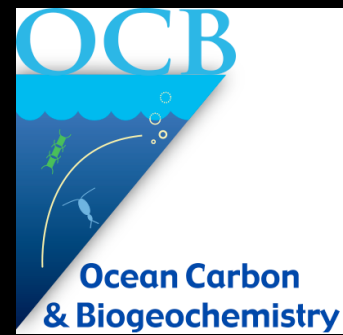




An International Study of the Marine Biogeochemical Cycles of Trace Elements and their Isotopes



Modelling ...

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GEOTRACES / OCB Workshop 1-4 August 2016, Lamont

Overarching Goal

- Set the scene for discussion of how models may contribute to a better understanding of TEI internal cycling
 - How do models represent key TEI internal cycling processes?
 - What are their assumptions?
 - How can we make progress?

Disclaimer(s) ...

- I will focus on:
 - General circulation style models
 - Nested within a physical framework
 - Important caveats associated with resolution
 - Models appropriate for decadal to centennial time scales
 - Resolving seasonality (resolving processes on hourly time scales)
 - Not box or inverse models
 - Processes relevant for bioactive TEs

So how do these models work?

```

zTL1(ji,jj,jk) = ztotlig(ji,jj,jk)
zkeq = fekeq(ji,jj,jk)
zfsatur = zTL1(ji,jj,jk) * 1E-9
ztf = trb(ji,jj,jk,jpfer)
! Fe' is the root of a 2nd order polynomial
zFe3(ji,jj,jk) = (- (1. + zfsatur * zkeq - zkeq * ztf)
&
+ SQRT( (1. + zfsatur * zkeq - zkeq * ztf)**2
&
+ 4. * ztf * zkeq) ) / ( 2. * zkeq )
zFe3(ji,jj,jk) = zFe3(ji,jj,jk) * 1E9
zFeL1(ji,jj,jk) = MAX(0., trb(ji,jj,jk,jpfer) * 1E9 - zFe3(ji,jj,jk))
END DO
END DO
ENDIF

! if no dust available
zdst = 0.
ICDIR NOVERRCHK
DO jk = 1, jpkml
ICDIR NOVERRCHK
DO jj = 1, jpj
ICDIR NOVERRCHK
DO ji = 1, jpi
zstep = xstep
# if defined key_degrad
zstep = zstep * facvcl(ji,jj,jk)
# endif

! Scavenging rate of iron. This scavenging rate depends on the load of particles of sea water.
! This parameterization assumes a simple second order kinetics (k[Particles][Fe]).
! Scavenging onto dust is also included as evidenced from the DUNE experiments.
-----
IF (ln_fechem ) THEN
zfeequi = ( zFe3(ji,jj,jk) + zFe2(ji,jj,jk) + zFeP(ji,jj,jk) ) * 1E-9
zfecoll = ( 0.3 * zFeL1(ji,jj,jk) + 0.5 * zFeL2(ji,jj,jk) ) * 1E-9
ELSE
IF (ln_fecoloid) THEN
zfeequi = zFe3(ji,jj,jk) * 1E-9
zhplus = max( rtrn, hi(ji,jj,jk) )
fe3sol = fesol(ji,jj,jk)
fe3sol = fesol(ji,jj,jk,1) * ( fesol(ji,jj,jk,2) * zhplus**2 &
&
+ fesol(ji,jj,jk,3) * zhplus + fesol(ji,jj,jk,4) )
&
&
+ fesol(ji,jj,jk,5) / zhplus )
zfecoll = max( ( 0.1 * zFeL1(ji,jj,jk) ) * 1E-9 , ( zFeL1(ji,jj,jk) * 1E-9 - fe3sol ) )
ELSE
zfeequi = zFe3(ji,jj,jk) * 1E-9
zfecoll = 0.5 * zFeL1(ji,jj,jk) * 1E-9
fe3sol = 0.
kfep = 0.
ENDIF
# if defined key_kriest
ztrc = ( trb(ji,jj,jk,jppoc) + trb(ji,jj,jk,jpcal) + trb(ji,jj,jk,jpgst) ) * 1.6
# else
ztrc = ( trb(ji,jj,jk,jppoc) + trb(ji,jj,jk,jpgoc) + trb(ji,jj,jk,jpcal) + trb(ji,jj,jk,jpgst) ) * 1.6
# endif

IF (ln_dust ) zdst = dust(ji,jj) / ( wdust / rday ) * tmask(ji,jj,jk) ! dust in kg/m2/s
zlam1b = 3.e-5 + xlamdust * zdst + xlam1 * ztrc
zscave = zfeequi * zlam1b * zstep

! Compute the different ratios for scavenging of iron
! to later allocate scavenged iron to the different organic pools
-----
zdenom1 = xlam1 * trb(ji,jj,jk,jppoc) / zlam1b
zdenom2 = xlam1 * trb(ji,jj,jk,jpgoc) / zlam1b

! Compute the proportion of filter feeders
zproport = (zgrazffep + zgrazffeg) / (rtrn + zgraztot)
! Compute fractionation of aggregates. It is assumed that
! diatoms based aggregates are more prone to fractionation
! since they are more porous (marine snow instead of fecal pellets)
zratio = trb(ji,jj,jk,jpgsi) / ( trb(ji,jj,jk,jpgoc) + rtrn )
zratio2 = zratio * zratio
zfrac = zproport * grazflux * zstep * wsbio4(ji,jj,jk)
&
* trb(ji,jj,jk,jpgoc) * trb(ji,jj,jk,jpmes)
&
* ( 0.2 + 3.8 * zratio2 / ( 1.*2 + zratio2 ) )
zfracfe = zfrac * trb(ji,jj,jk,jpbfe) / (trb(ji,jj,jk,jpgoc) + rtrn)

zgrazffep = zproport * zgrazffep
zgrazffeg = zproport * zgrazffeg
zgrazfff = zproport * zgrazfff
zgrazfff = zproport * zgrazfff
zgraztot = zgrazd + zgrazn + zgrazn + zgrazpoc + zgrazffep + zgrazffeg
zgraztot = zgrazd * quotan(ji,jj,jk) + zgrazn + zgrazn * quotan(ji,jj,jk)
&
+ zgrazpoc + zgrazffep + zgrazffeg
zgraztotf = zgrazf + zgraznf + zgrazn * ferat3 + zgrazpof + zgrazfff + zgrazfff

zgraztot = zgrazd + zgrazn + zgrazn + zgrazpoc + zgrazffep
! Compute the proportion of filter feeders
zproport = zgrazffep / ( zgraztot + rtrn )
zgrazffep = zproport * zgrazffep
zgrazfff = zproport * zgrazfff
zgraztot = zgrazd + zgrazn + zgrazn + zgrazpoc + zgrazffep
zgraztotn = zgrazd * quotan(ji,jj,jk) + zgrazn + zgrazn * quotan(ji,jj,jk) + zgrazpoc + zgrazffep
zgraztotf = zgrazf + zgraznf + zgrazn * ferat3 + zgrazpof + zgrazfff

! Total grazing ( grazing by microzoa is already computed in p4zmicro )
IF (lk_tomput ) zgrazing(ji,jj,jk) = zgraztot

! Mesozooplankton efficiency
-----
zgrasrat = ( zgraztotf + rtrn ) / ( zgraztot + rtrn )
zgrasratn = ( zgraztotn + rtrn ) / ( zgraztot + rtrn )
zepshtert = MIN( 1., zgrasratn, zgrasrat / ferat3 )
zepshterv = zepshtert * MIN( epshter2, ( 1. - unass2 ) * zgrasrat / ferat3, ( 1. - unass2 ) * zgrasratn )
zgrarem2 = zgraztot * ( 1. - zepshterv - unass2 ) &
&
+ ( 1. - epshter2 - unass2 ) / ( 1. - epshter2 ) * ztortz2
zgrarer2 = zgraztot * MAX( 0., ( 1. - unass2 ) * zgrasrat - ferat3 * zepshterv ) &
&
+ ferat3 * ( ( 1. - epshter2 - unass2 ) / ( 1. - epshter2 ) * ztortz2 )
zgrapoc2 = zgraztot * unass2

! Update the arrays TRA which contain the biological sources and sinks
zgransig = zgrarem2 * sigma2
tra(ji,jj,jk,jppo4) = tra(ji,jj,jk,jppo4) + zgransig
tra(ji,jj,jk,jpnh4) = tra(ji,jj,jk,jpnh4) + zgransig
tra(ji,jj,jk,jpdoc) = tra(ji,jj,jk,jpdoc) + zgrarem2 - zgransig
# if defined key_ligand
tra(ji,jj,jk,jplgw) = tra(ji,jj,jk,jplgw) + (zgrarem2 - zgransig) * ldocz
z2l1gprod(ji,jj,jk) = (zgrarem2 - zgransig) * ldocz

tra(ji,jj,jk,jpox) = tra(ji,jj,jk,jpox) - zout * zgransig
tra(ji,jj,jk,jpfer) = tra(ji,jj,jk,jpfer) + zgrarer2
zfezoo2(ji,jj,jk) = zgrarer2
tra(ji,jj,jk,jpdic) = tra(ji,jj,jk,jpdic) + zgransig
tra(ji,jj,jk,jptal) = tra(ji,jj,jk,jptal) + rno3 * zgransig

zmortz2 = ztortz2 + zrespz2
zmortzoc = unass2 / ( 1. - epshter2 ) * ztortz2 + zrespz2
tra(ji,jj,jk,jpmes) = tra(ji,jj,jk,jpmes) - zmortz2 + zepshterv * zgraztot
tra(ji,jj,jk,jpdia) = tra(ji,jj,jk,jpdia) - zgrazd

```

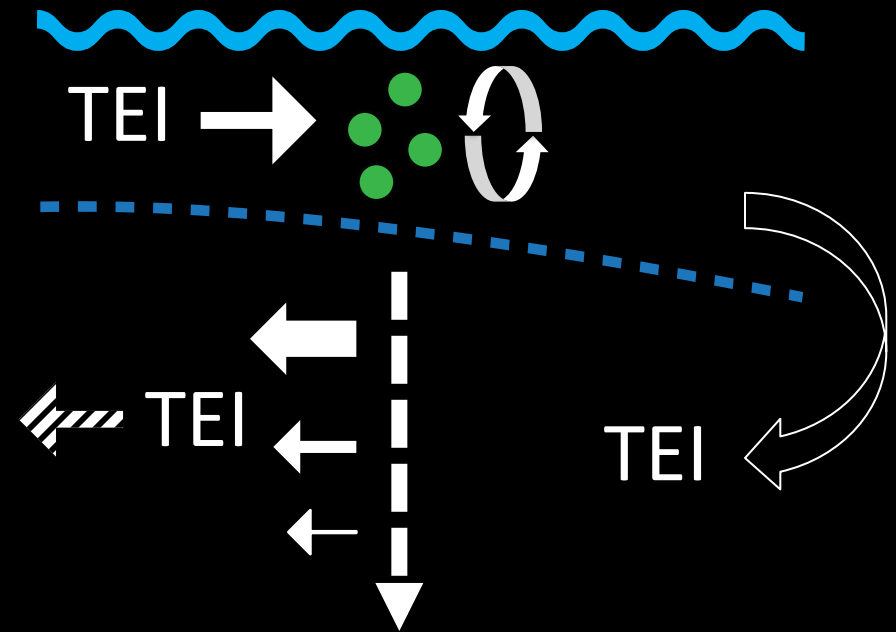
What are the main processes at play?

