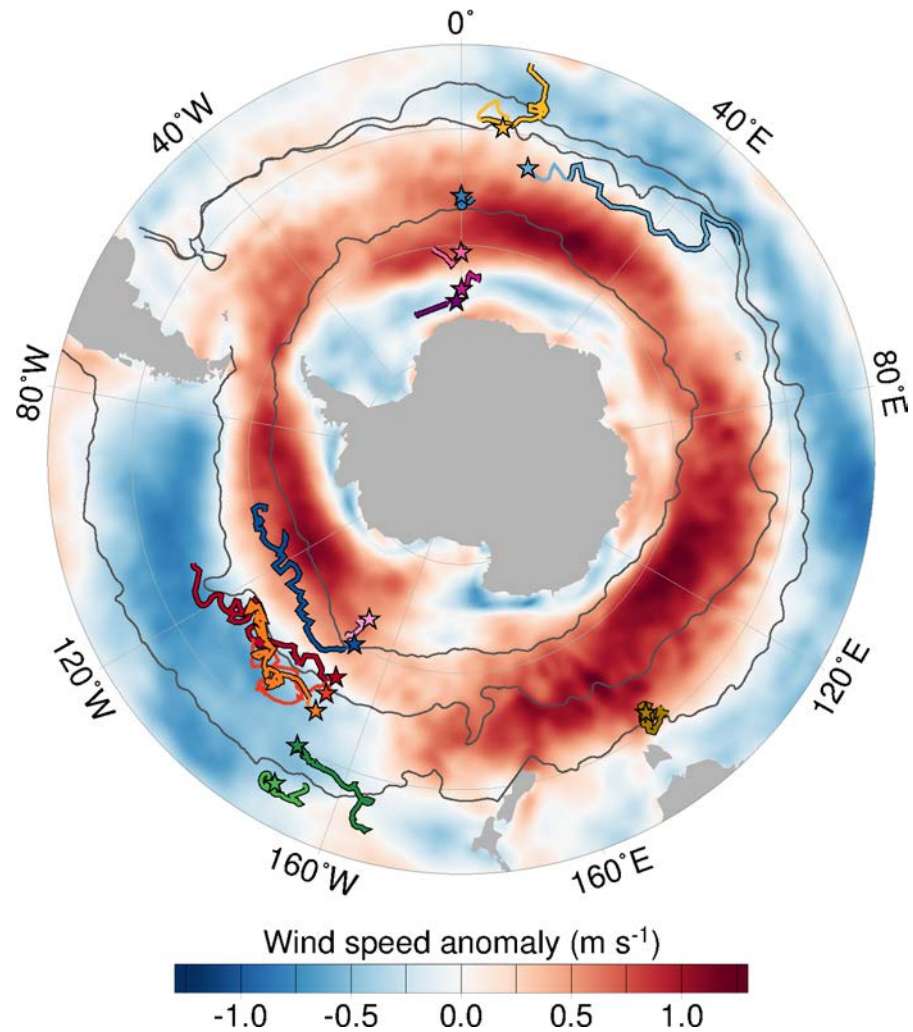
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  - Method for converting pH to pCO<sub>2</sub> looks good.
- Hypothesis 2: Climatological baseline should have a stronger Southern Ocean source
  - Our results support this (if there were data in winter, the baseline would shift up)
  - However, there must be a larger sink elsewhere to compensate the Southern Ocean source and maintain large ocean carbon sink
- Hypothesis 3: 2014-present is anomalous
  - This appears to be the case, but past history suggest this can only explain ~0.5 Pg C y<sup>-1</sup> of 1.5 Pg C y<sup>-1</sup> anomaly we find

# Landschützer et al. (2014) climatological air-sea flux estimate for 1998-2011



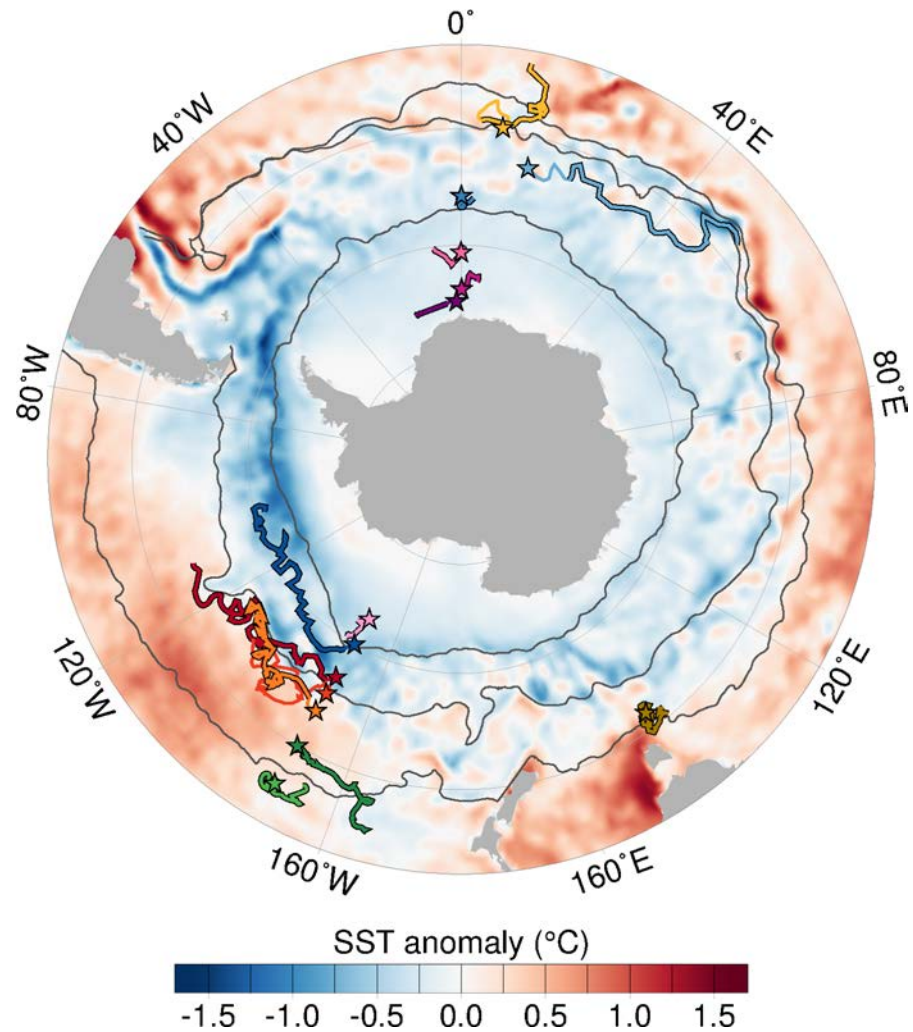
## Wind speed anomaly (April 2014 to 2016)

- The ACC region has increased winds which would imply more upwelling & thus outgassing.
- ERA-interim wind fields, calculated as the mean over Apr 2014 - Apr 2016 minus the mean over Apr 1979 - Apr 2016



## SST Anomaly (April 2014 to 2016)

- The subtropics are warmer, consistent with increased outgassing.
- The ACC region has colder temperatures, consistent with higher upwelling.
- ERA-interim SST fields, calculated as the mean over Apr 2014 - Apr 2016 minus the mean over Apr 1979 - Apr 2016





# Conclusion: Southern Ocean CO<sub>2</sub> flux to atmosphere is greater than previous estimates

- Hypothesis 1: flawed methodology
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## Future work

- SOCCOM
  - Keep adding floats!
  - SOSE
- Test consistency with other constraints
  - Atmospheric observations & atmospheric inversion models
  - Interior ocean DIC observations & ocean inverse, SOSE, etc.
  - How would addition of float data affect the Landschützer neural network map?

# What is driving changes? – ML budget

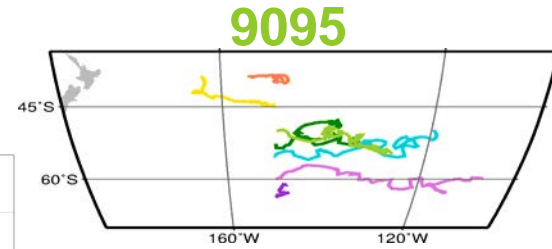
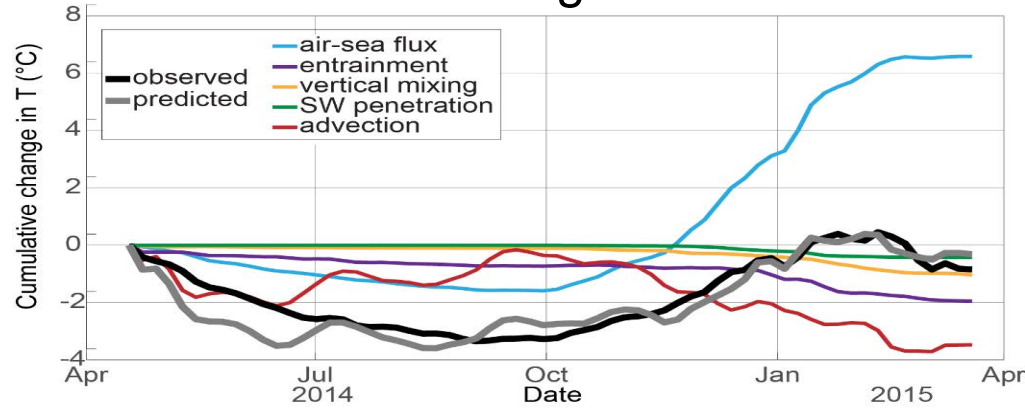
<b>Change in <math>C_T^{ml}</math> =</b>	Estimated from float obs
<b>Entrainment +</b>	Observed $dMLD/dt$ , $C_T(z)$
<b>Vertical mixing +</b>	Observed $dC_T/dz$ , assume $K_z (1 \times 10^{-4} m s^{-2})$
<b>Air-sea flux +</b>	ECMWF wind speed, observed $\Delta pCO_2$ , Wanninkhof 2014 coefficient
<b>Dilution +</b>	ECMWF Evaporation, Precipitation
<b>Advection +</b>	AVISO $\mathbf{u}_{geo}$ , ECMWF wind $\mathbf{u}_{ek}$ , AVHRR SST, linear regression with GLODAPv2
<b>Biology</b>	Residual optimized with nitrate budget

# Mixed layer budget

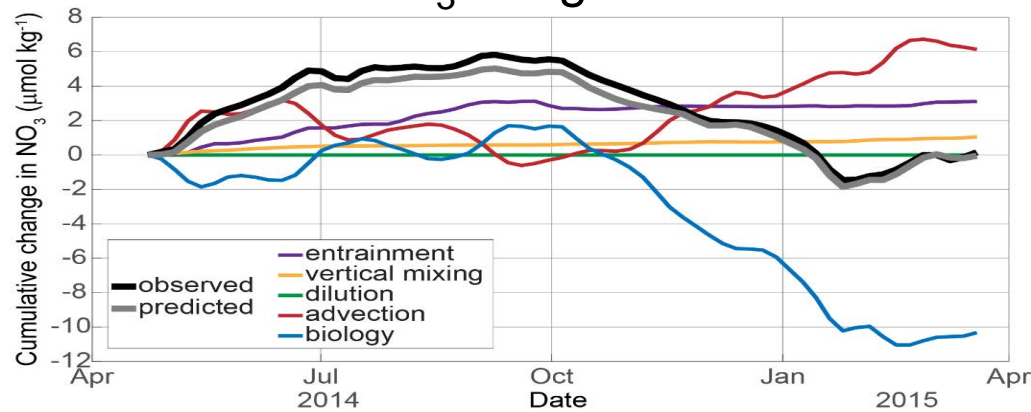
C:N = 6.0

NCP = 2.3 molC m<sup>-2</sup> m<sup>-y</sup>

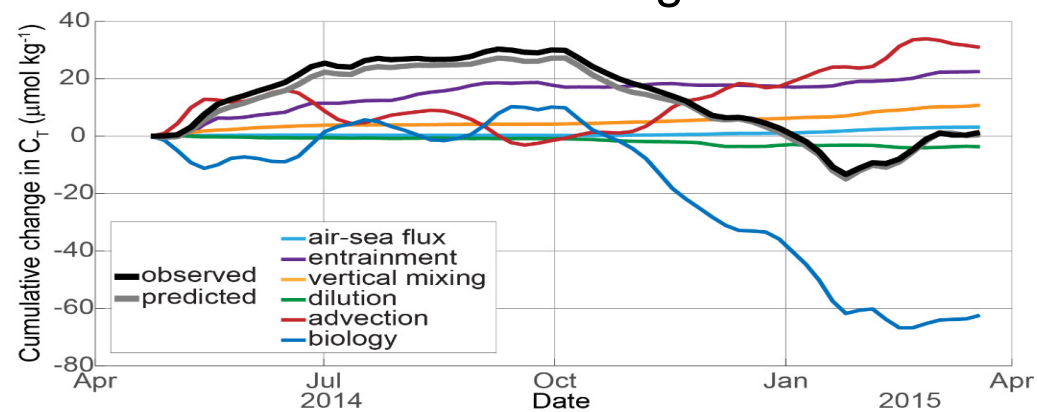
T budget



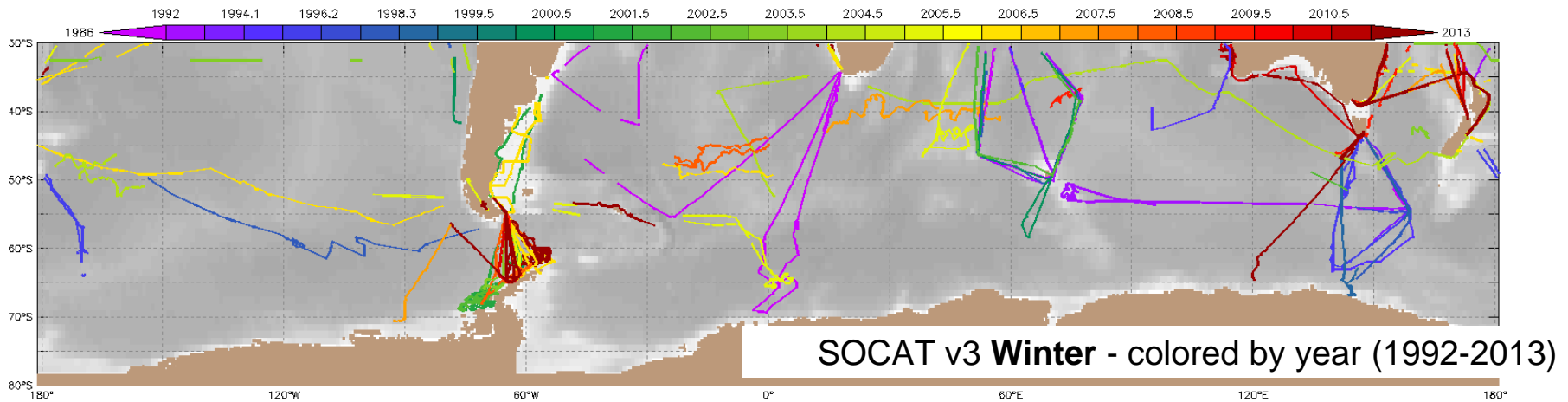
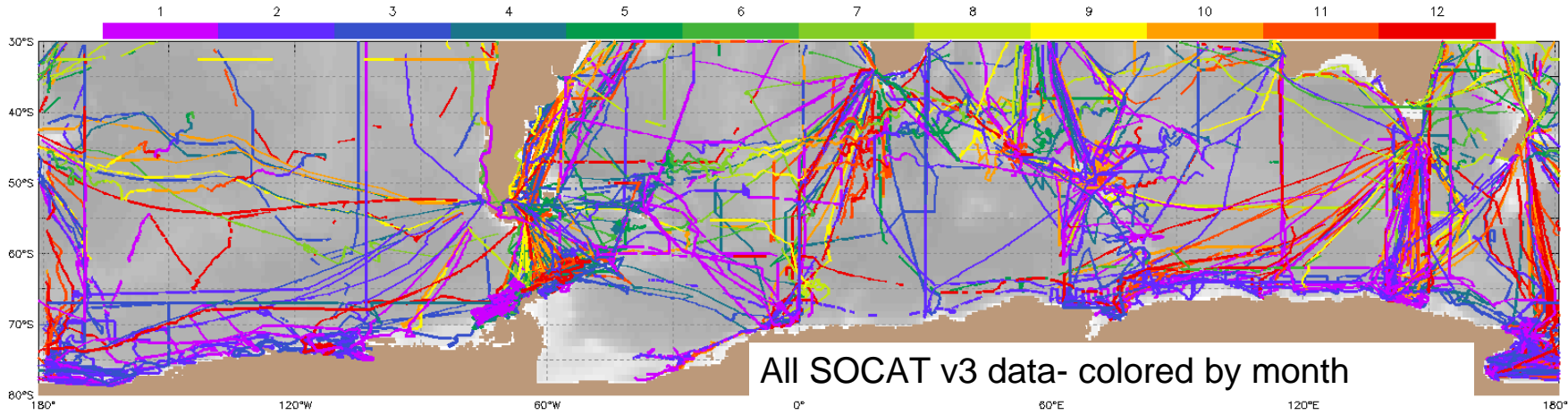
NO<sub>3</sub> budget