Surface uptake and cycling

Sinking and regeneration

Scavenging

Subduction and transport



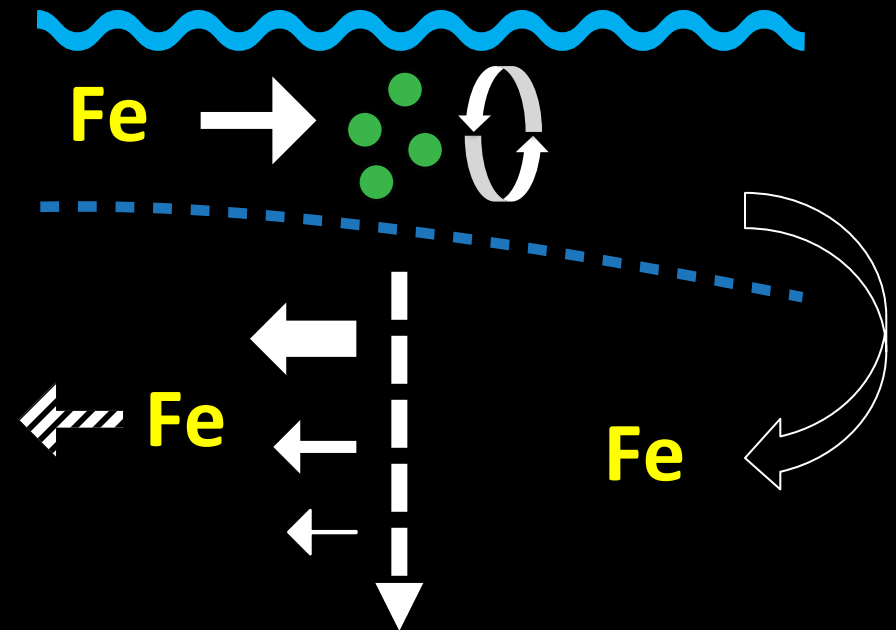
What are the main processes at play?

Surface uptake and cycling

~~Sinking and regeneration~~

Scavenging

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Necessary focus here on Fe ...

Evolution of internal cycling in GCMs

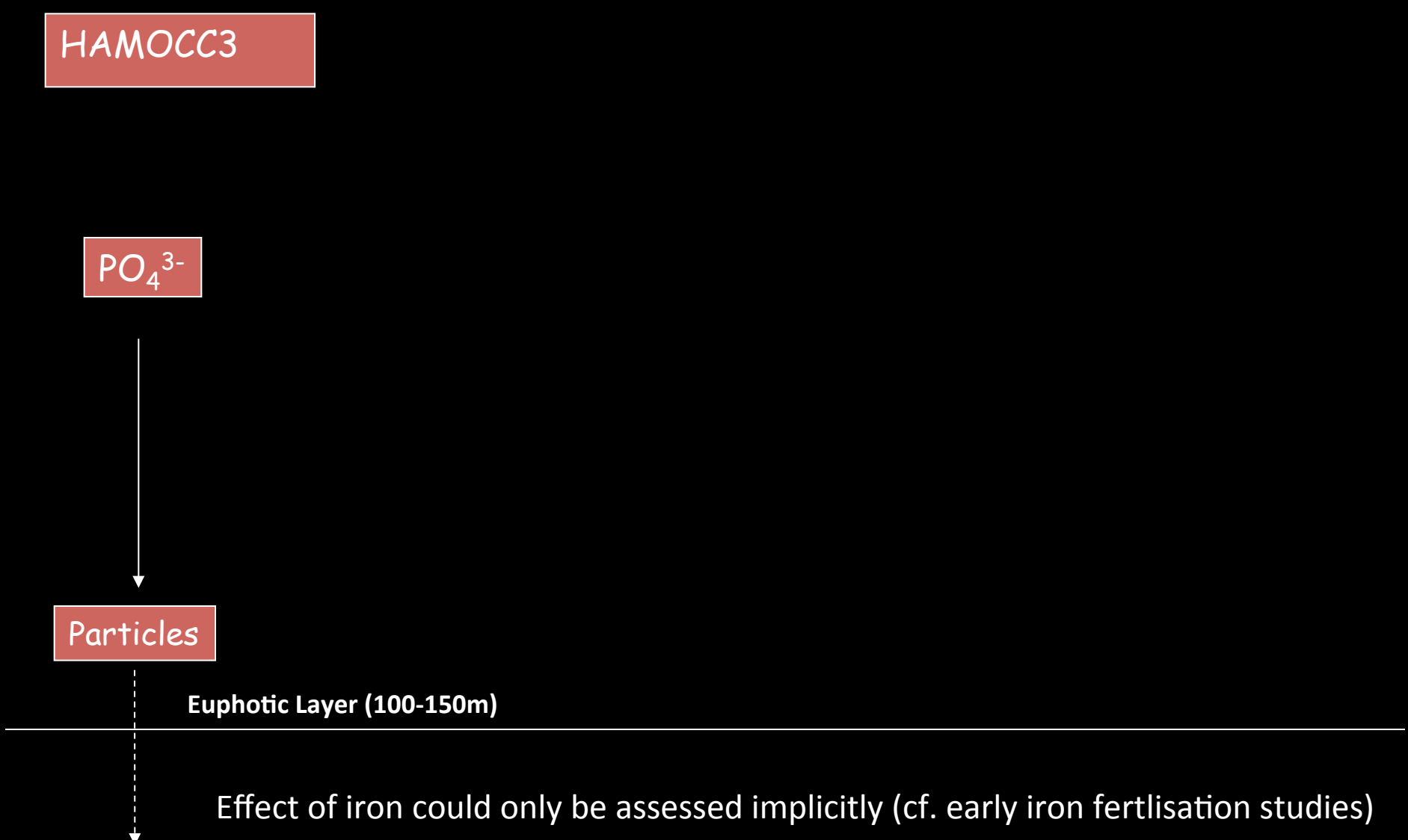
HAMOCC3

PO_4^{3-}

Particles

Euphotic Layer (100-150m)

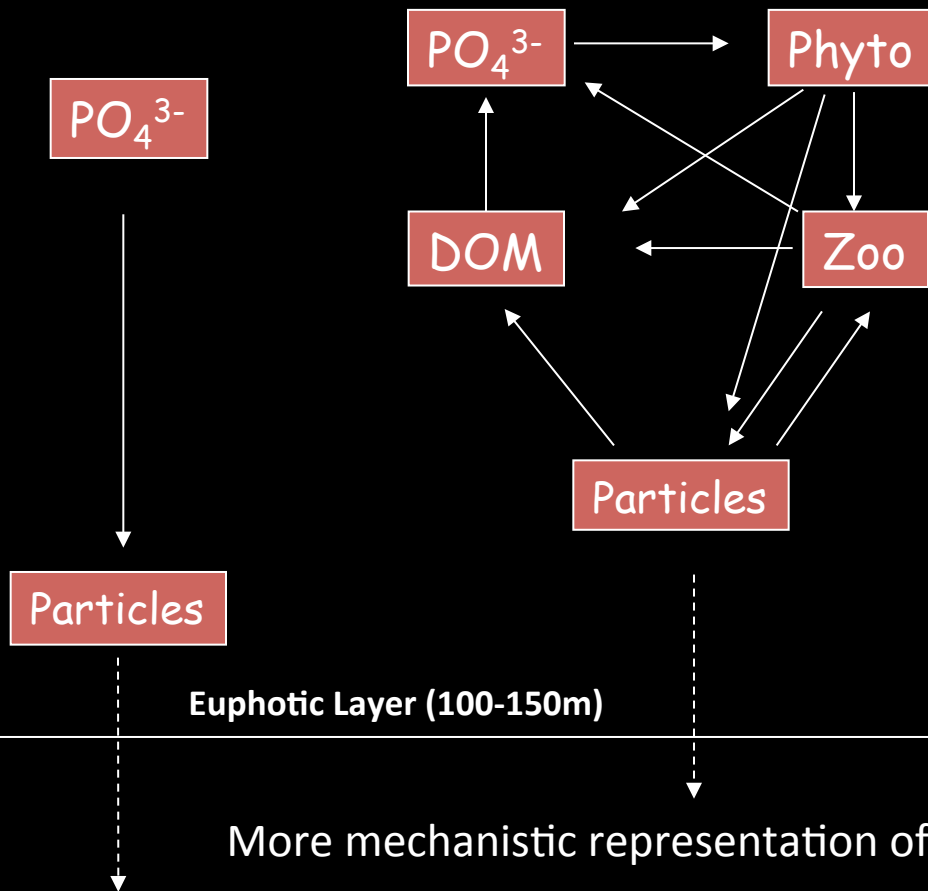
Effect of iron could only be assessed implicitly (cf. early iron fertilisation studies)