Carbon budget

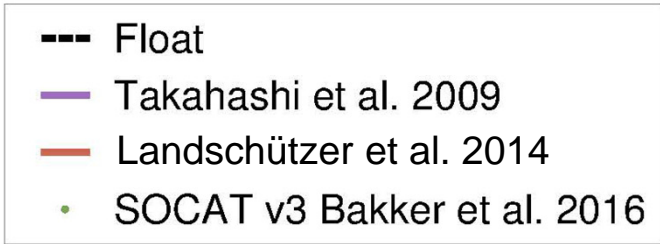


# SOCAT v. 3 data (Bakker et al., 2016)

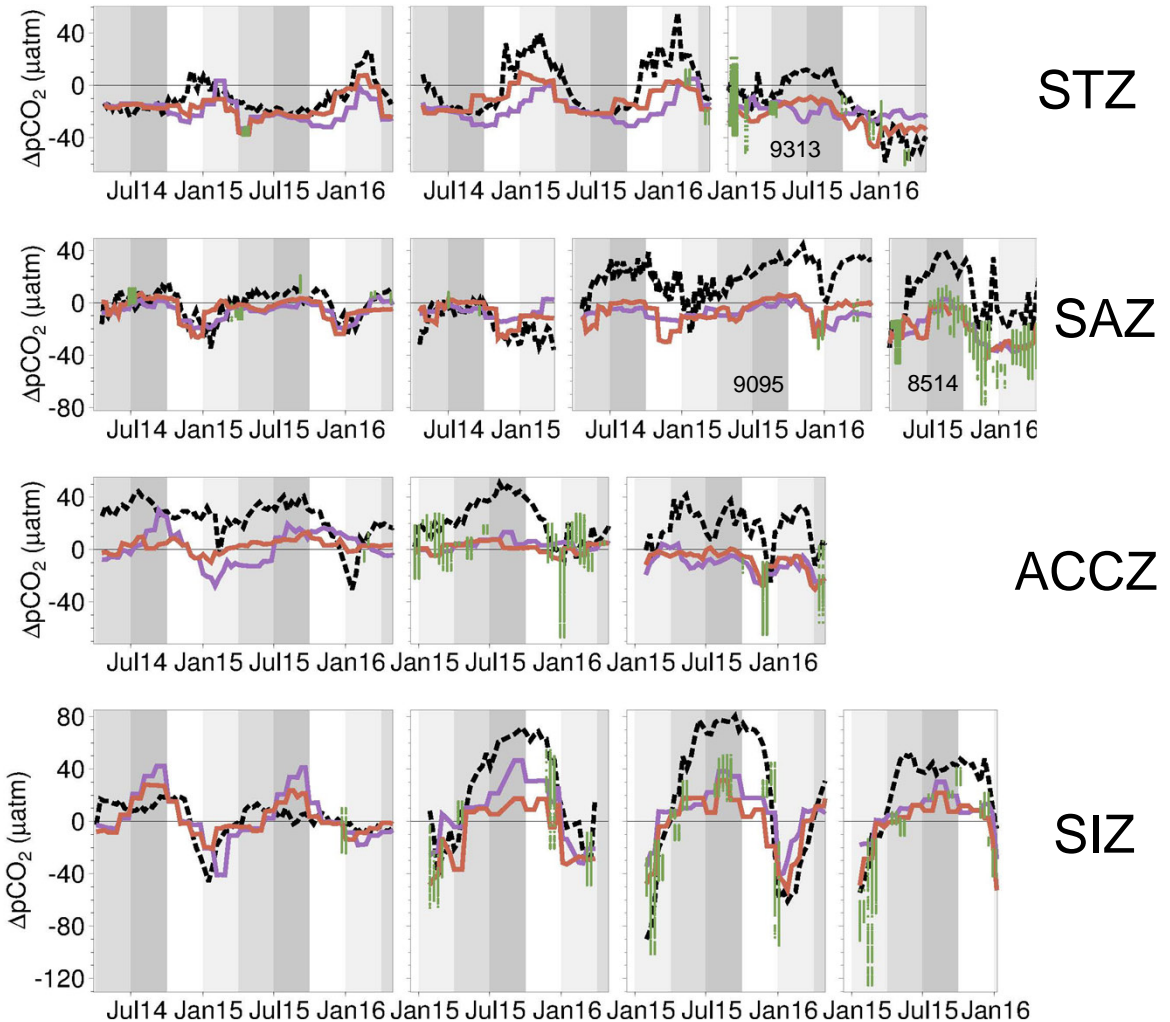




# Comparison of estimates of $\Delta p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$

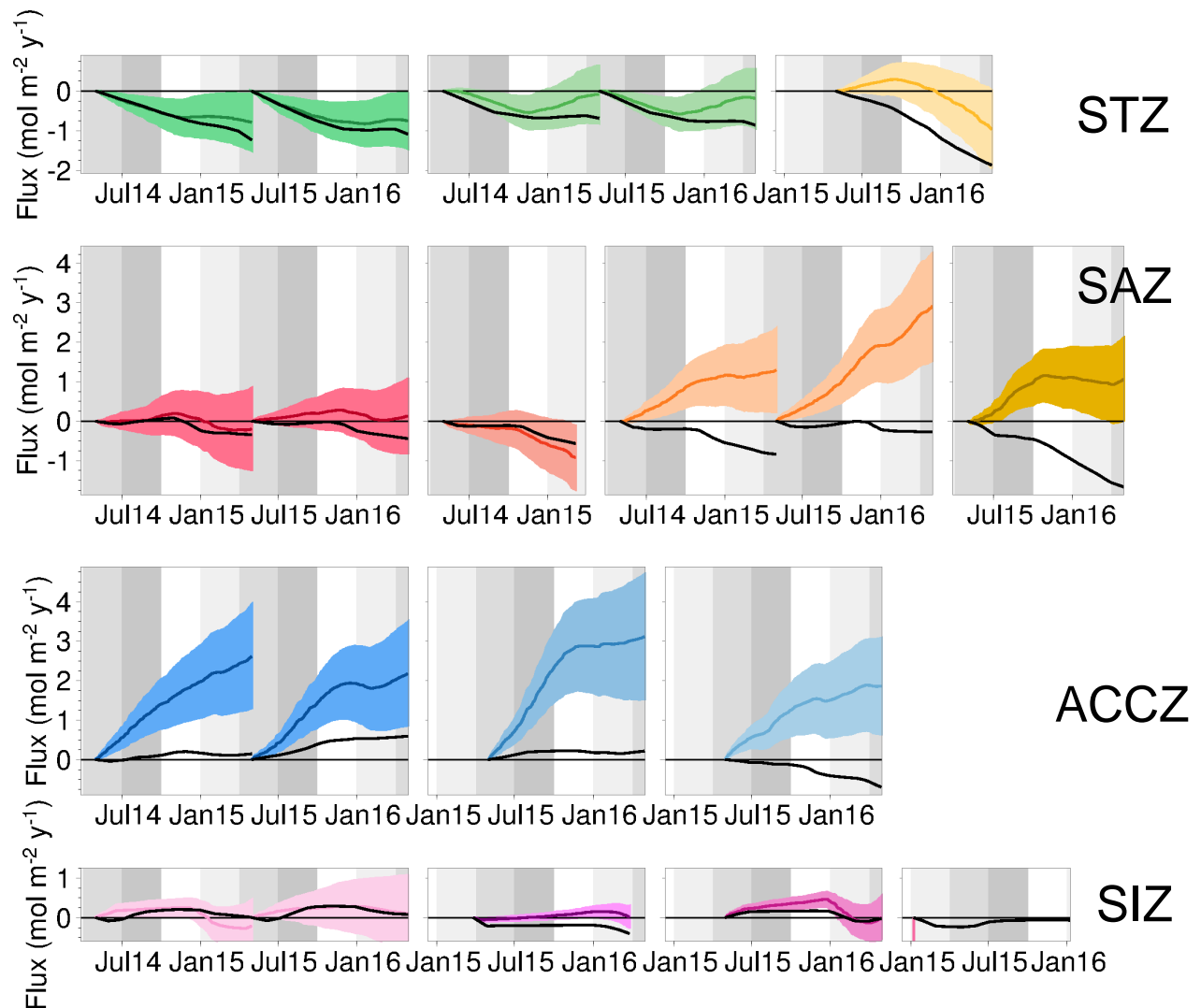


- **Note 1:** STZ 9313 crosses front
- **Note 2:** Indian Ocean Float (8514; second row, far right) is an active eddy region. Other properties - T, S, O<sub>2</sub>, NO<sub>3</sub> - indicate this float is seeing waters different than the waters sampled by SOCAT
- **Note 3:** Floats in bottom row are in SIZ and the disagreement between floats and shipbased estimates is during the ice covered season.



# Cumulative annual air-sea CO<sub>2</sub> flux comparison

- STZ
  - Agreement between estimates is good except float 9313 (yellow) which crosses front
  - Accumulates carbon when water is cold (winter & spring)
- ACCZ
  - Agreement between estimates is bad.
  - Floats show ocean losing carbon, mostly when water is cold
- SAZ and SIZ



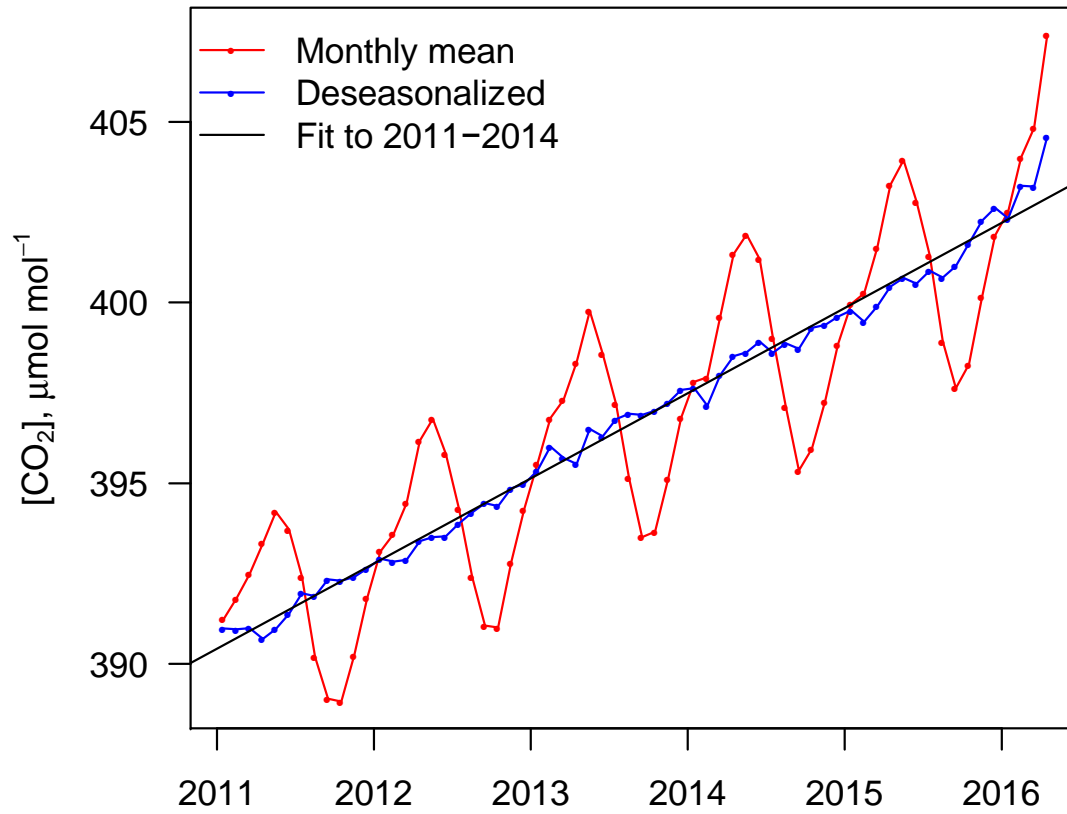
# Climate is anomalous

Gray et al., in preparation

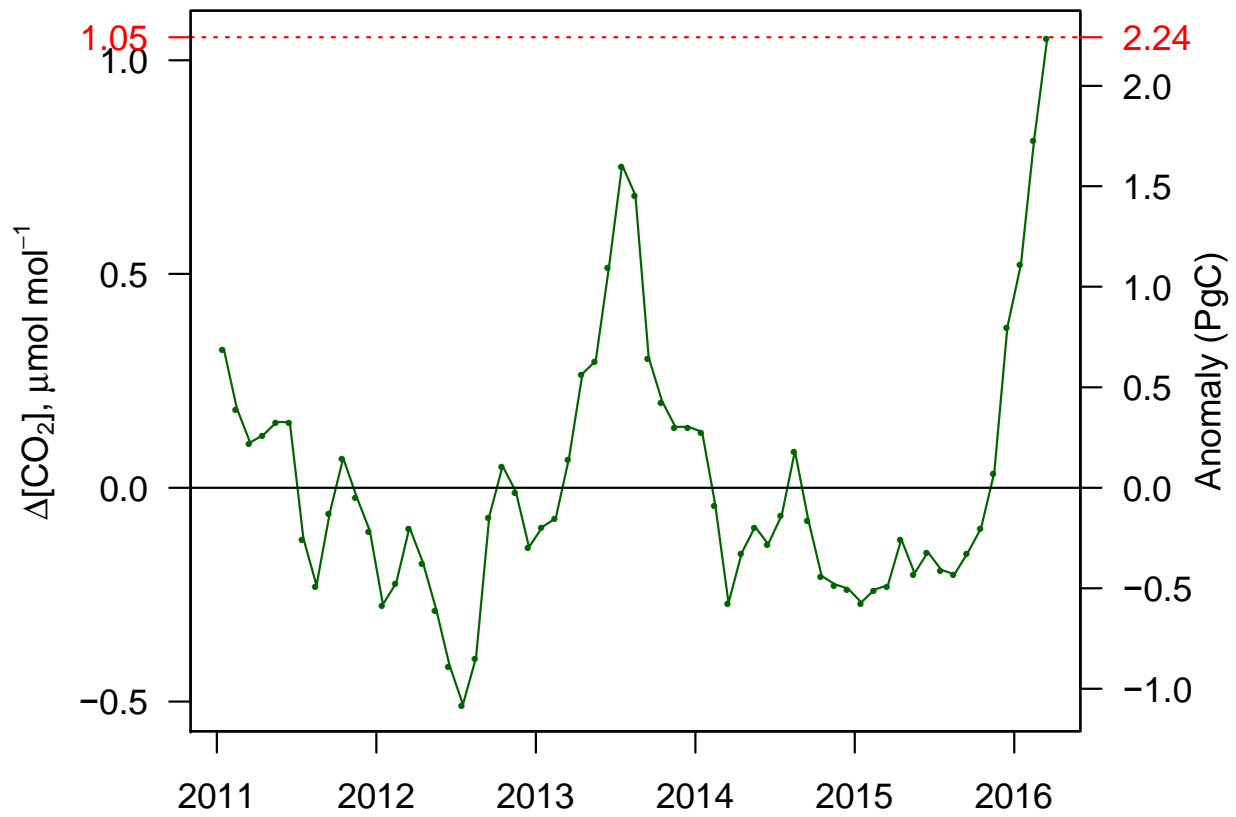
But where is all the extra carbon  
going?

Jacobson (pers. comm.)

Monthly mean CO<sub>2</sub> at Mauna Loa



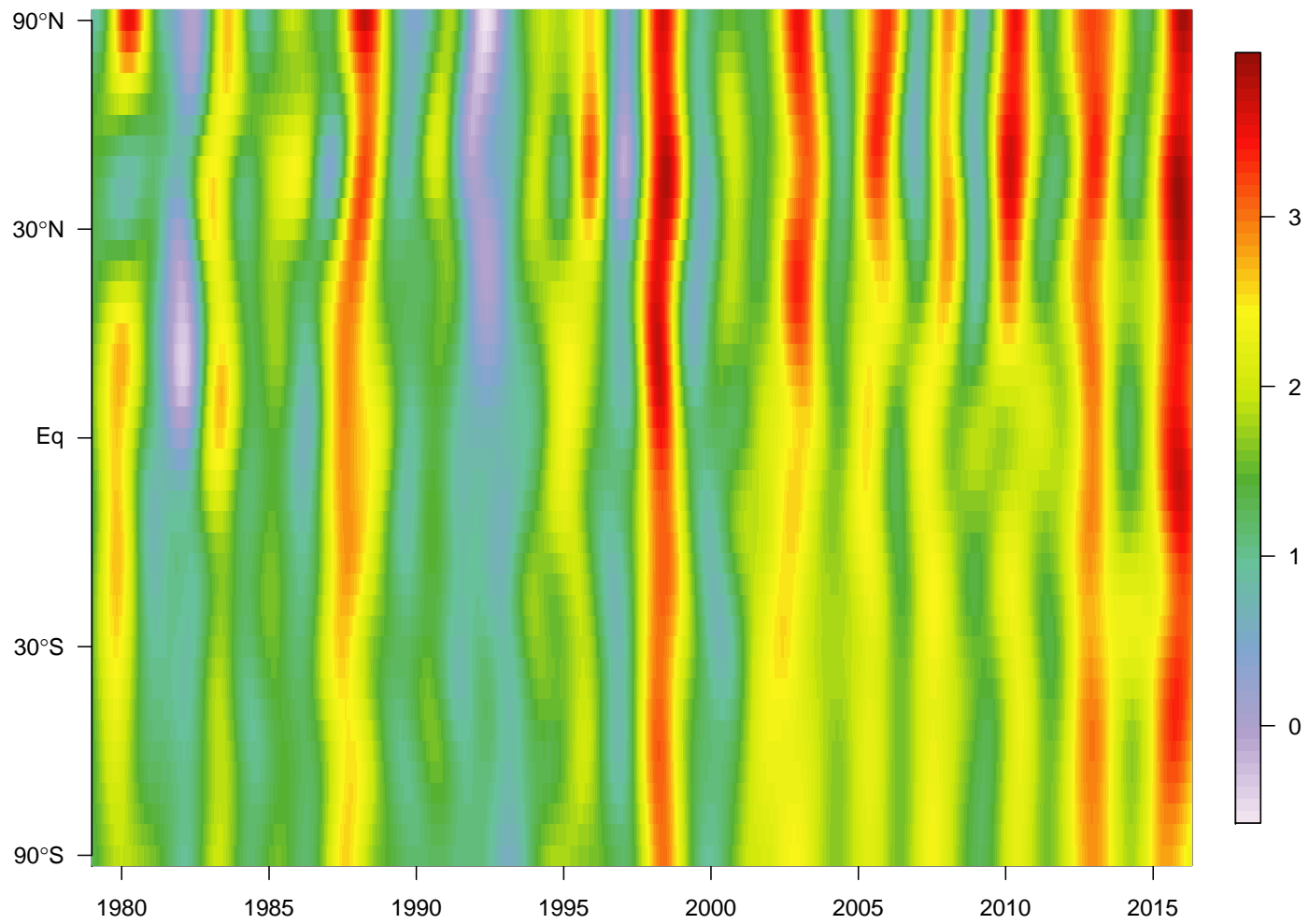
### Anomalous global MBL CO<sub>2</sub>



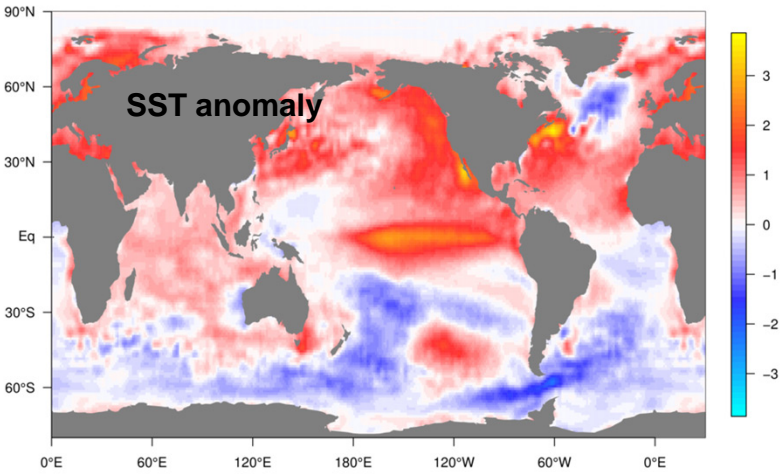
Note timing



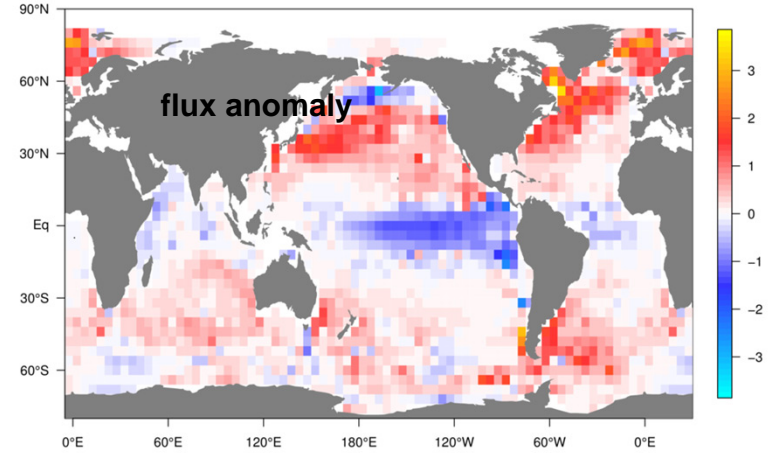
CO<sub>2</sub> zonal-mean growth rate ( $\mu\text{mol yr}^{-1}$ )



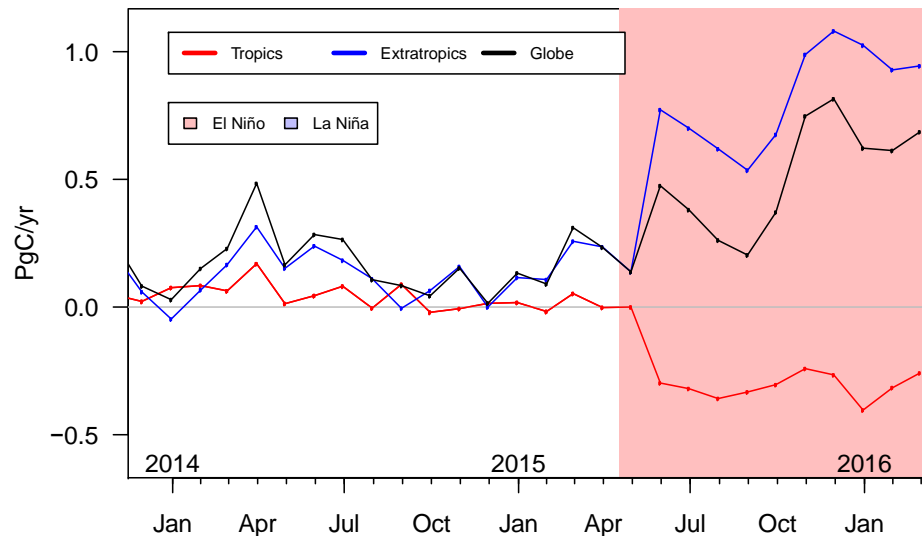
2015-09 through 2016-04 SSTA (degrees C)



2015-09 through 2016-04 AOML CO<sub>2</sub> flux anomaly (mol m<sup>-2</sup> yr<sup>-1</sup>)



AOML monthly air-sea CO<sub>2</sub> flux anomaly



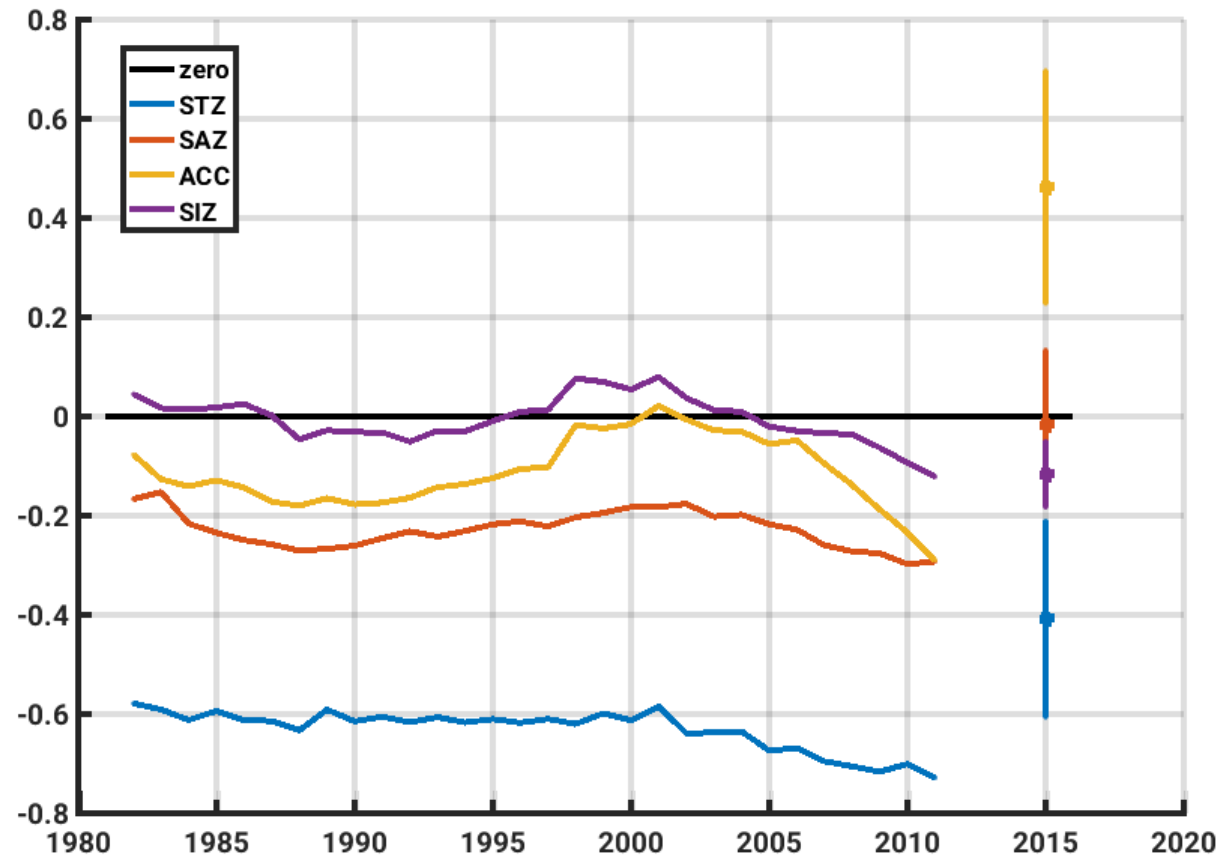
El Niño-driven anomalous CO<sub>2</sub> (PgC)

	1997-98	2015-16
Atmosphere	1.9 to 2.2 <sup>1</sup>	2.2 to 3.6 <sup>1</sup>
Oceans	-0.5 <sup>2</sup> to -0.7 <sup>3</sup>	(-0.3 to) 0.4 <sup>2</sup>