Evolution of internal cycling in GCMs

HAMOCC3

NPZD

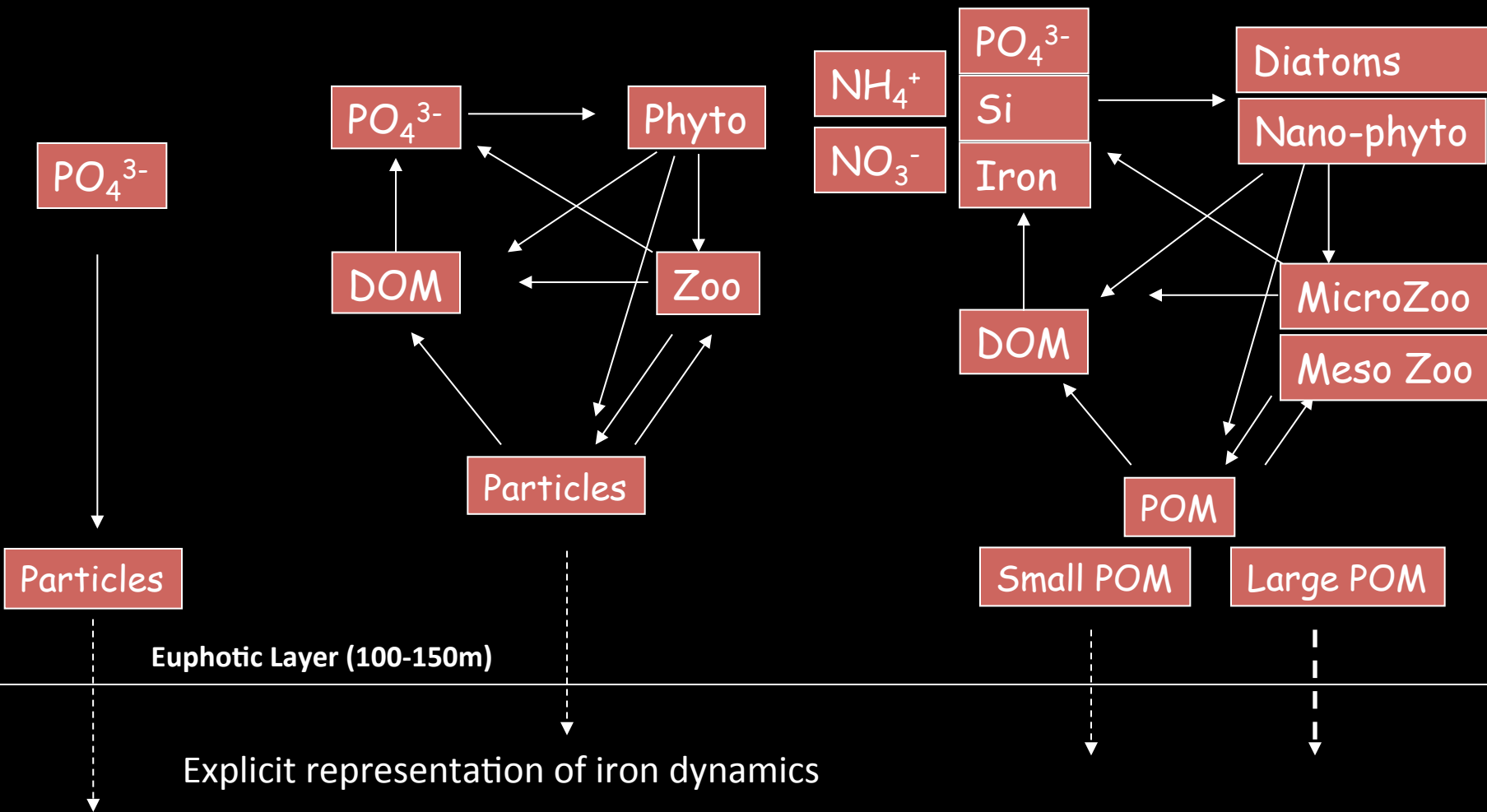


Evolution of internal cycling in GCMs

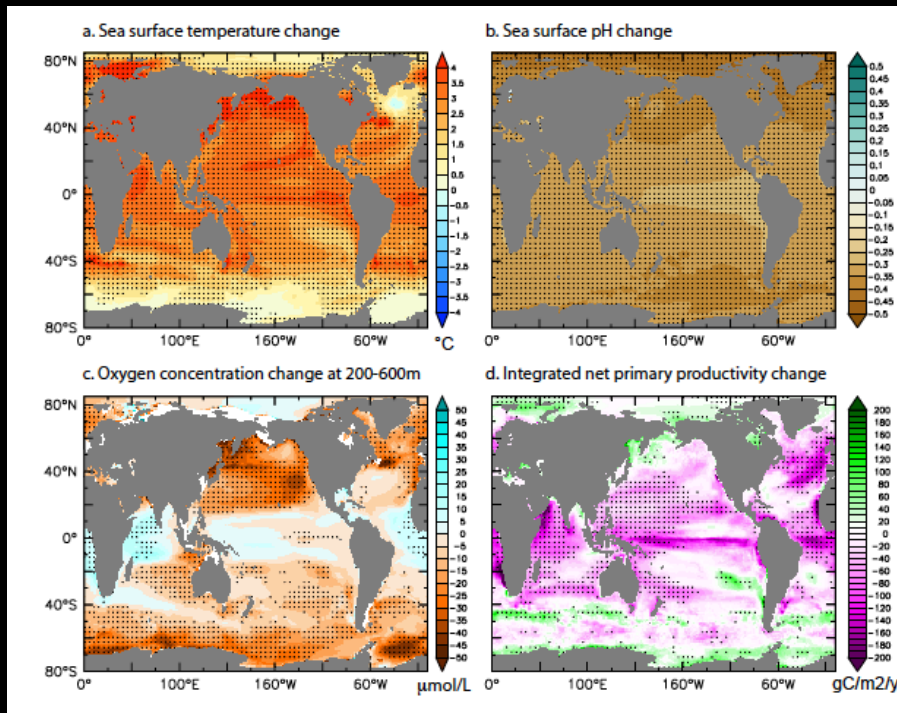
HAMOCC3

NPZD

OBGCM



We rely on models ...

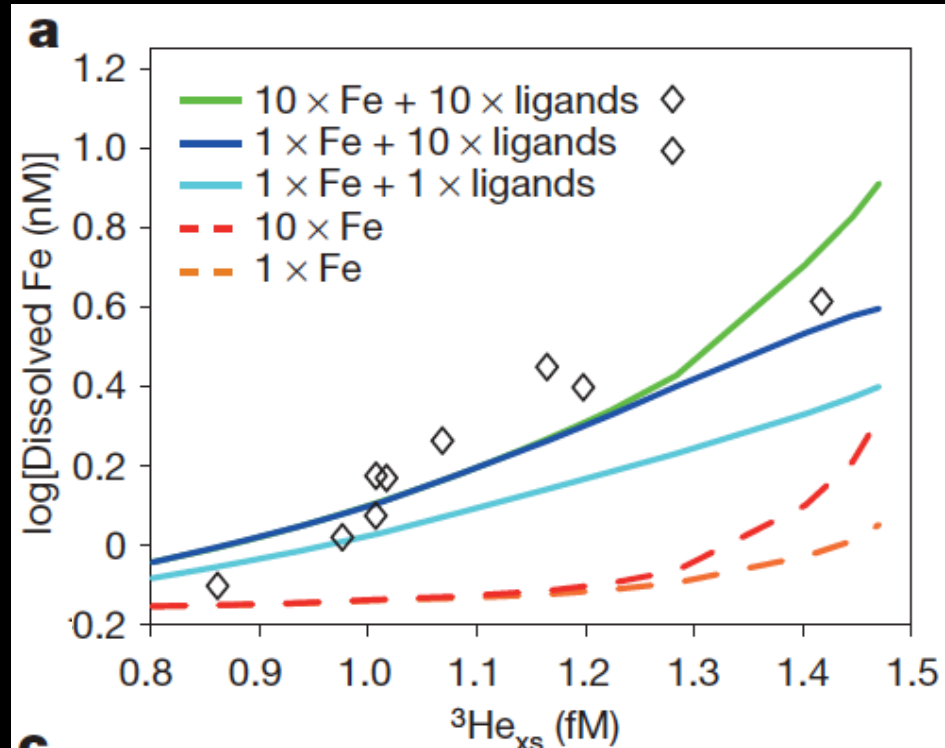


Biogeosciences, 10, 6225–6245, 2013
 www.biogeosciences.net/10/6225/2013/
 doi:10.5194/bg-10-6225-2013
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Multiple stressors of ocean ecosystems in the 21st century: projections with CMIP5 models

L. Bopp¹, L. Resplandy¹, J. C. Orr¹, S. C. Doney⁵, J. P. Dunne⁶, M. Gehlen¹, P. Halloran³, C. Heinze^{6,9,10}, T. Ilyina⁴, R. Séférian^{1,5}, J. Tjiputra^{6,9,10}, and M. Vichi⁷



LETTER

doi:10.1038/nature14577

Basin-scale transport of hydrothermal dissolved metals across the South Pacific Ocean

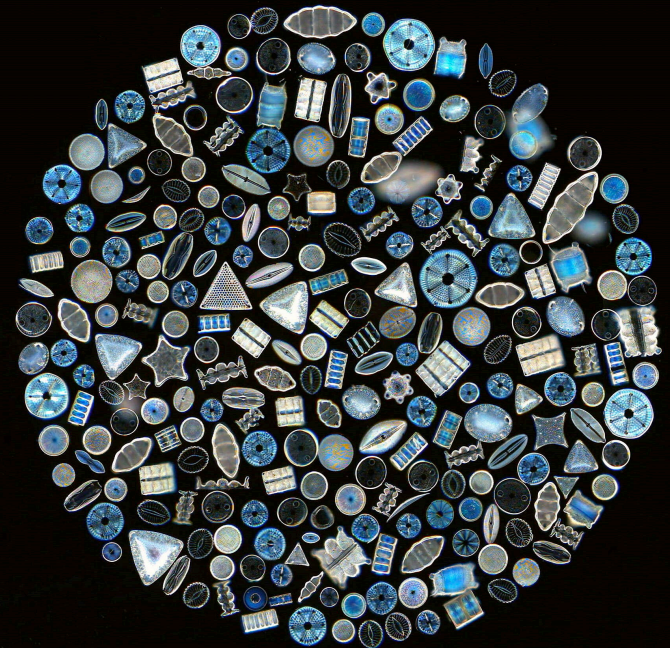
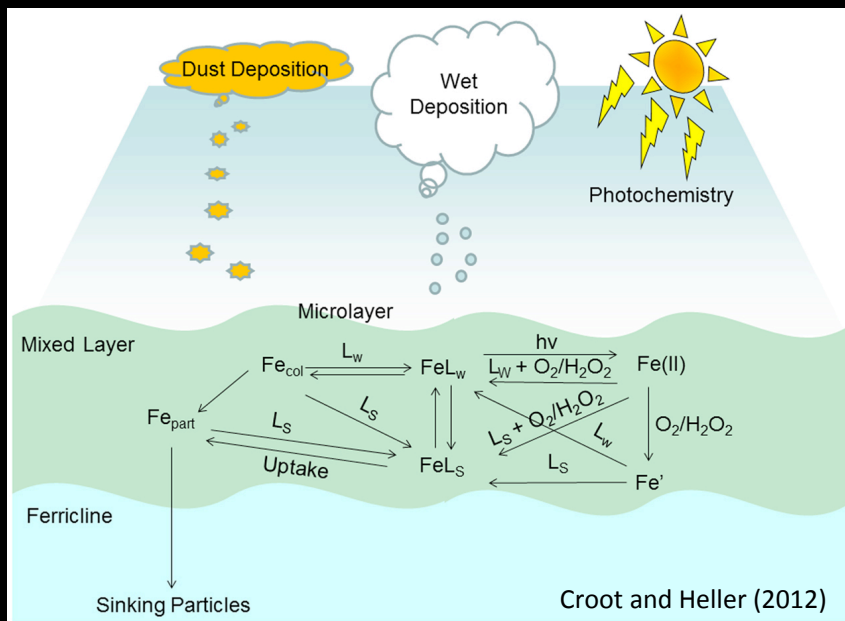
Joseph A. Resing¹, Peter N. Sedwick², Christopher R. German³, William J. Jenkins³, James W. Moffett⁴, Bettina M. Sohst⁵ & Alessandro Tagliabue⁶

Projecting the future

Testing hypotheses

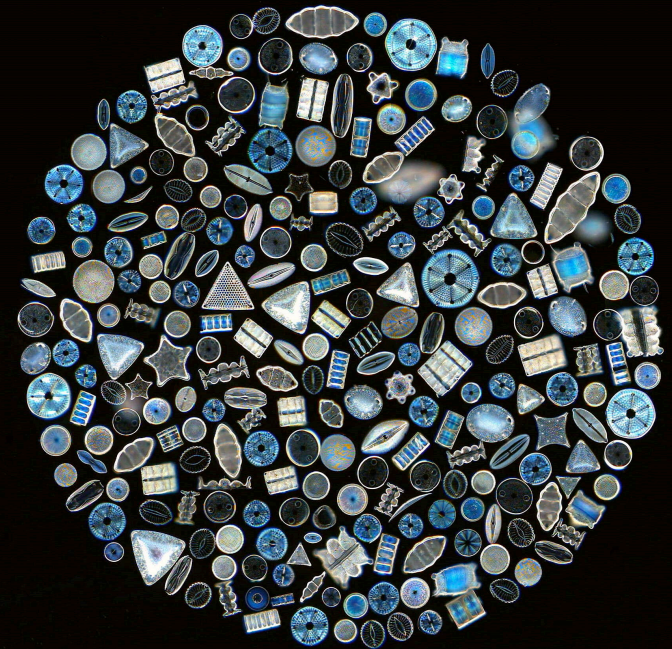
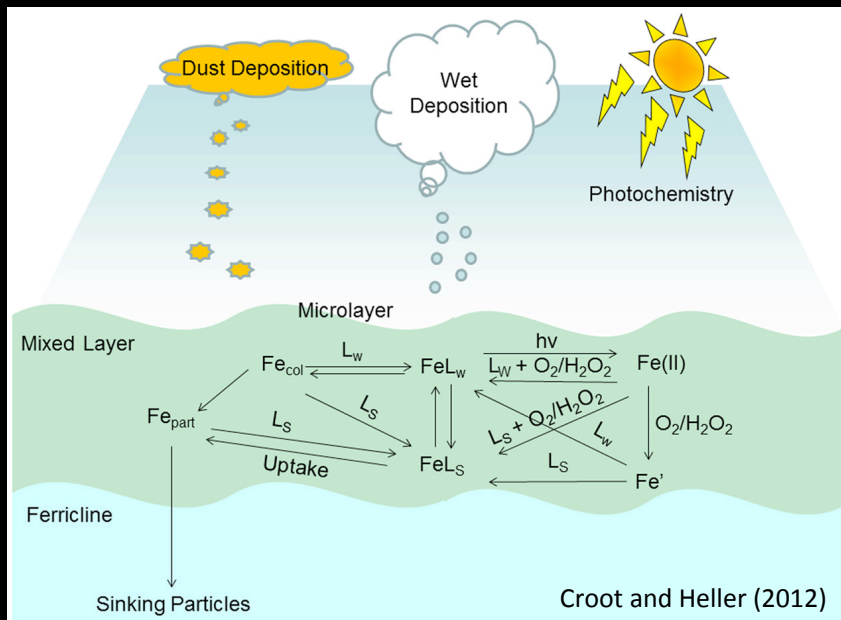
Models are underpinned by *trade offs*

- We *want* lots of processes to reflect the complexity of nature ...