1. This work
2. NOAA AOML monthly pCO<sub>2</sub>
3. Chavez *et al.* (Science, 1999)

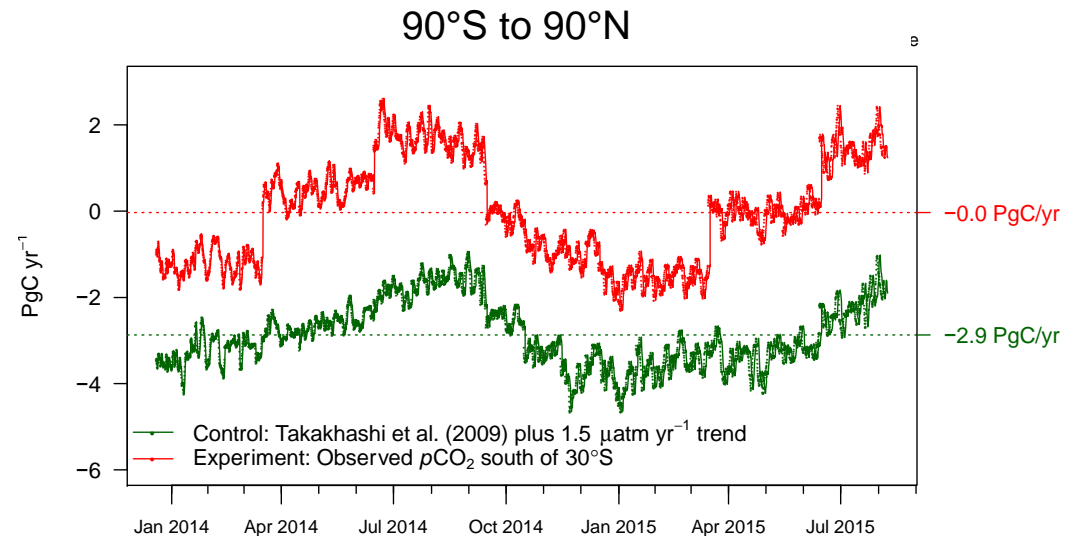
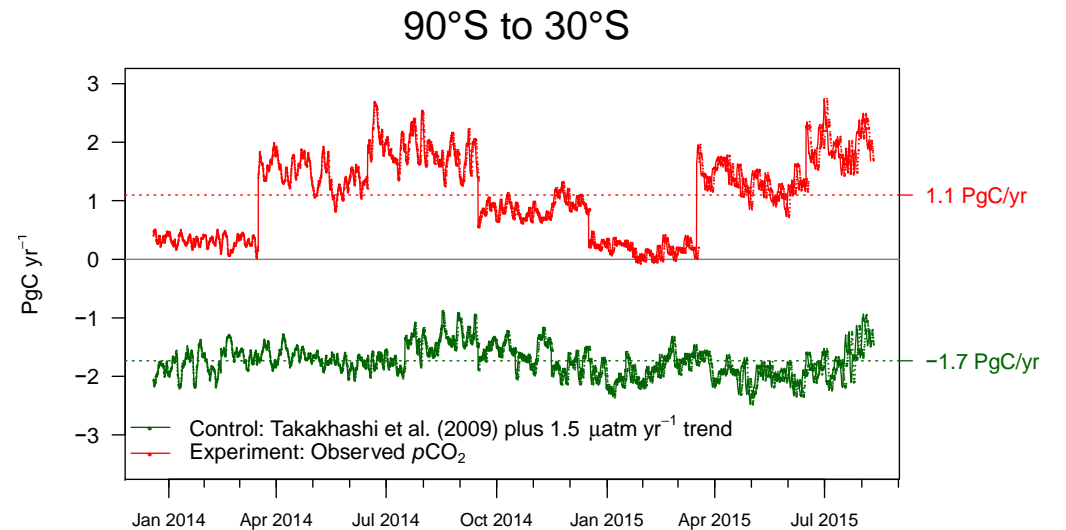
## Time history of CO<sub>2</sub> fluxes (PgC y<sup>-1</sup>)

- The 1982 to 2011 time history is from Landschützer et al. (2014)
- The individual points are from SOCCOM floats
- STZ = subtropical zone
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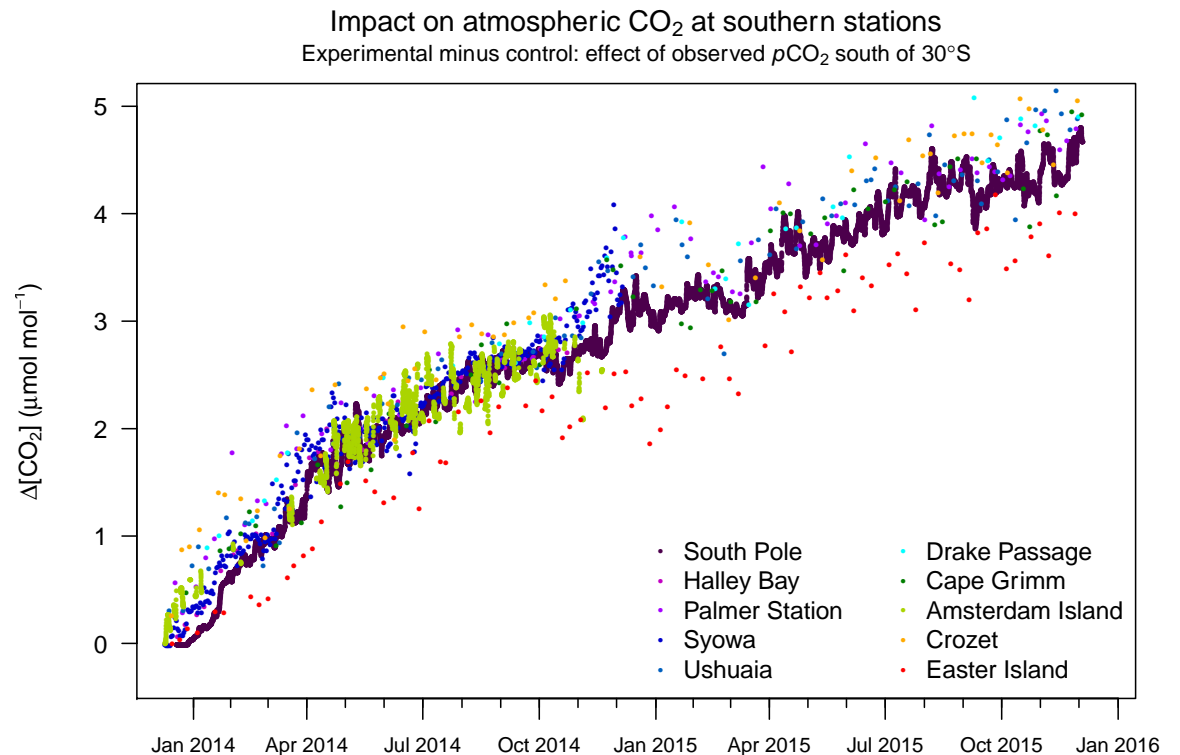
# Ocean fluxes calculated by TM5

- From Jacobson (pers. comm.)
- These fluxes are “like” priors using TM5 to compute them.
- TM5-4DVar (Transport Model 5 – Four-Dimensional Variational model), is an atmospheric transport model used for CO<sub>2</sub> flux estimation (cf. Basu et al., 2013)
- Gas exchange parameters:
  - solubility and Schmidt No. from WOA09,
  - ERA-interim winds,
  - NOAA marine boundary layer atmospheric pCO<sub>2</sub> with extrapolated trend.
- Numbers on the right-hand axis are the time-mean fluxes



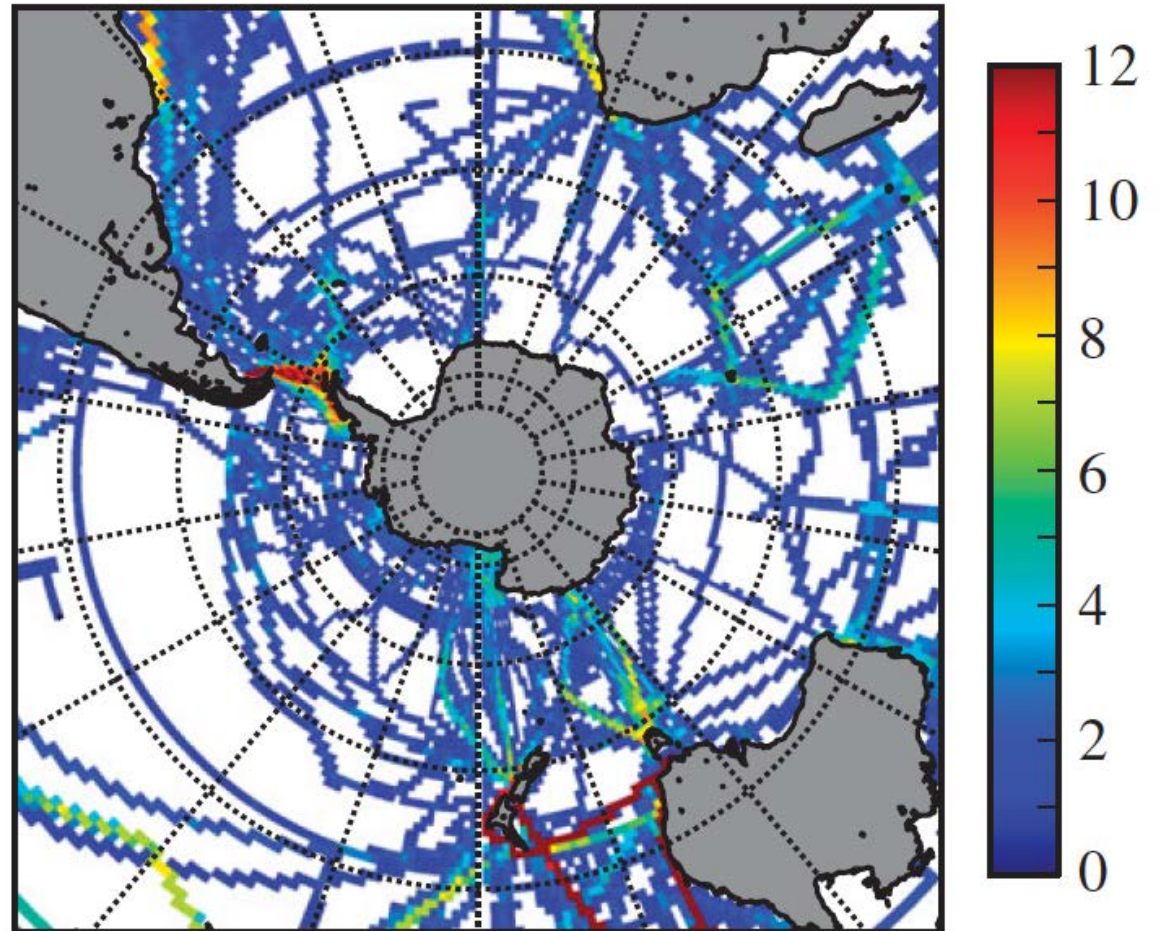
## Impact on atmospheric $p\text{CO}_2$ south of $30^\circ\text{S}$

- impact on atmospheric observations of the "experimental" air-sea fluxes in the Southern Ocean.
- The impact is cumulative over time but reaches about 3 ppm after one year.
- No evidence from CarbonTracker that we're missing this large a  $\text{CO}_2$  flux difference.
- Jacobson (pers. comm.)



## Underway sampling of $p\text{CO}_2$ from ships

- The color scale shows the months of the year with surface  $p\text{CO}_2$  measurements
- Based on all measurements between 1970 to 2011 binned in  $1^\circ$  squares.
- White = no data

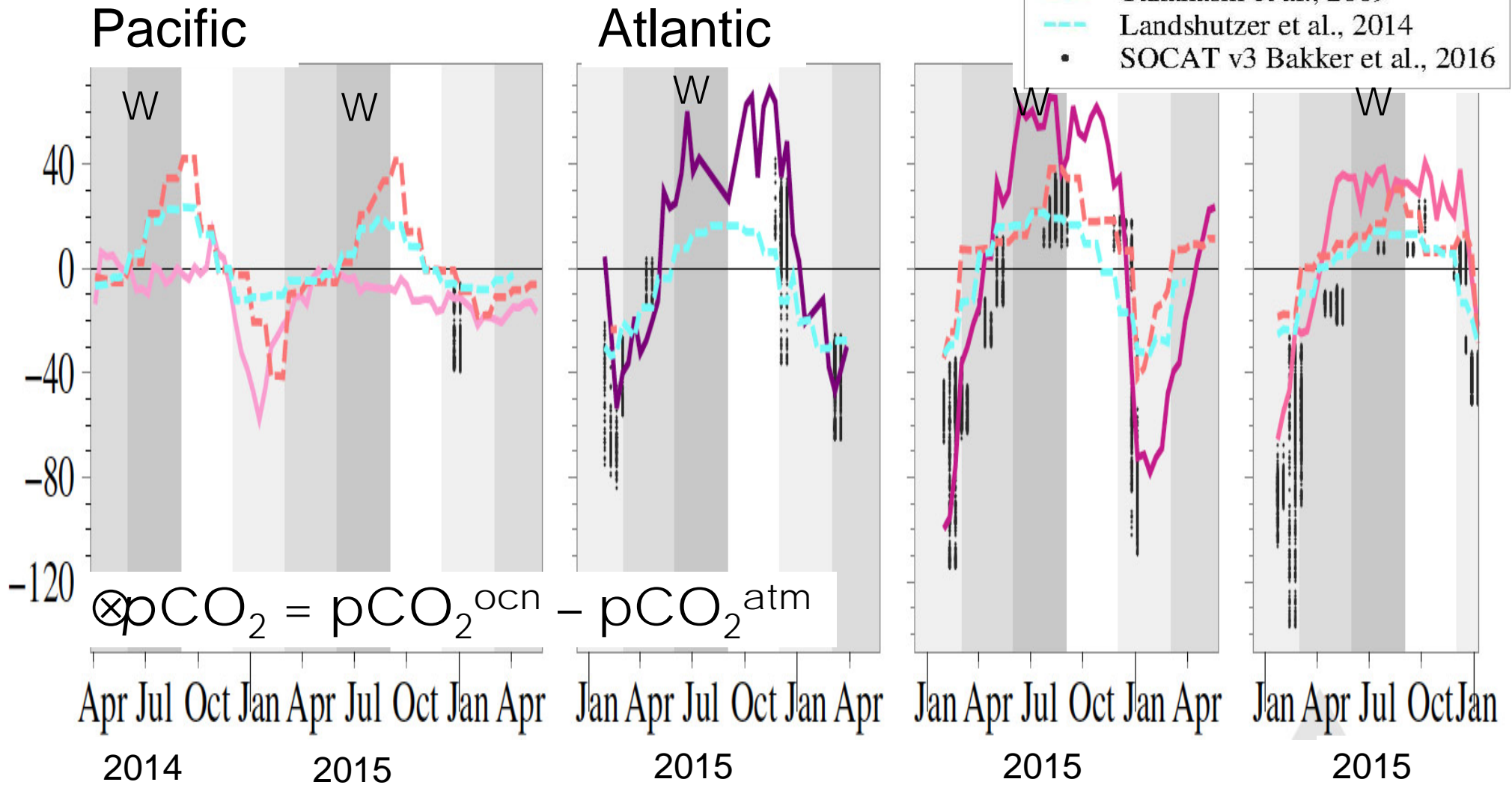




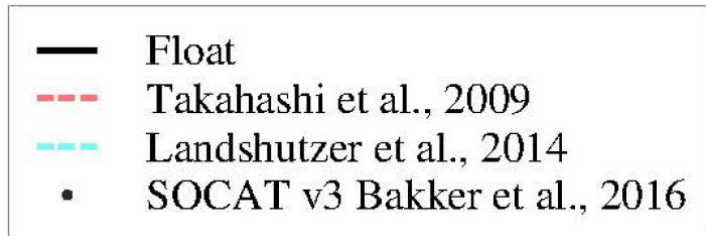
# Verification of CO<sub>2</sub> estimation

Williams et al. (in preparation)