Models are underpinned by *trade offs*

- We *want* lots of processes to reflect the complexity of nature ...



- But we face two trade offs:
 - How do we parameterise the model?
 - It may become too slow to be useful

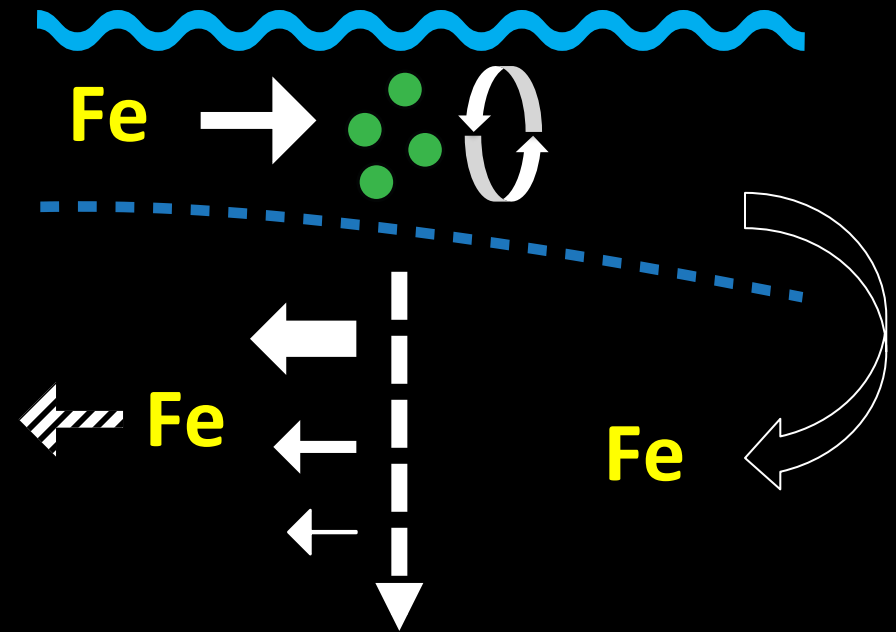
What are the main processes at play?

Surface uptake and cycling

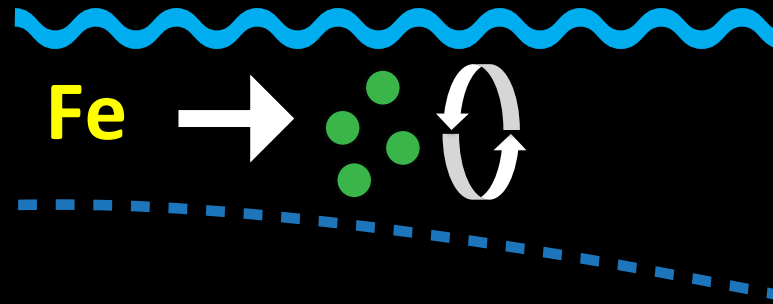
Sinking and regeneration

Scavenging

Subduction and transport

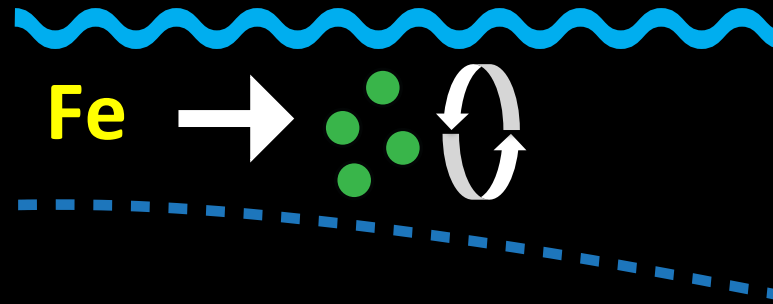


Surface uptake and cycling



Three main concepts to deal with:

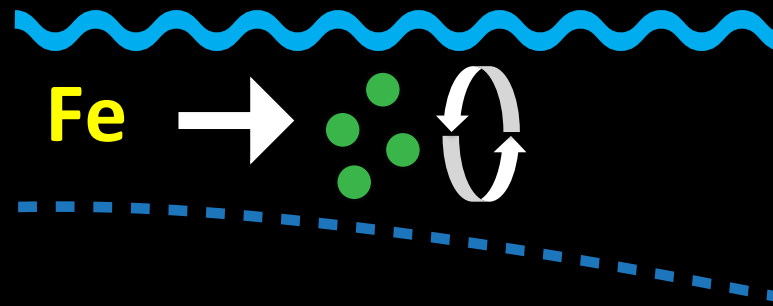
Surface uptake and cycling



Three main concepts to deal with:

1. Uptake
2. Impact on Growth rate
3. Recycling

Surface uptake and cycling

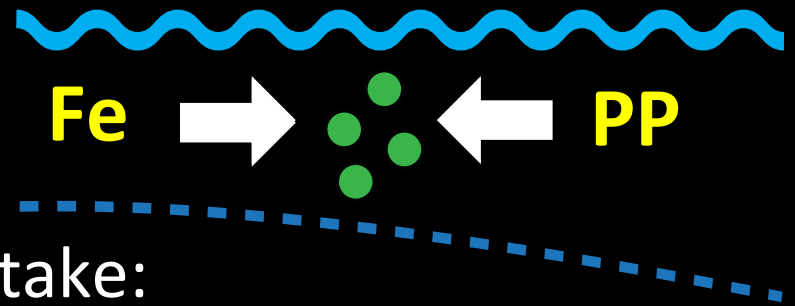


Three main concepts to deal with:

- 1. Uptake**
2. Impact on Growth rate
3. Recycling

Uptake by microbes

Two main philosophies

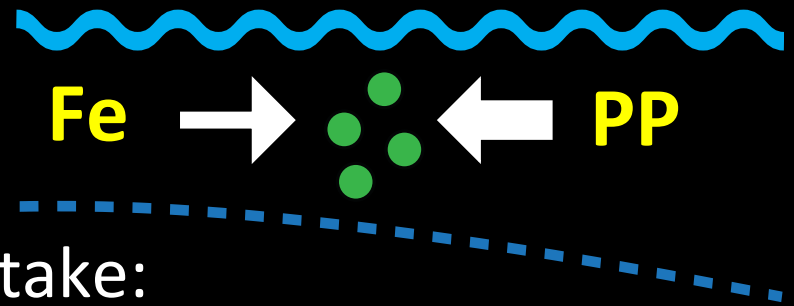


1. **Coupled** (aka “Redfield”) uptake:

Fe uptake derived from rate of primary production via imposed stoichiometry:

Uptake by microbes

Two main philosophies



1. **Coupled** (aka “Redfield”) uptake:

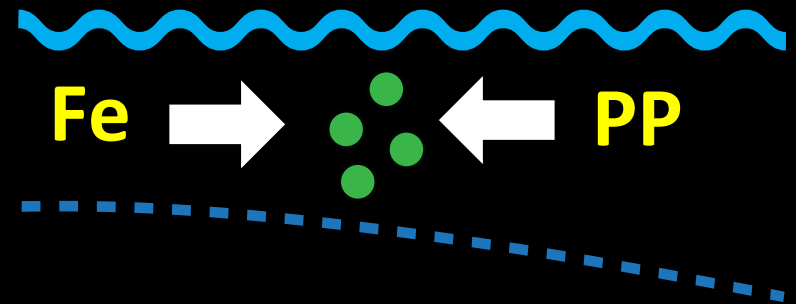
Fe uptake derived from rate of primary production via imposed stoichiometry

2. **Decoupled** (aka “Michaelis Menten”) uptake:

Fe uptake derived independent of primary production via kinetic equations

“Coupled” uptake by microbes

Main assumptions:



- Uptake of Fe is ‘slave’ to primary production
- Often driven by fixed Fe/C stoichiometry
- > Fe uptake varies proportionally with PP and growth
- > Fixed Fe/C stoichiometry attractive from an efficiency standpoint (cf. trade offs)

“Decoupled” uptake by microbes

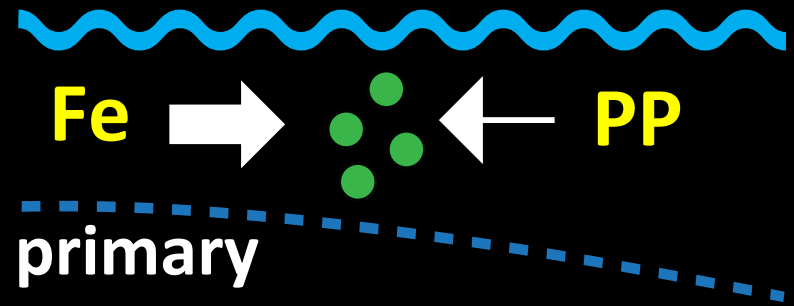
Main assumptions:

- Uptake of Fe is independent of primary production

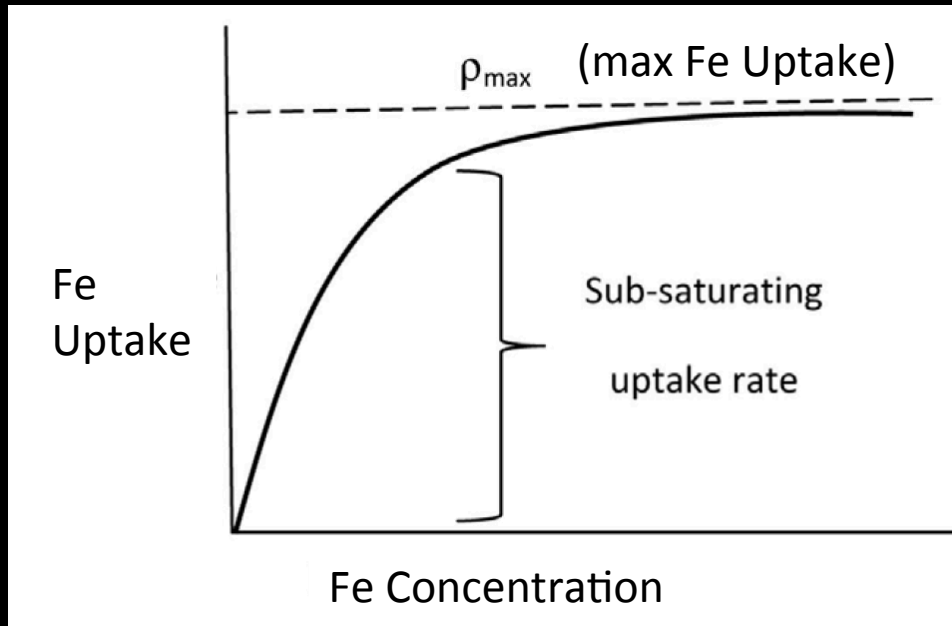
- Driven by Michaelis Menten Kinetics

- > Allows Fe uptake to continue when PP is limited (e.g. by light)

- > More computationally expensive



“Decoupled” uptake by microbes



Adapted from Shaked and Lis (2012)

$$\rho = \rho_{MAX} \frac{Fe}{Fe + K_s} + \text{Luxury} + \text{Surge}$$

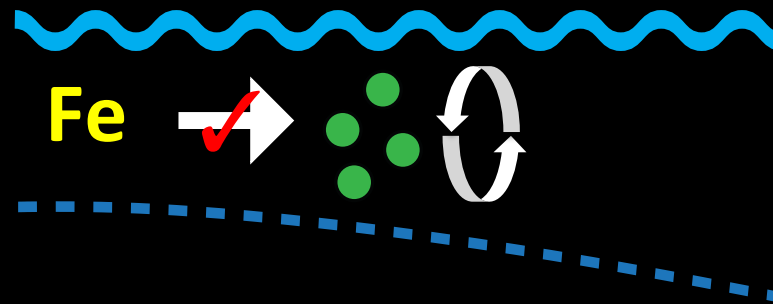
sensu Morel (1987, J Phyc)

“Decoupled” uptake by microbes

Key parameters:

- Affinity for (specific?) forms of Fe
- Imposed maximum cellular quota
- Relative increase in Fe uptake at low Fe

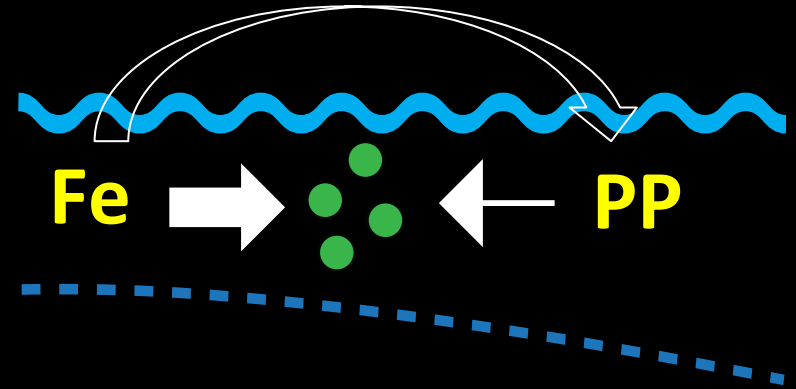
Surface uptake and cycling



Three main concepts to deal with:

1. Uptake
- 2. Impact on Growth rate**
3. Recycling

Growth Limitation



Two main philosophies

1. Monod Limitation:

Simplest ; external concentration drives growth rate

2. Quota (or Droop) Limitation:

Complex ; internal quota drives growth rate

Growth Limitation

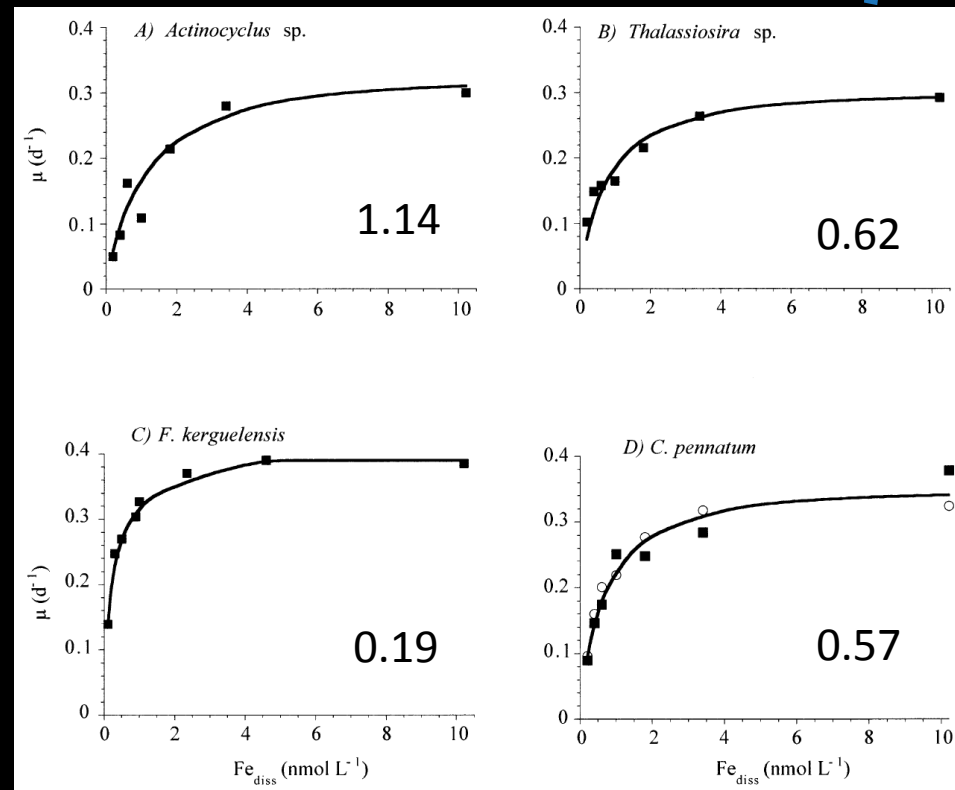
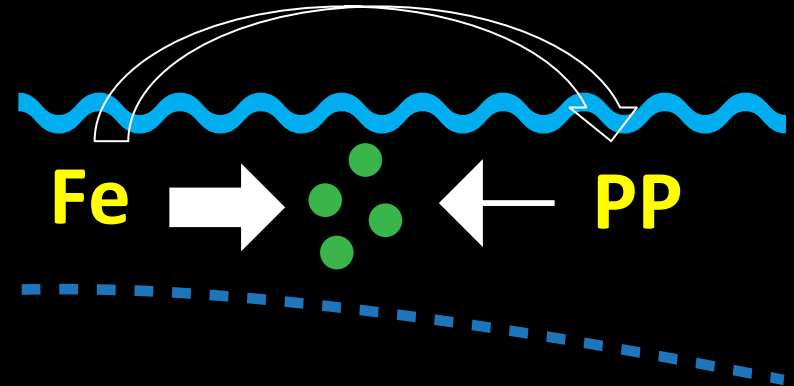
Two main philosophies

1. Monod Limitation:

$$\mu = \mu_{MAX} \frac{Fe}{Fe + K_{\mu}}$$

Key parameter is K_{μ} ; can be derived from experimental studies

Timmermans et al (2004, L&O)



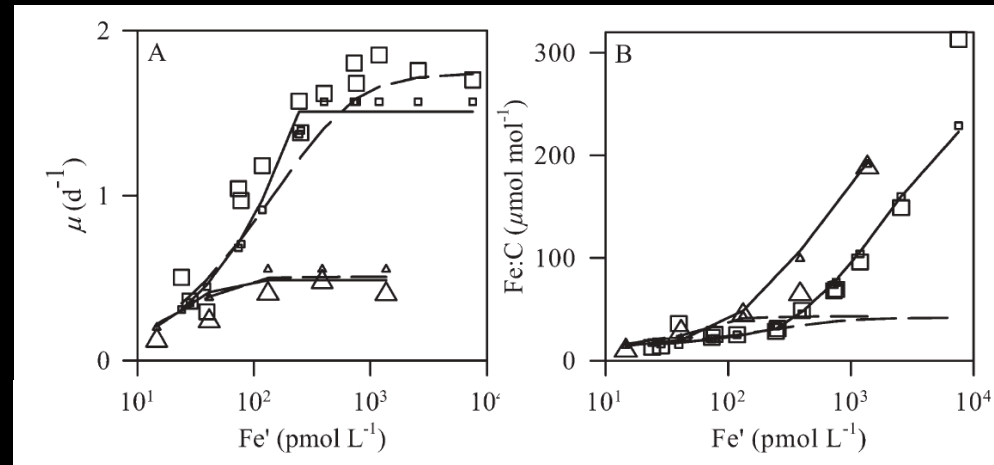
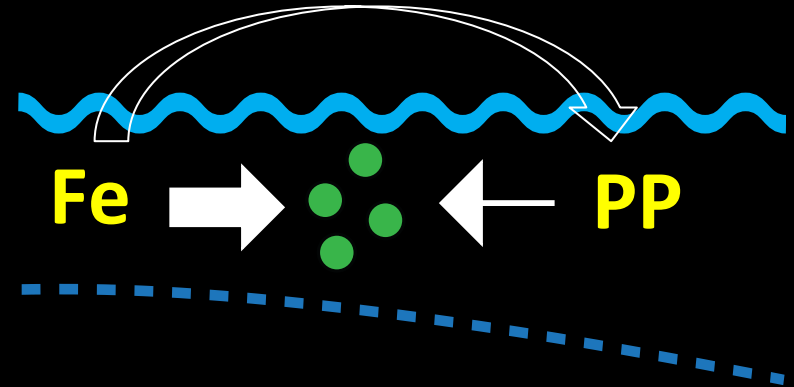
Growth Limitation

Two main philosophies

2. Quota Limitation

$$\mu = \mu_{MAX} \frac{Q - Q_{REQ}}{Q_{OPT}}$$

Key parameters are Q_{REQ} and Q_{OPT} ; can be derived from first principles / physiology / optimisation



Buitenhuis and Geider (2010, L&O)

Growth Limitation

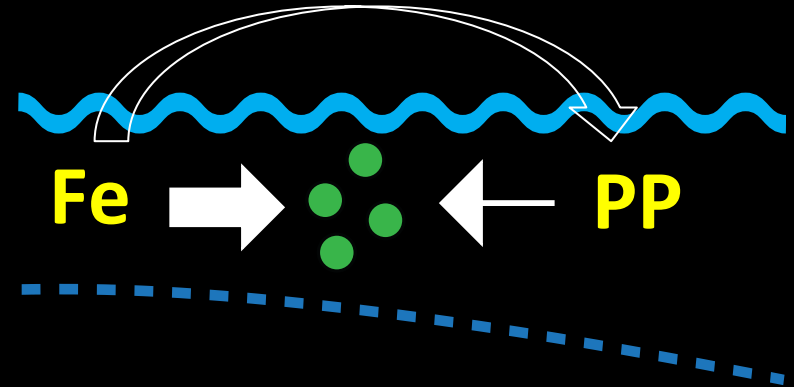
Two main philosophies

2. Quota Limitation

$$\mu = \mu_{MAX} \frac{Q - Q_{REQ}}{Q_{OPT}}$$

Key parameters are Q_{REQ} and Q_{OPT} ; can be derived from **first principles** / physiology / optimisation

after Raven (1990, J Phyc) Flynn and Hipkin (1999, JPR)



Contributions from:

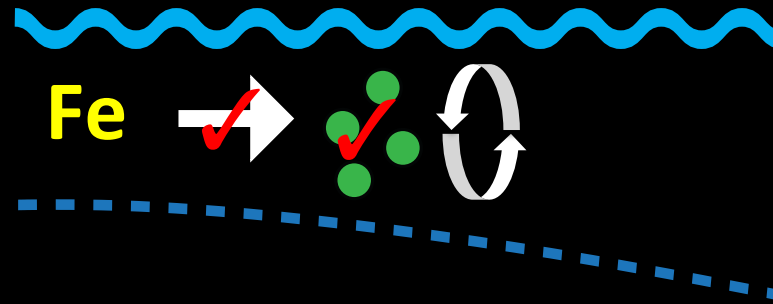
Fe/Chl per PSU

Nitrate reduction

Respiration

Provides mechanistic links to physiology!