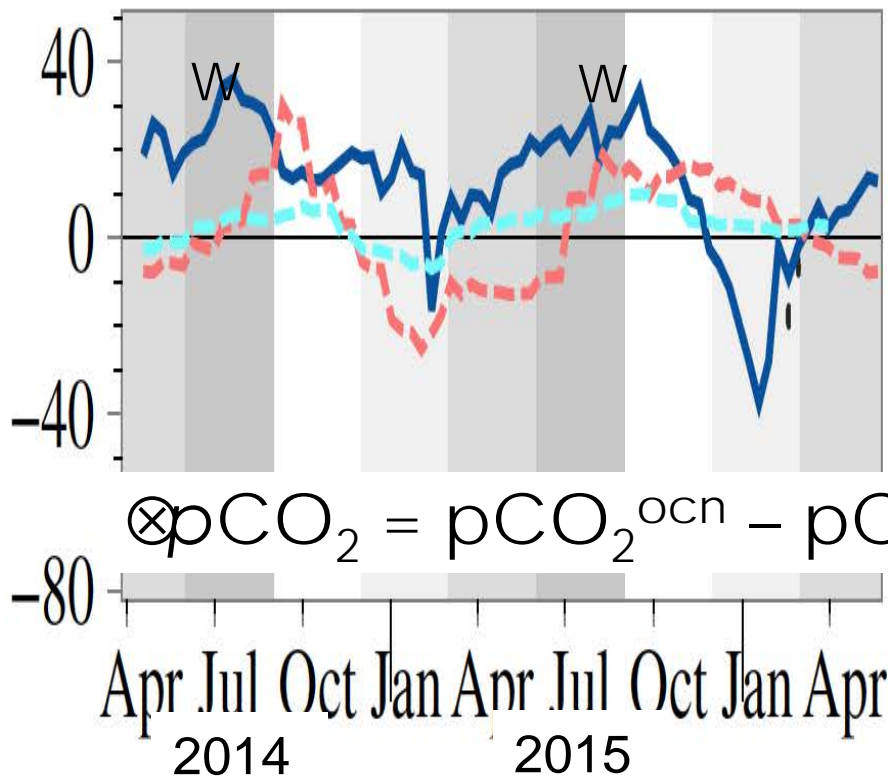
# $\Delta p\text{CO}_2$ in seasonal ice zone



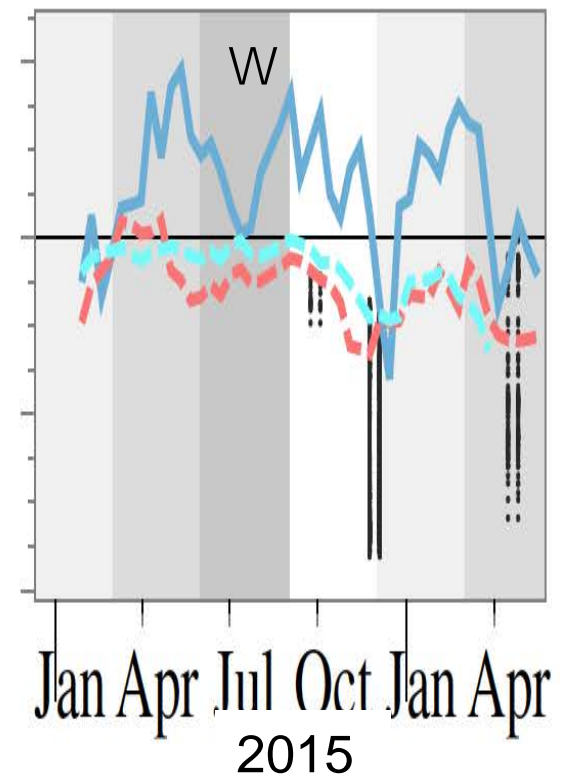
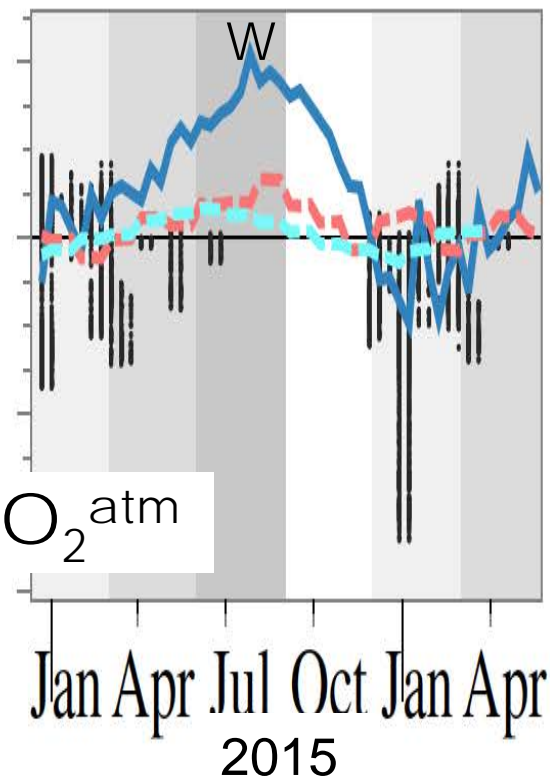
# $\Delta p\text{CO}_2$ in polar front zone



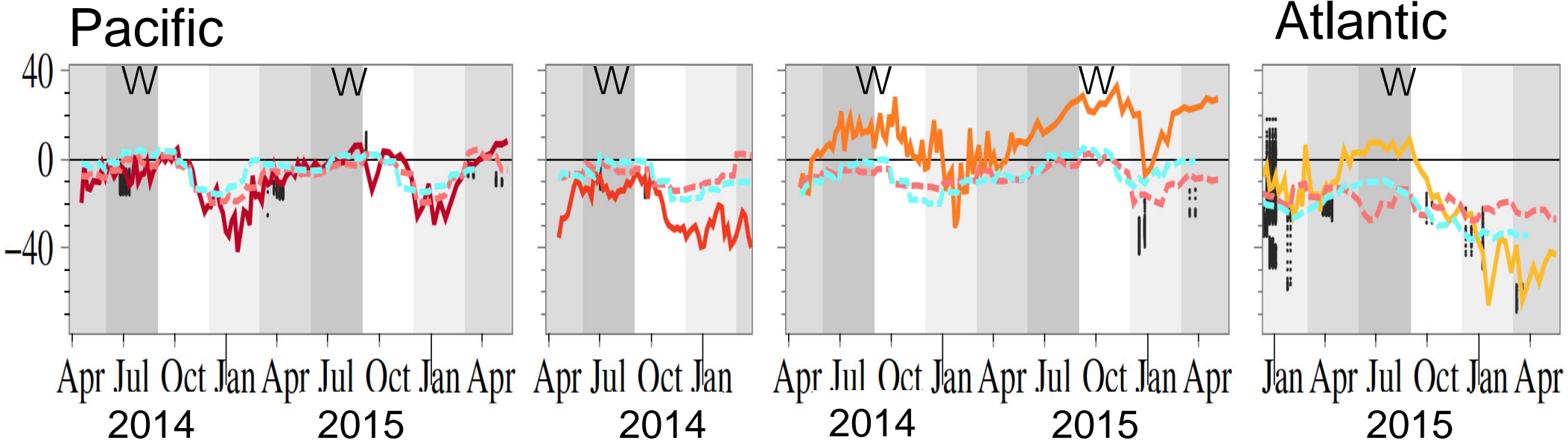
## Pacific



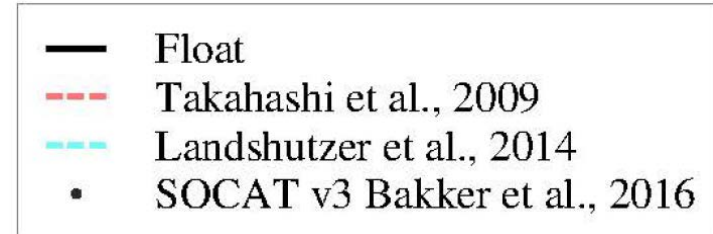
## Atlantic



# $\Delta p\text{CO}_2$ in subantarctic zone

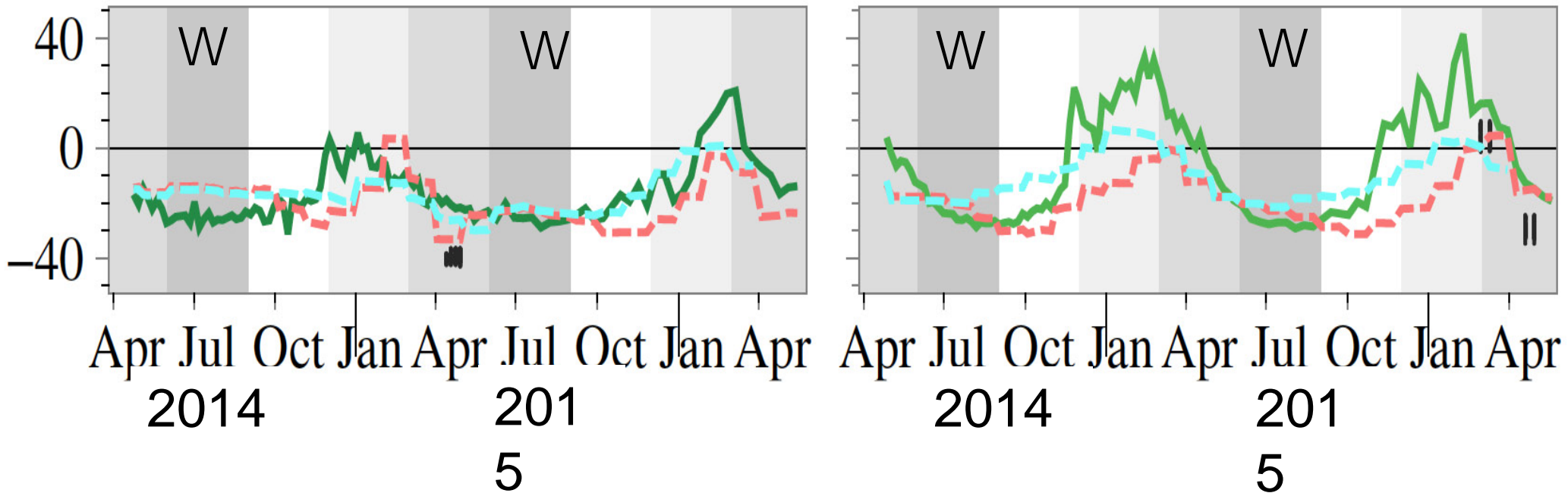


$$\Delta p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$$

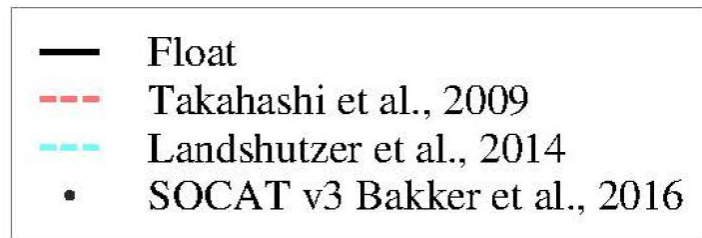


# $\Delta p\text{CO}_2$ in subtropical zone

Pacific



$$\otimes p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$$



# CO<sub>2</sub> flux estimate

We are interested in the flux of CO<sub>2</sub> out of ocean

$$F = k_g * K_0 * \Delta p\text{CO}_2$$

- $k_g$  = wind speed squared dependent gas transfer velocity from Wanninkhof (2014)
- $K_0$  = solubility constant
- $\Delta p\text{CO}_2 = p\text{CO}_2^{\text{ocn}} - p\text{CO}_2^{\text{atm}}$

Calculation of  $\Delta p\text{CO}_2$

- $p\text{CO}_2^{\text{atm}}$  from Cape Grim monthly data



## Annual net oceanic CO<sub>2</sub> uptake

- Positive indicates net outgassing, negative is uptake.
- Bold numbers are statistically different from zero.
- Italic numbers in Landschützer column are the total uptake obtained by subsampling where the float profiles are, then integrating, averaging, and multiplying by the same areas.
- The total for the Takahashi estimate is wrong. It should be -0.8, not -1.03.

Zone	Total PgC y <sup>-1</sup>	Takahashi PgC y <sup>-1</sup>	Landschützer PgC y <sup>-1</sup>
STZ	<b>-0.2 ± 0.08</b>	-0.4	-0.5 <i>-0.3</i>
SAZ	0.08 ± 0.1	-0.2	-0.2 <i>-0.2</i>
ACCZ	<b>0.8 ± 0.2</b>	-0.2	-0.07 <i>0.02</i>
SIZ	0.03 ± 0.05	0.02	-0.02 <i>-0.02</i>
	<b>0.7 ± 0.2</b>	-1.03	-0.8 <i>-0.5</i>

## CO<sub>2</sub> fluxes (mol m<sup>-2</sup> y<sup>-1</sup>)

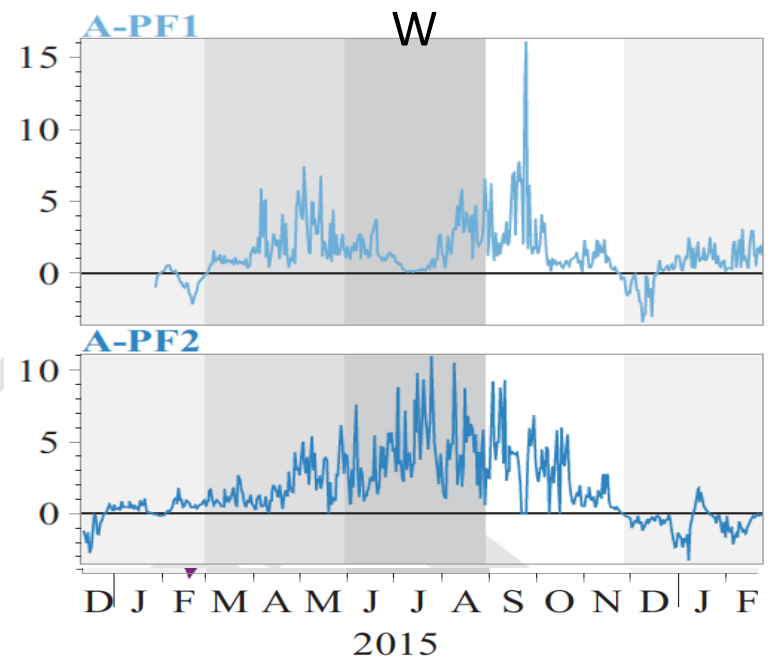
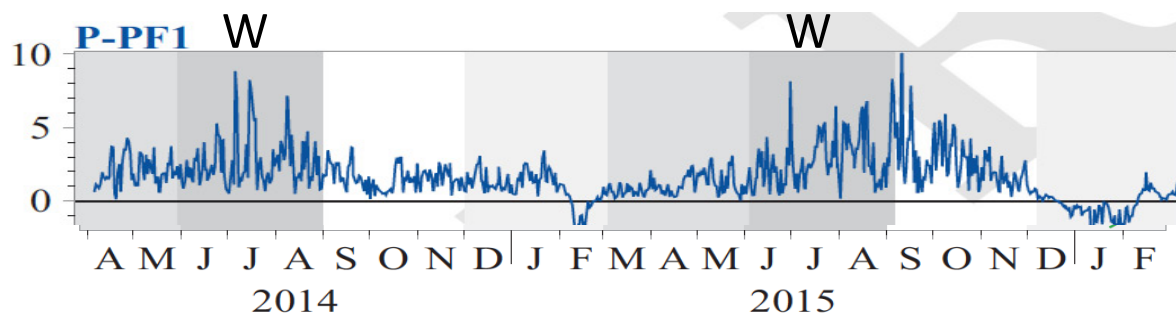
Pacific		Atlantic	
Subtropics (Significant uptake)			
P-ST1	<b>-0.9 ± 0.7</b>		
P-ST2	<b>-1.6 ± 0.7</b>		
Subantarctic Zone			
P-SA1	-0.1 ± 1.1	A-SA1	-0.8 ± 0.9
P-SAA3	-1.0 ± 1.1		
Polar Front Zone (Significant degassing)			
P-PF1	<b>1.6 ± 1.1</b>	A-PF1	<b>1.5 ± 1.2</b>
		A-PF2	<b>2.2 ± 1.3</b>
Seasonal Ice Zone			
P-SI1	-0.6 ± 0.6	A-SI1	0.2 ± 0.7
		A-SI2	-0.7 ± 0.3
		A-SI3	-0.3 ± 0.2