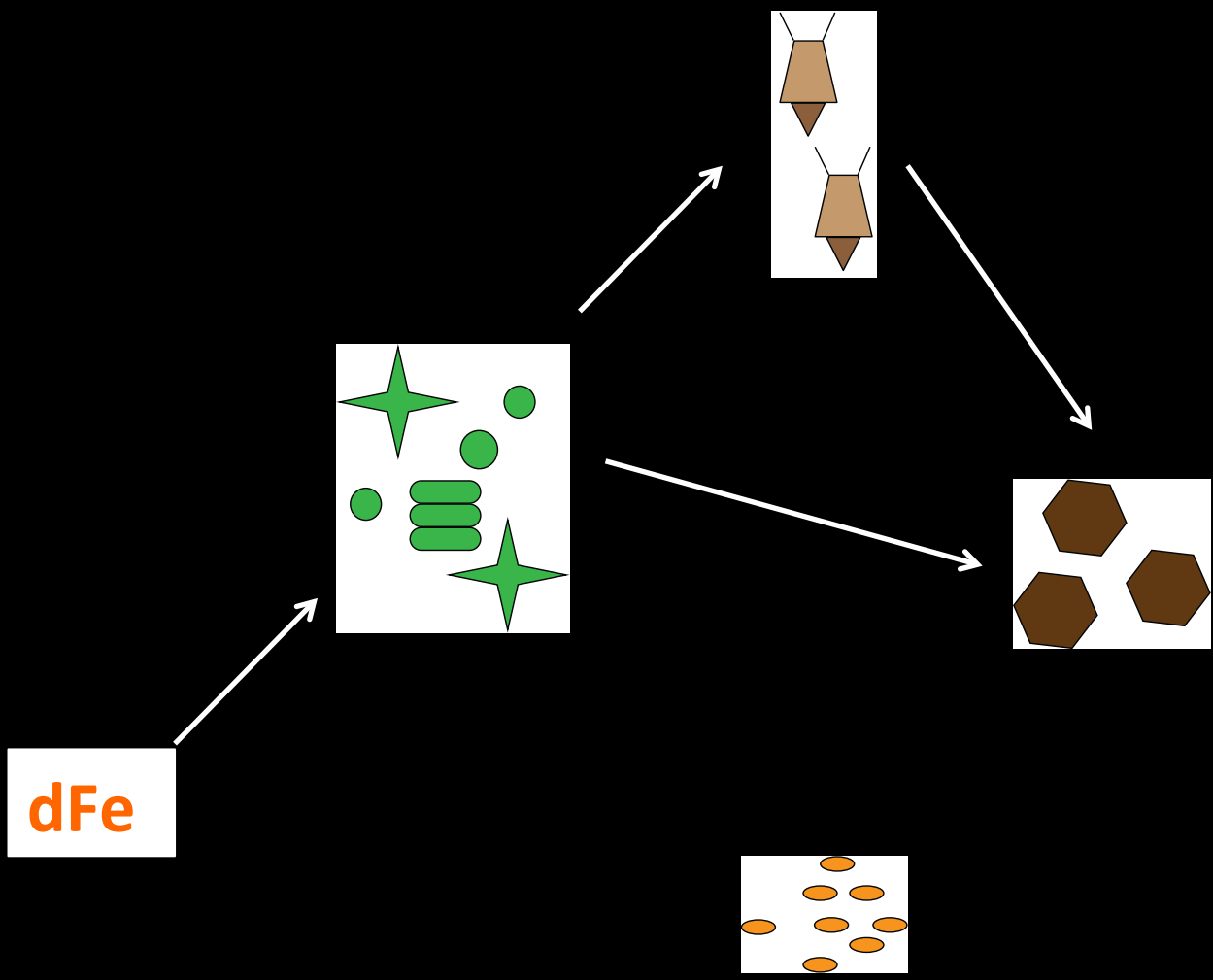
Surface uptake and cycling



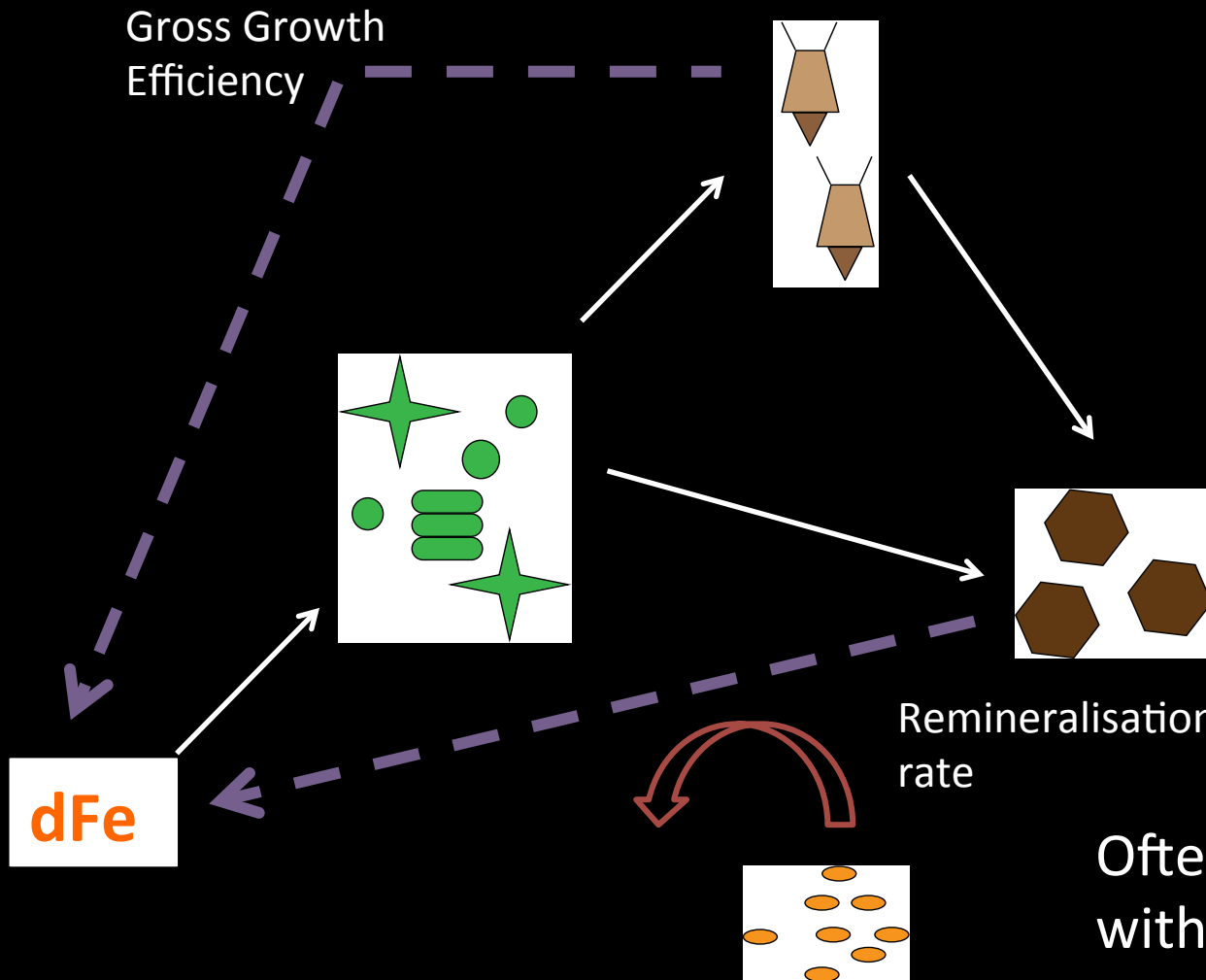
Three main concepts to deal with:

1. Uptake
2. Impact on Growth rate
- 3. Recycling**

Recycling

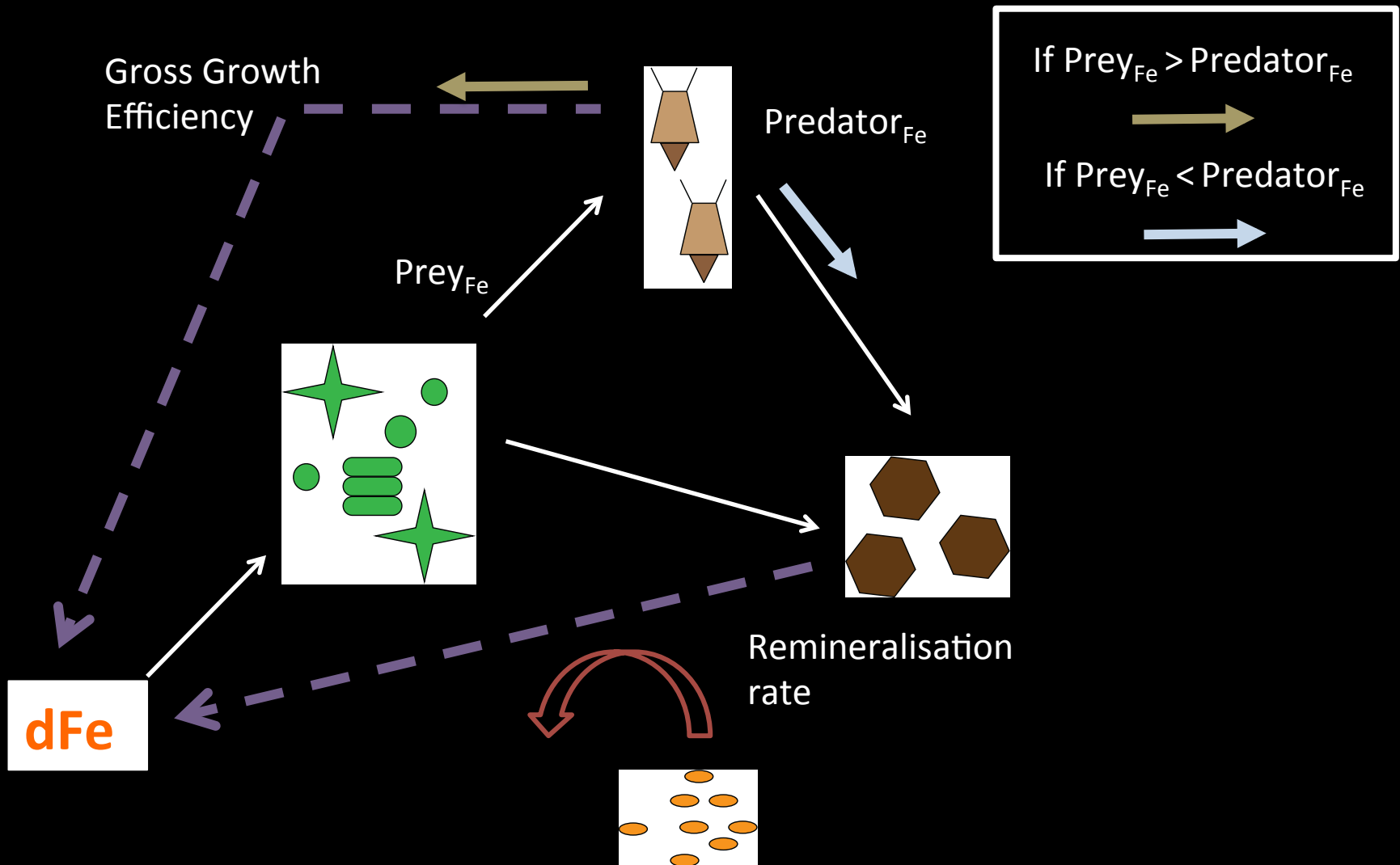


Recycling

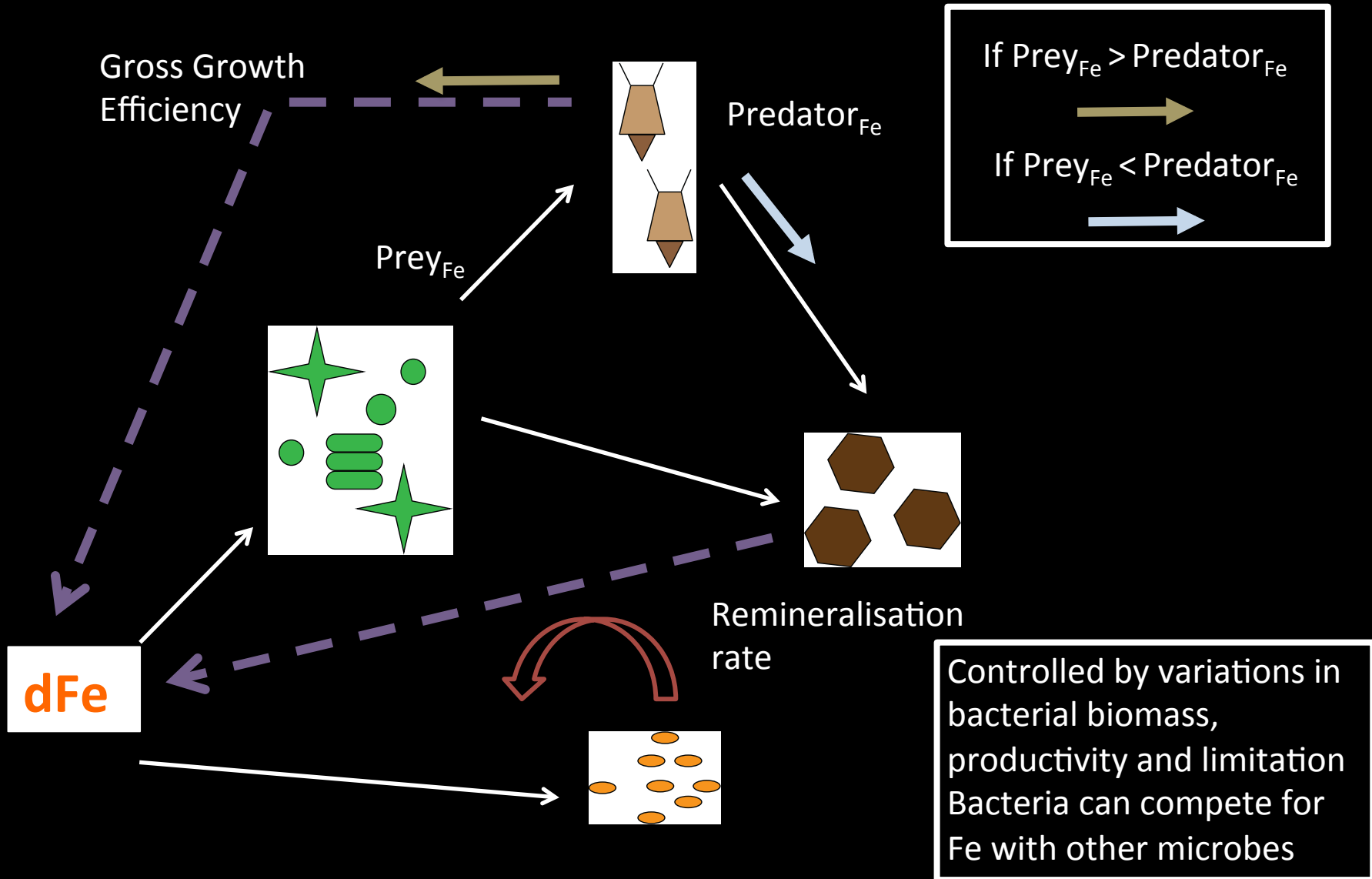


Often simply modelled with fixed rates

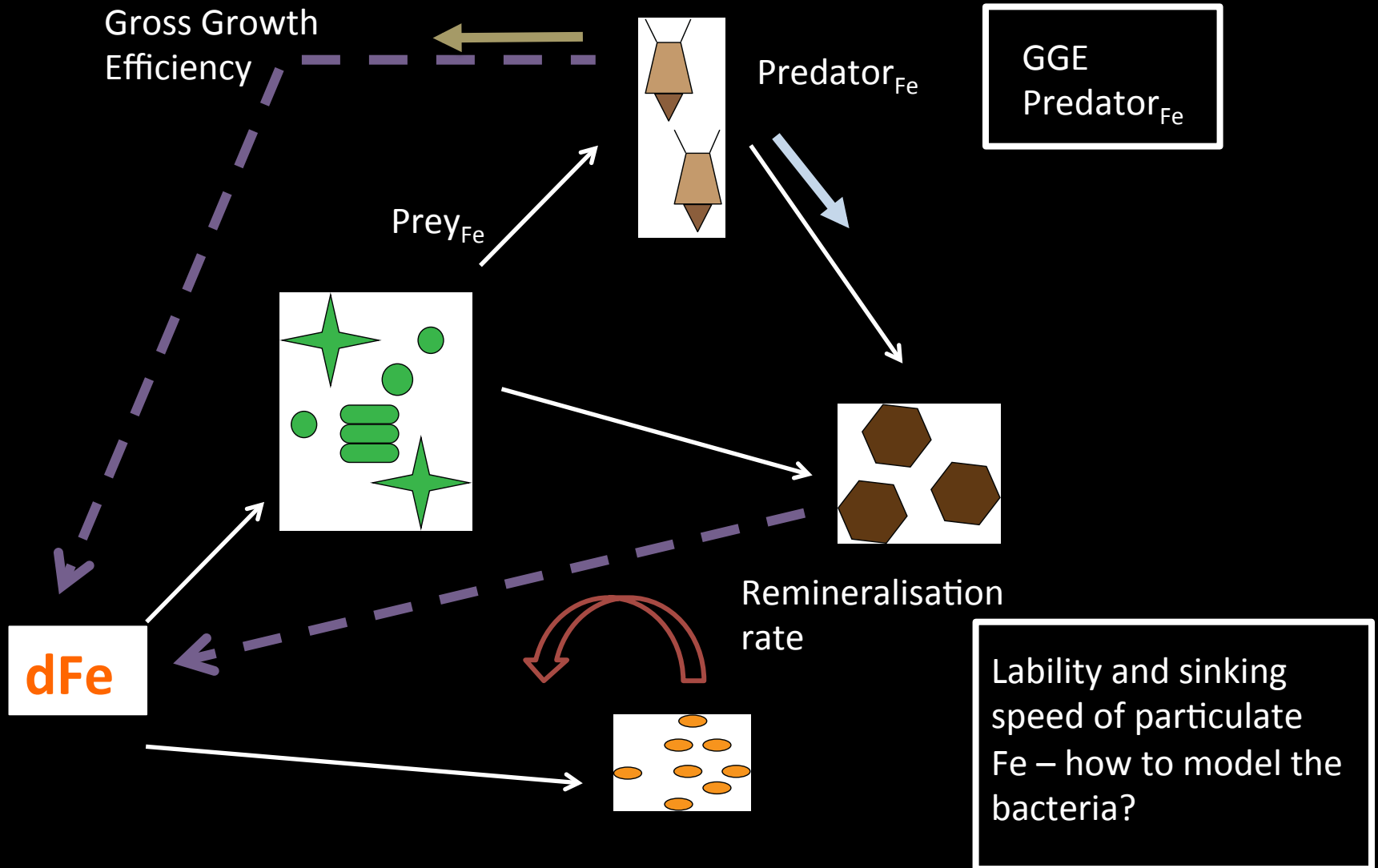
Recycling



Recycling



Recycling – Key Parameters



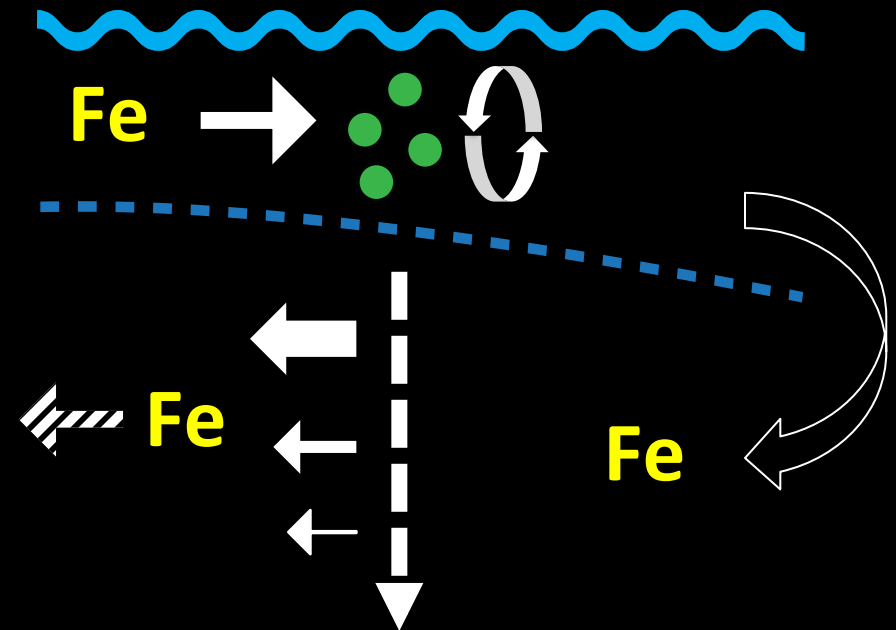
What are the main processes at play?

✓ Surface uptake and cycling

~~Sinking and~~ ✓ regeneration

Scavenging

Subduction and transport



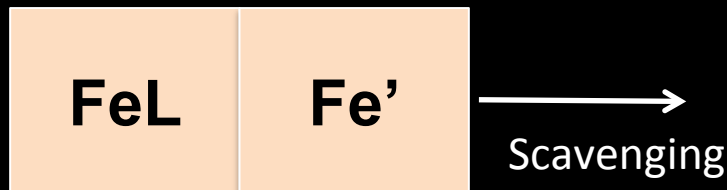
Scavenging

- Like other TElS, iron is lost from the dissolved pool by 'abiotic processes'
- Fundamental role for organic ligands
- Two main issues:
 - Speciation of Fe
 - 'Loss' of Fe from the dissolved pool

Speciation

- Key is to model how much Fe is complexed and how much is 'free'

Threshold Model

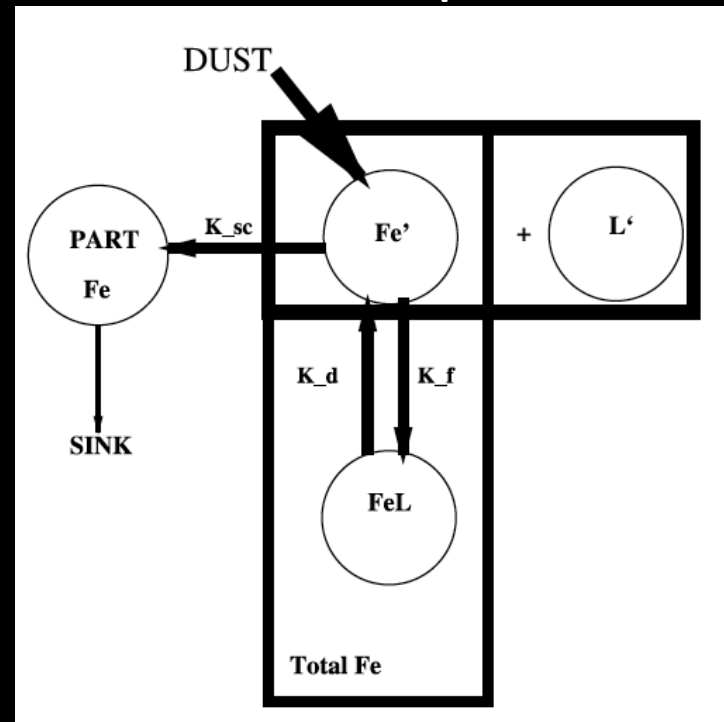


Ligand Concentration

Johnson et al. (1997, Mar Chem)