For scale, the average anthropogenic CO<sub>2</sub> uptake is  $0.46 \pm 0.14$  mol m<sup>-2</sup> y<sup>-1</sup>

# CO<sub>2</sub> fluxes in Polar Front zone (mol m<sup>-2</sup> y<sup>-1</sup>)

	Pacific		Atlantic	
y <sup>-1</sup>	P-PF1	1.6 ± 1.1	A-PF1	1.5 ± 1.2
			A-PF2	2.2 ± 1.3

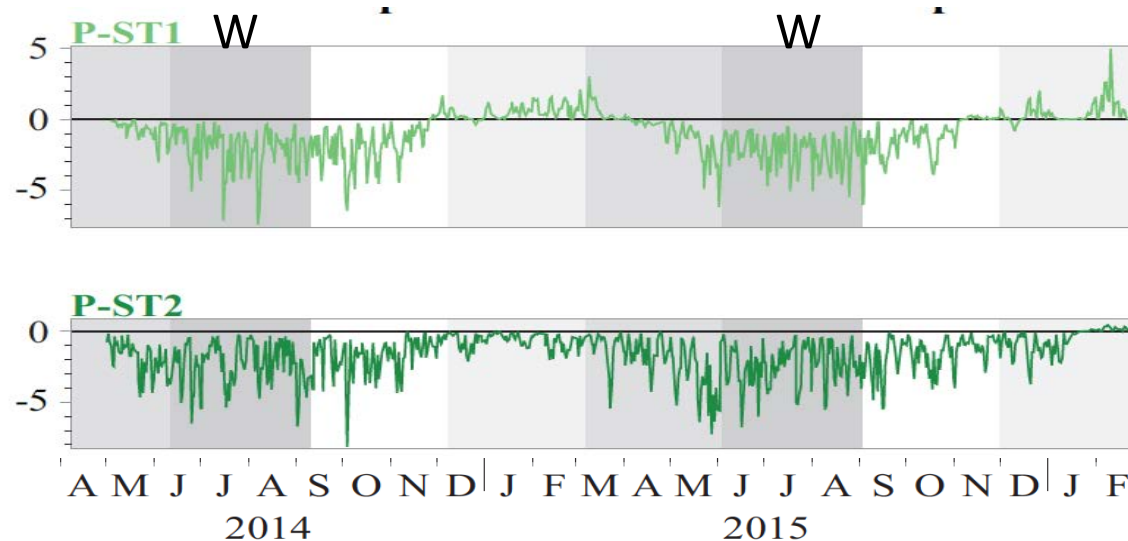
Anthropogenic CO<sub>2</sub> uptake is 0.46 ± 0.14 mol m<sup>2</sup> y<sup>-1</sup>



# CO<sub>2</sub> fluxes in subtropics (mol m<sup>-2</sup> y<sup>-1</sup>)

Pacific subtropics	
P-ST1	-0.9 ± 0.7
P-ST2	-1.6 ± 0.7

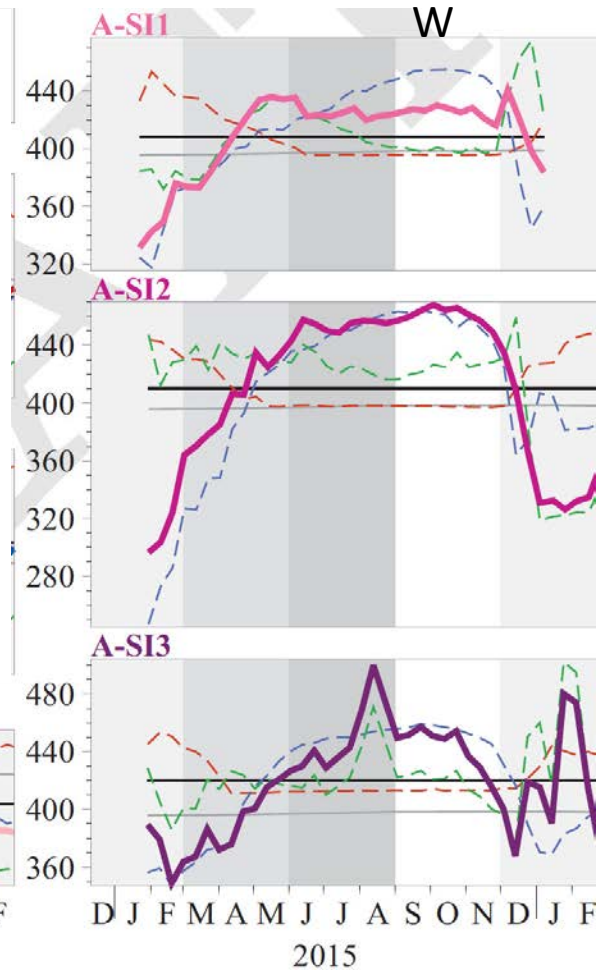
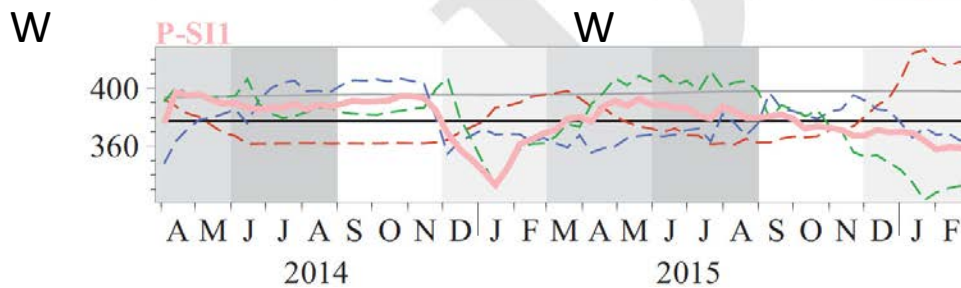
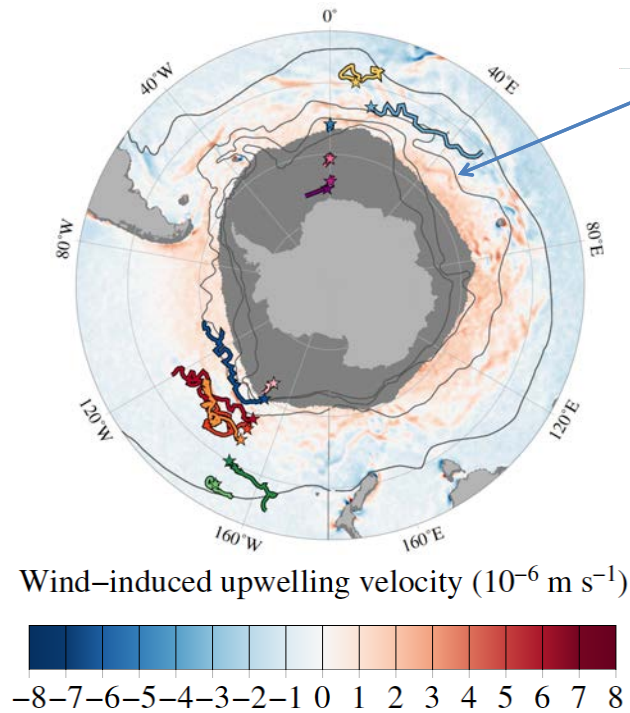
Anthropogenic CO<sub>2</sub> uptake is 0.46 ± 0.14 mol m<sup>2</sup> y<sup>-1</sup>



# Causes of CO<sub>2</sub> seasonality

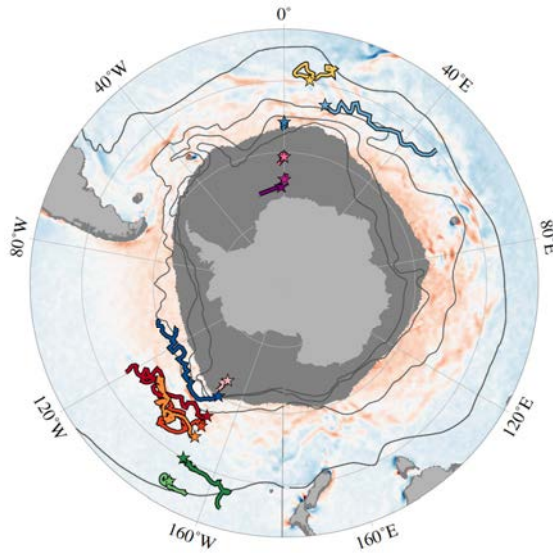
Gray et al. (in preparation)

# pCO<sub>2</sub> in seasonal ice zone

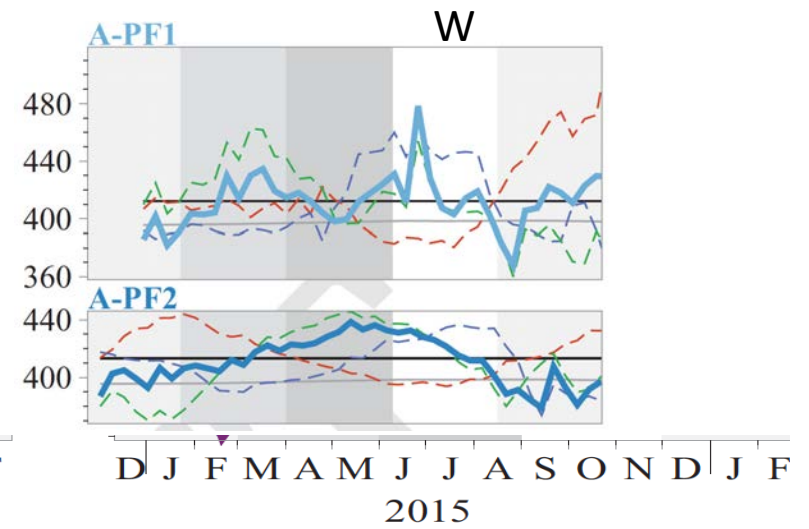
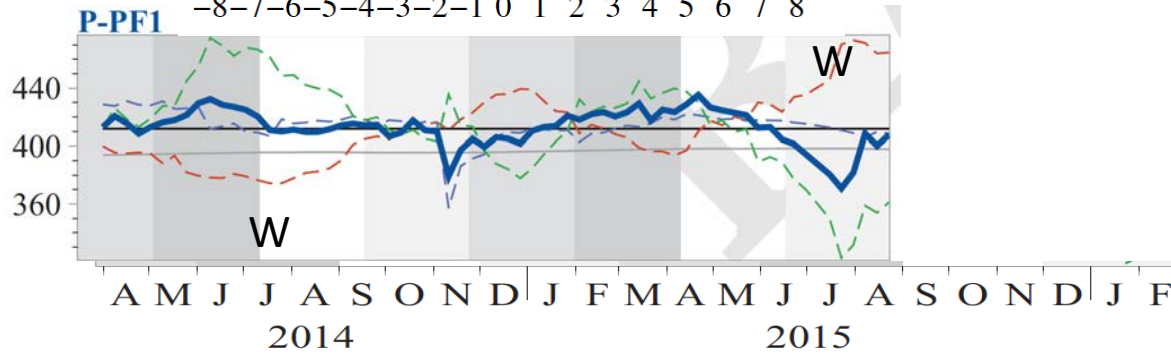
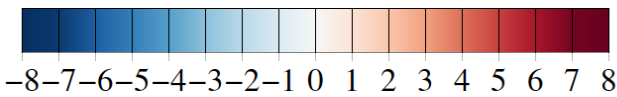




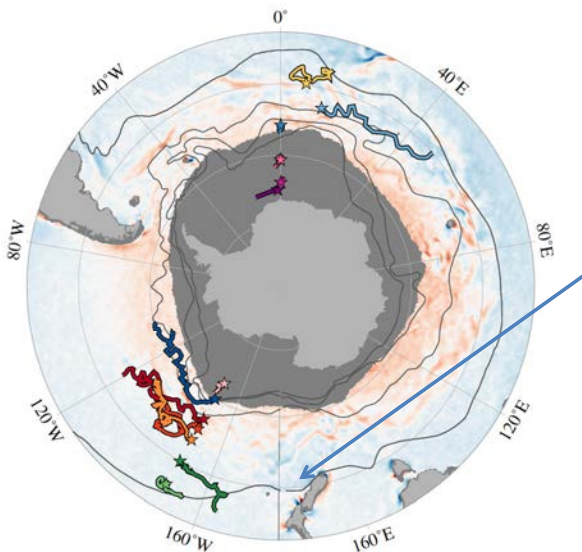
# pCO<sub>2</sub> in Polar Front zone



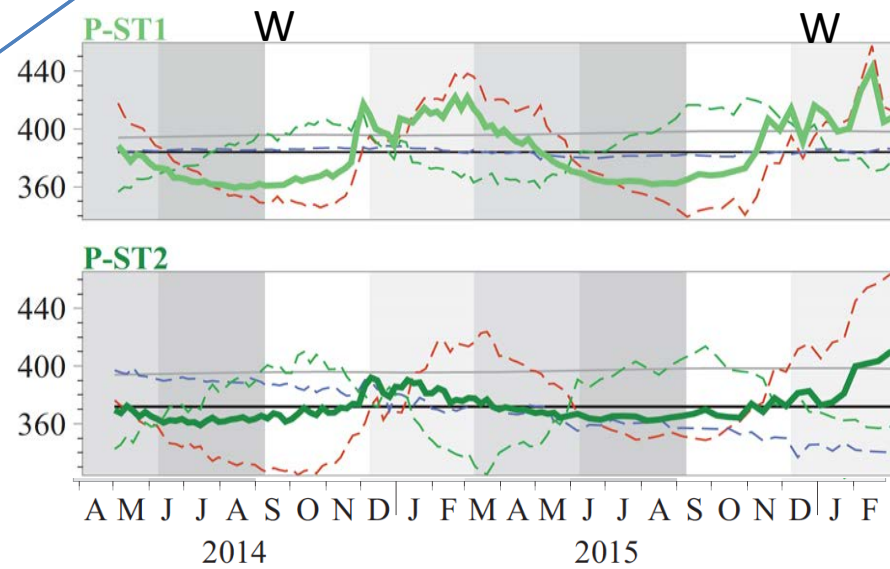
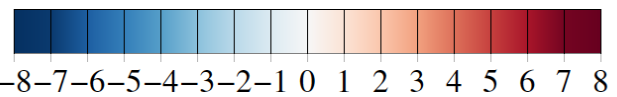
Wind-induced upwelling velocity ( $10^{-6} \text{ m s}^{-1}$ )



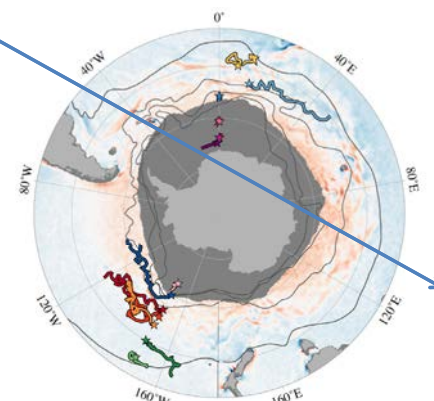
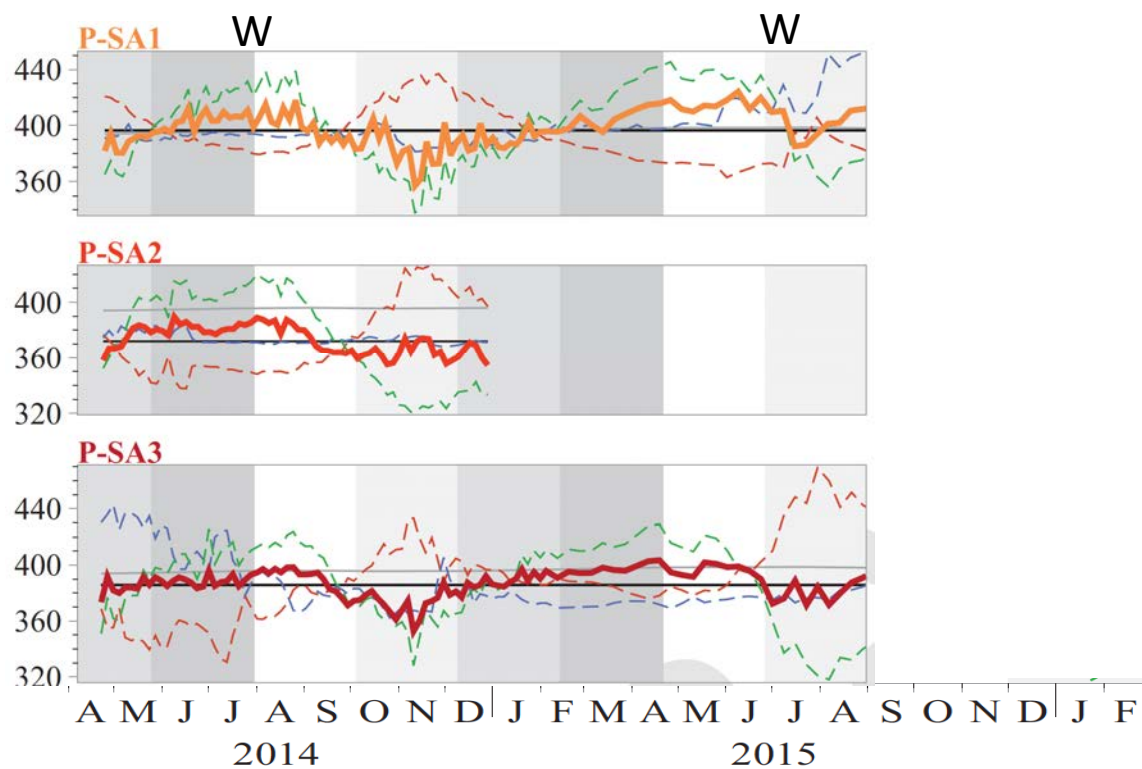
# pCO<sub>2</sub> in subtropical gyre (Pacific only)



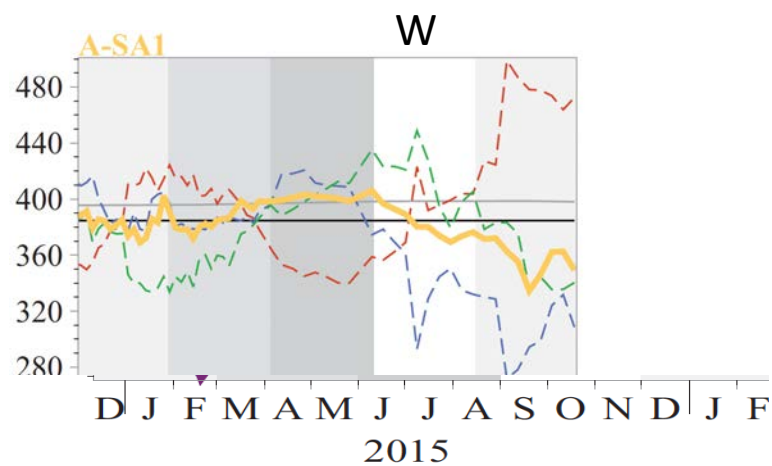
Wind-induced upwelling velocity ( $10^{-6} \text{ m s}^{-1}$ )



# pCO<sub>2</sub> in SubAntarctic Front zone



Wind-induced upwelling velocity (10<sup>-6</sup> m s<sup>-1</sup>)



# The role of the Southern Ocean in the global climate system.

1

It accounts for **67-98%** of the excess heat that is transferred from the atmosphere into the ocean each year.

2

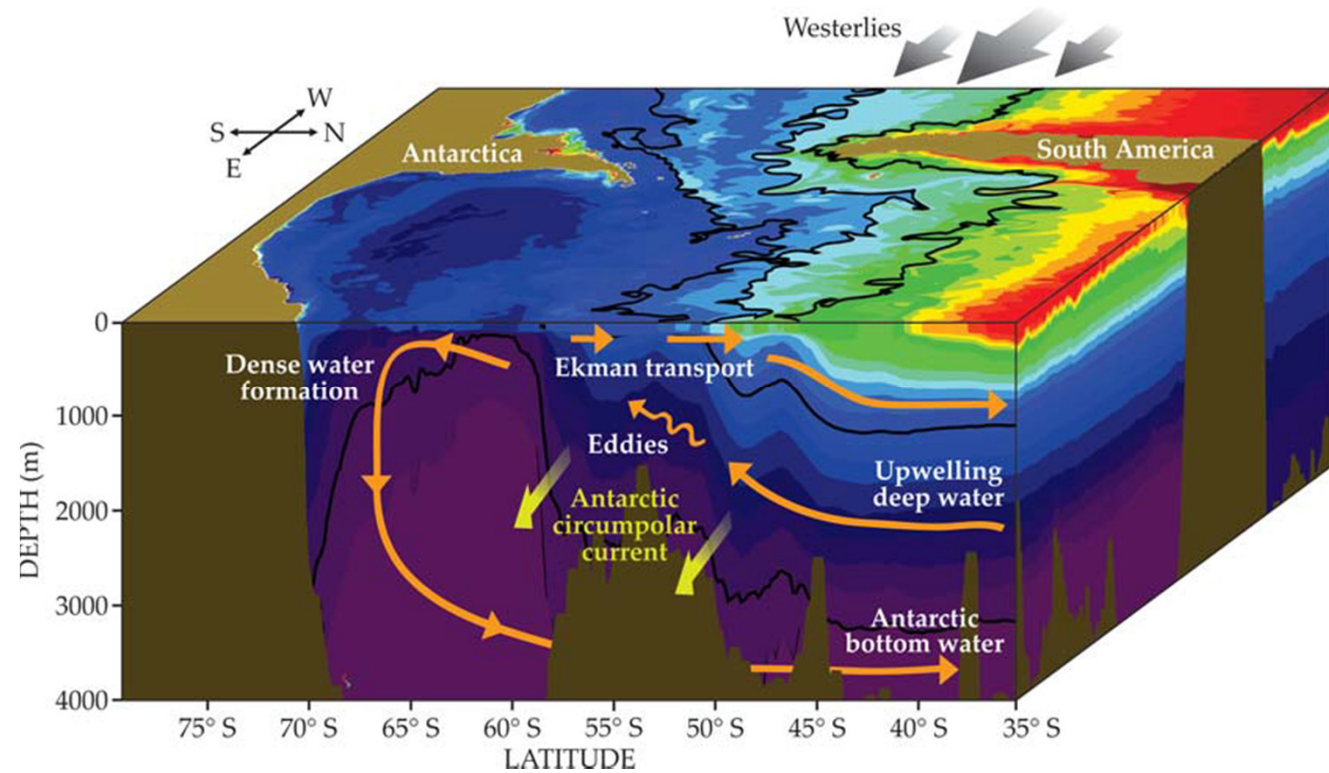
It accounts for **up to half** of the annual oceanic uptake of anthropogenic carbon dioxide from the atmosphere.

3

Vertical exchange in the Southern Ocean is responsible for supplying nutrients that fertilize **three-quarters** of the biological production in the global ocean north of 30°S.

# Why is the Southern Ocean so important?

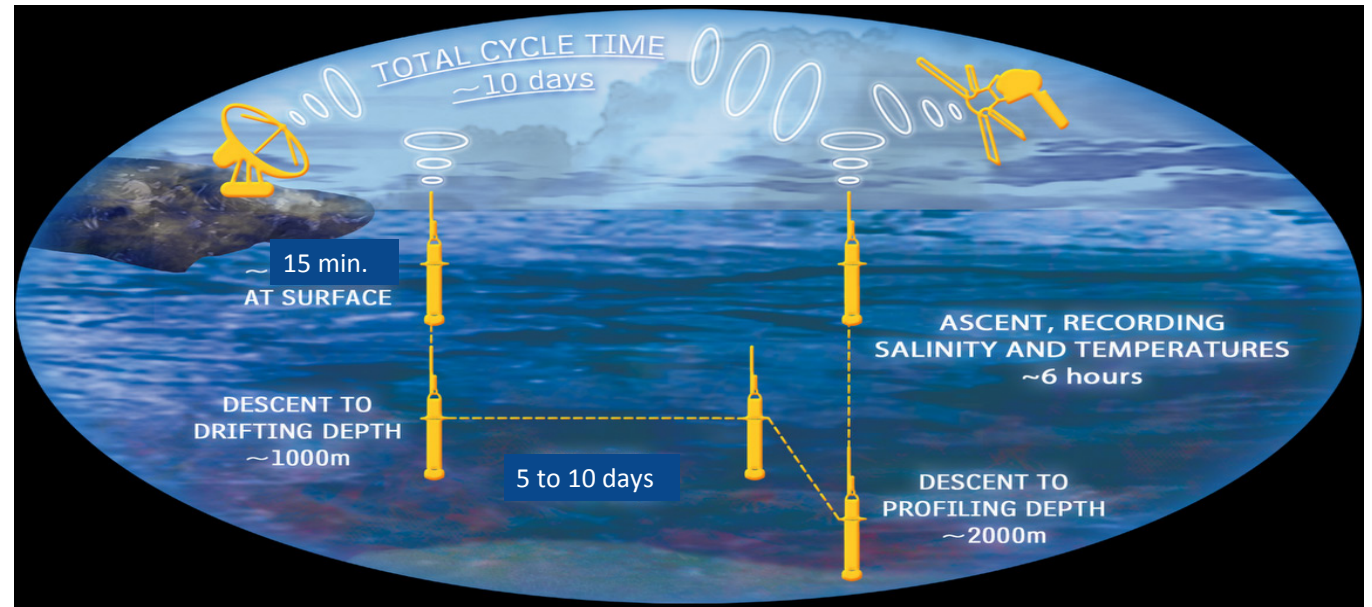
- “Window to the deep ocean”
- Southern Ocean is the only place where there is direct upwelling from very deep waters to the sea surface over a very large region



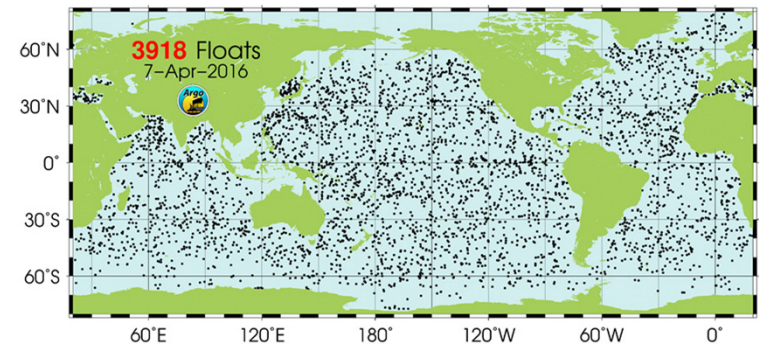
Morrison et al. (2015) based on Talley & Oibers (2014, after Speer et al. 2000, and NRC 2011)



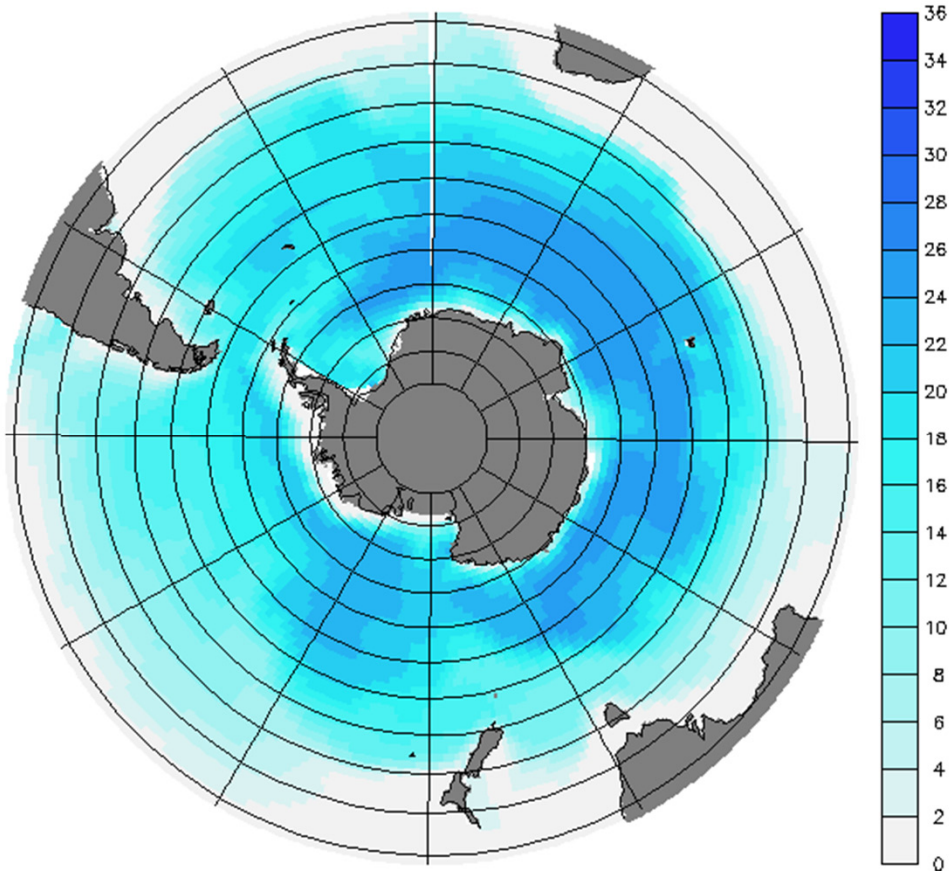
# Argo Floats



- Argo profiling floats
  - have a 4 to 7 year lifetime,
  - Measure T & S from ~2000 m to the surface every 5 to 10 days.
  - data direct to Internet.



# Undersampling of Nitrate



Nitrate ( $\mu\text{mol/kg}$ ): Year/Month: 1983/01  
JFM = white, AMJ = black, JAS = black, OND = white

Animation of actual Southern Ocean nitrate observations

- Background is World Ocean Atlas monthly climatology.
- End of animation shows additional measurements that would have been made if Argo floats had nitrate sensors

Courtesy of R. Slater, Princeton U.



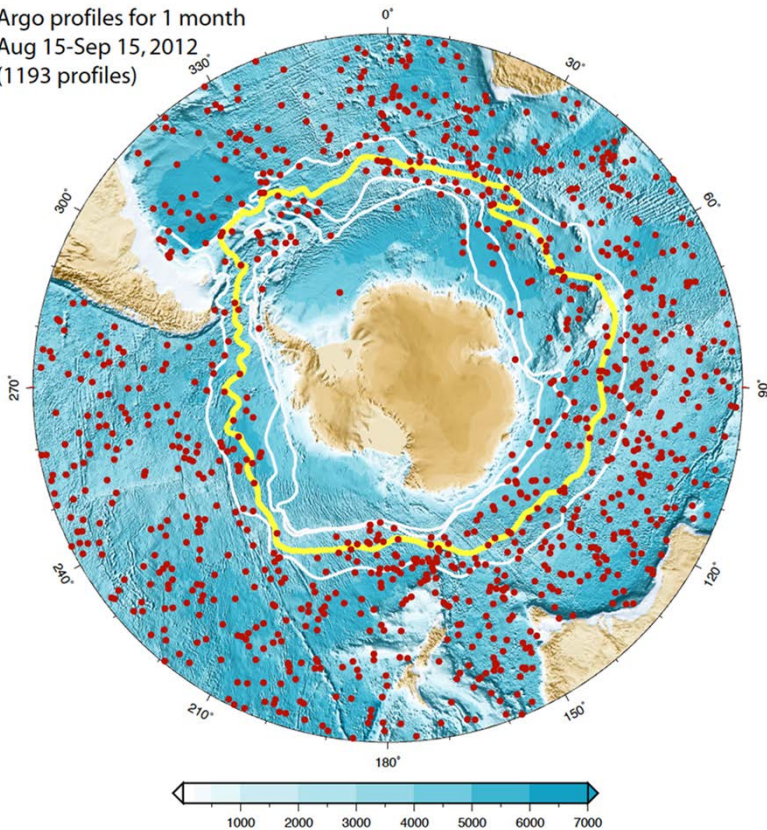
# The Opportunity

## (1) A paradigm shift – Transformative observing system

- Argo floats (currently required to measure only temperature & salinity)

- Biogeochem sensors now developed for
  - pH
  - Nitrate,
  - oxygen, and
  - optics (FLBB) – *funded by NASA*

Argo profiles for 1 month  
Aug 15-Sep 15, 2012  
(1193 profiles)

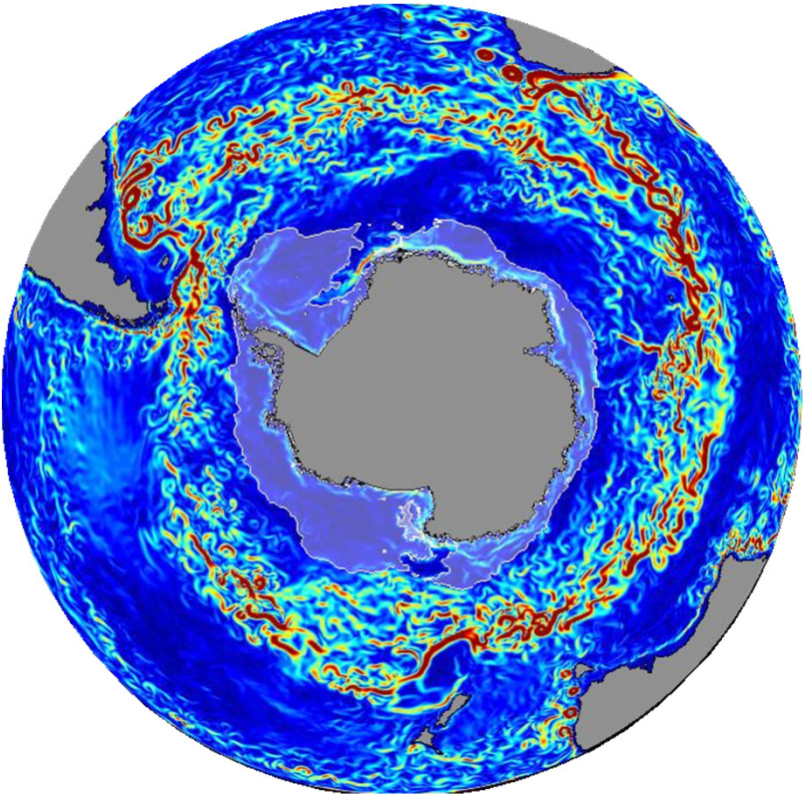


**SOCCOM**

# The Opportunity

## (2) Transformative observational analysis methods

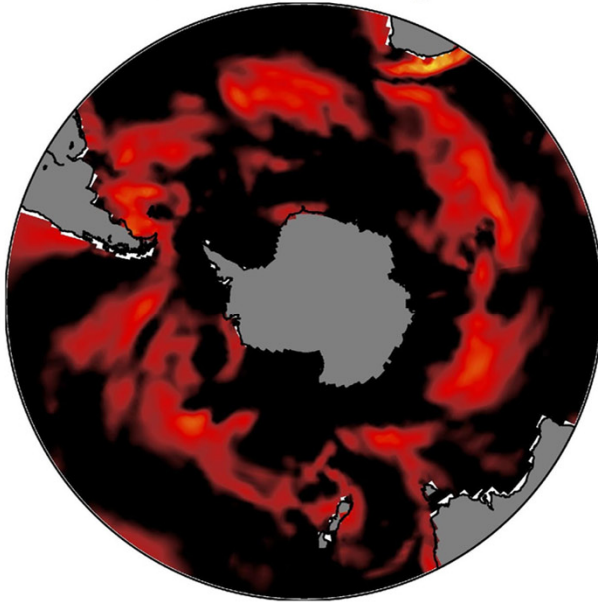
- Southern Ocean State Estimation (SOSE) using data fitting to produce full 4D estimates of ocean properties



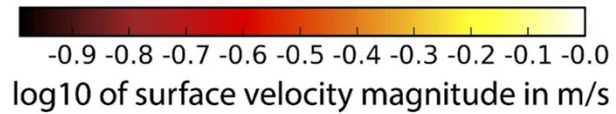
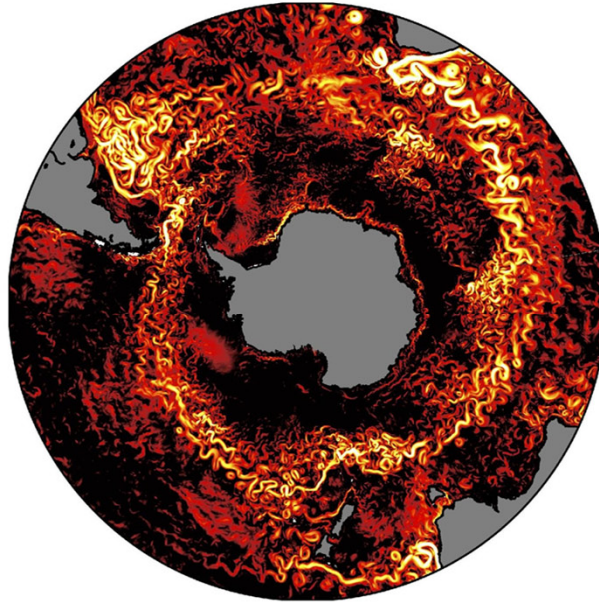
**SOCCOM**

# The Opportunity

1° ocean resolution  
(GFDL CM2-1deg)



0.1° ocean resolution  
(GFDL CM2.6)



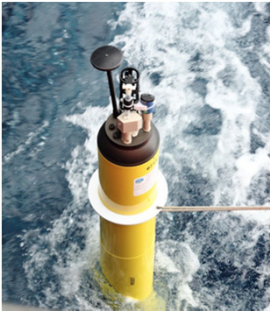
(3) Eddy rich-high resolution climate models



**SOCCOM**

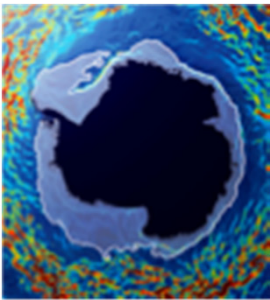


# What is SOCCOM?



SOCCOM's mission is to drive a transformative shift in our understanding of the role of the Southern Ocean in climate change and biogeochemistry by:

Extending sparse Southern Ocean biogeochemical observations by deploying a robotic observing system composed of ~200 autonomous BGC Argo floats that will provide nearly continuous coverage in time and horizontal space over the entire Southern Ocean, as well as vertical coverage deep into the water column.



Using these observations to analyze and improve a new generation of high resolution ( $1/10^\circ$ ) earth system models to both increase our understanding of the Southern Ocean's current workings and make better projections of the future trajectory of the Earth's climate and biogeochemistry.



Educating a new generation of ocean scientists trained in both ocean observation and simulation, and develop a sophisticated outreach effort to disseminate results to the broadest possible community.

# Directorate

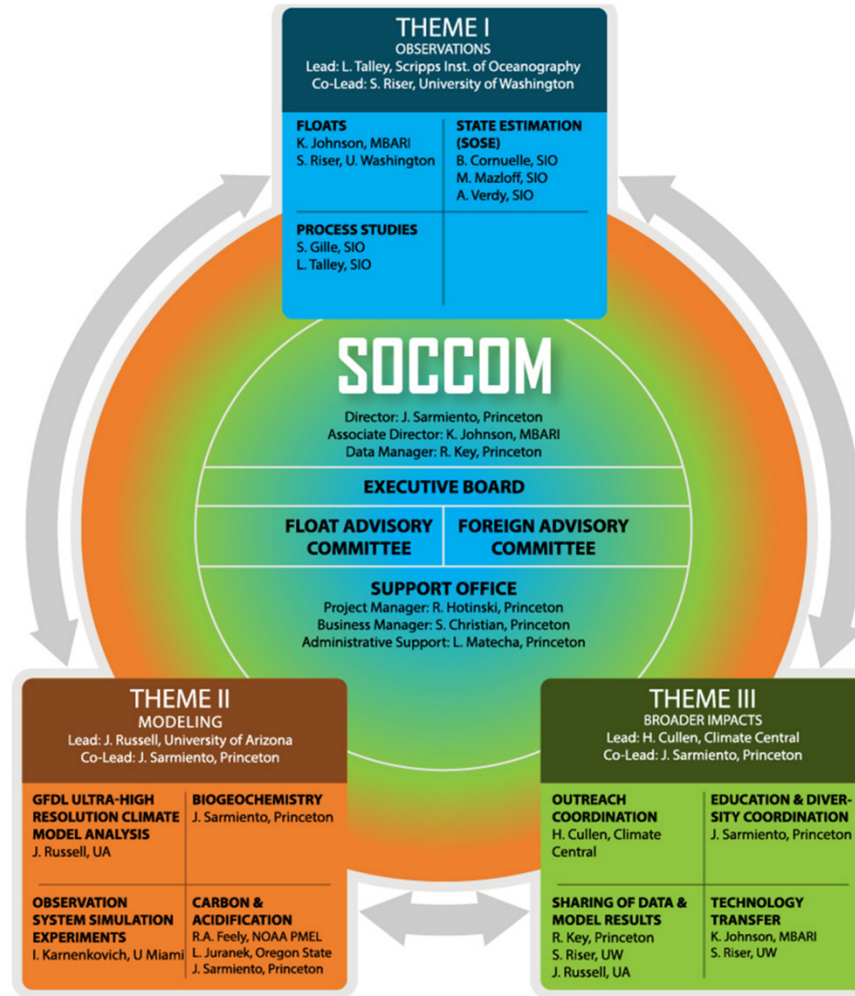


**Director**  
Jorge Sarmiento,  
Princeton

**Associate  
Director**  
Ken Johnson,  
MBARI



**Project  
Manager**  
Roberta  
Hotinski,  
Princeton



# Executive Board

## Theme I Observations



Lynne Talley, UCSD



Steve Riser, U.  
Washington

## Theme II Modeling



Joellen Russell  
U. Arizona

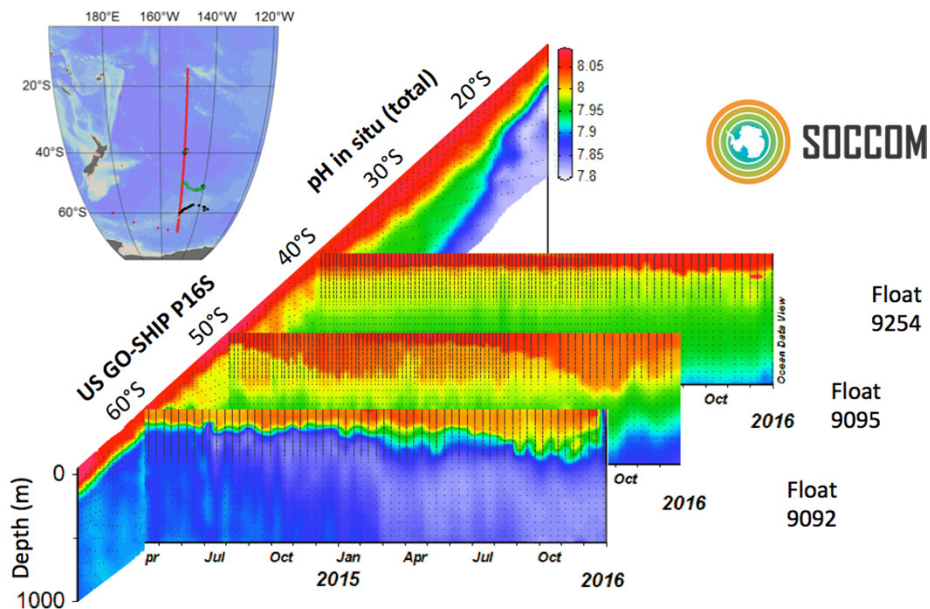
## Broader Impacts

Heidi Cullen,  
Climate Central



**SOCCOM**

# SOCCOM float observations



Extension of decadal-scale GO-SHIP carbon observations into the seasonal and interannual domain

Some details:

- (1) Calibration – goal is to produce climate quality data
- (2) All data immediately available on the web
- (3) Ice avoidance software
- (4) Southern Ocean State Estimate (SOSE):
  - Least squares fit to data by the adjoint method (M. Mazloff as part of ECCO consortium funded by the NSF)
  - Earlier development of BGC model was supported by NASA CMS
  - Nominal resolution is  $1/6^\circ$