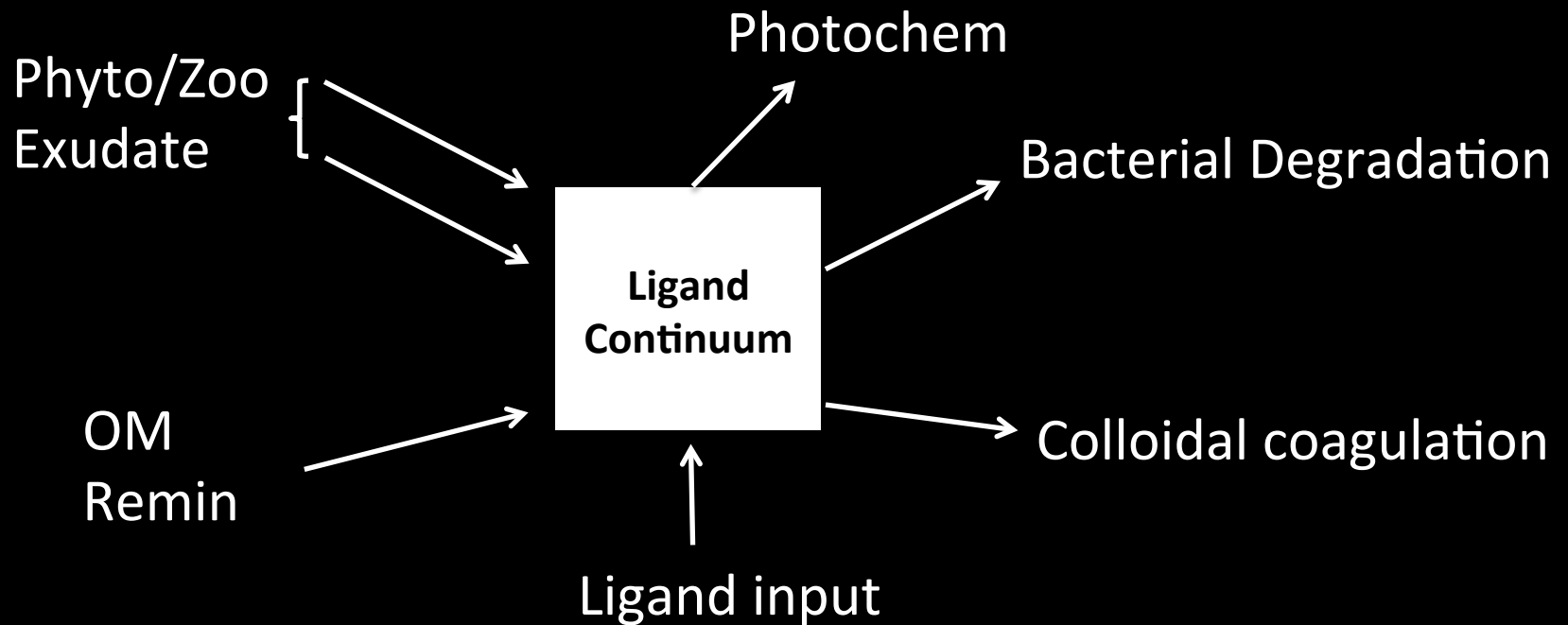
Requires information on
[L] and K_{FeL}

Complexation Model



Parekh et al. (2004, GBC)

Dynamic Ligands

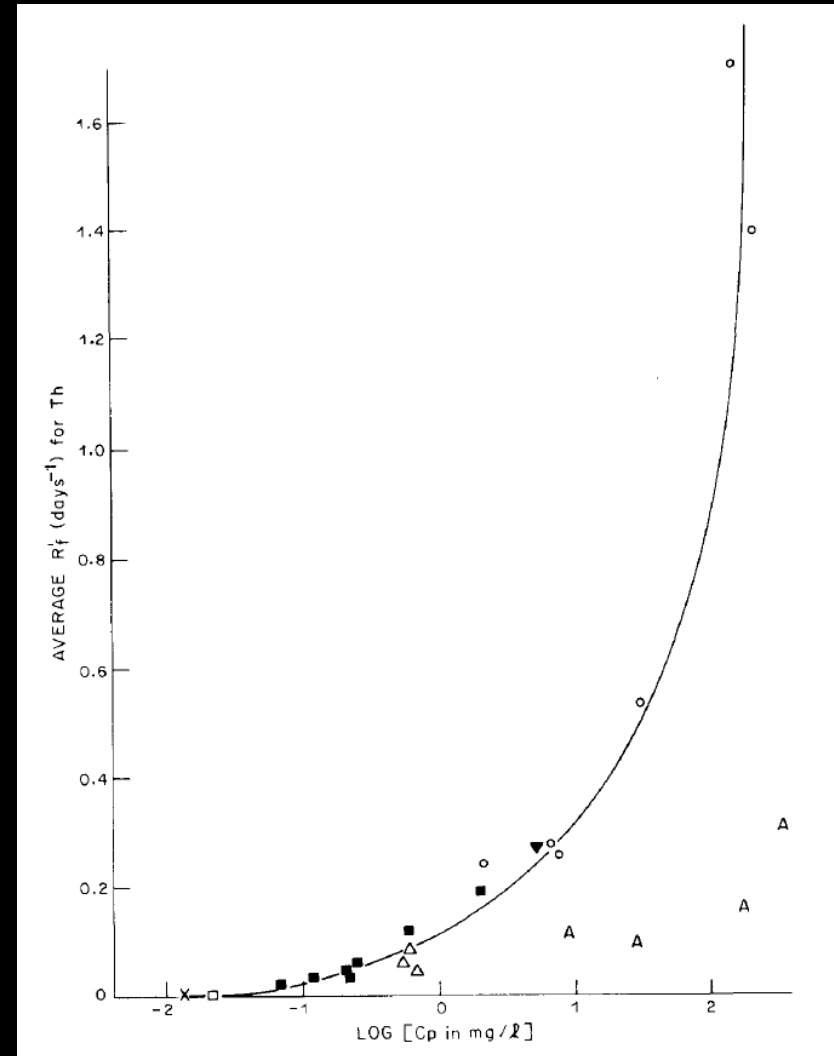


More attractive from a mechanistic standpoint, but requires more parameter choices (trade off)

Benchmarking these against other tracers (e.g. DOC production, O₂ consumption) is more helpful for models

Loss from the dissolved pool

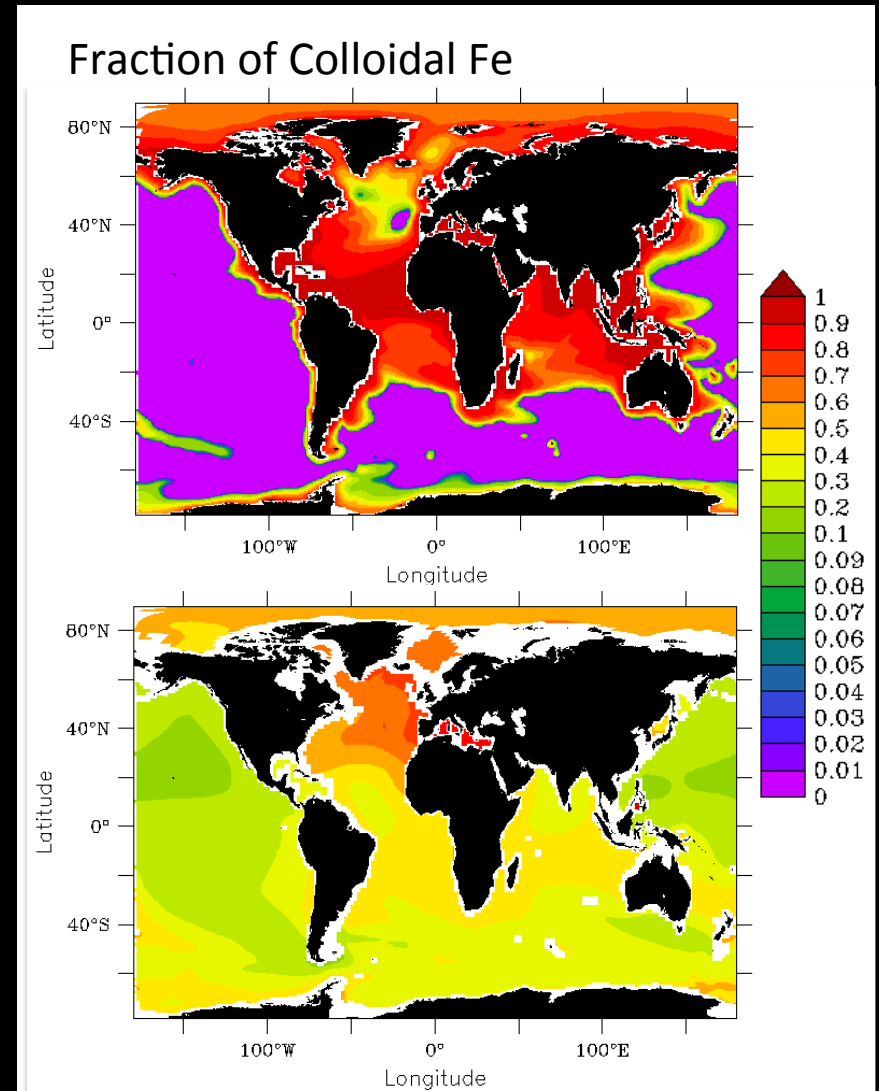
- Although a function of particle load, many models still use fixed rates
- At its most complex, scavenging is modelled as a function of particle load
- Crucial is the relative role played by each particle type (small and large POC, biogenic Si, calcite, lithogenics)
- Insights from other TEIs (e.g Pa/Th) potentially transformative



Honeyman et al (1988, DSR)

Colloidal Fe

- Important new datasets emerging on colloidal TEIs
- Usually ignored in models
- Where represented, relies on equilibrium speciation from laboratory studies
- Need to understand and represent unique roles for soluble and colloidal TEIs to further their representation in models



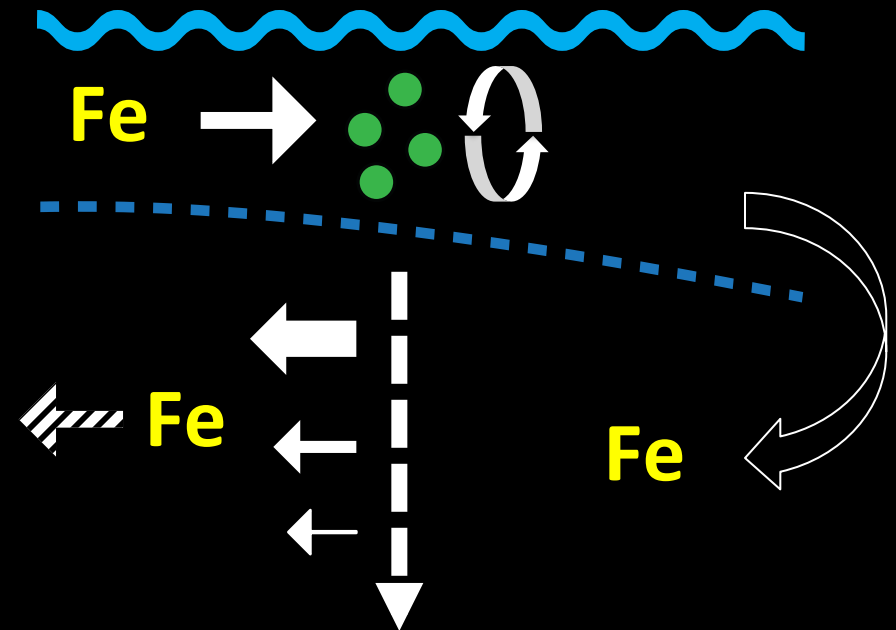
What are the main processes at play?

✓ Surface uptake and cycling

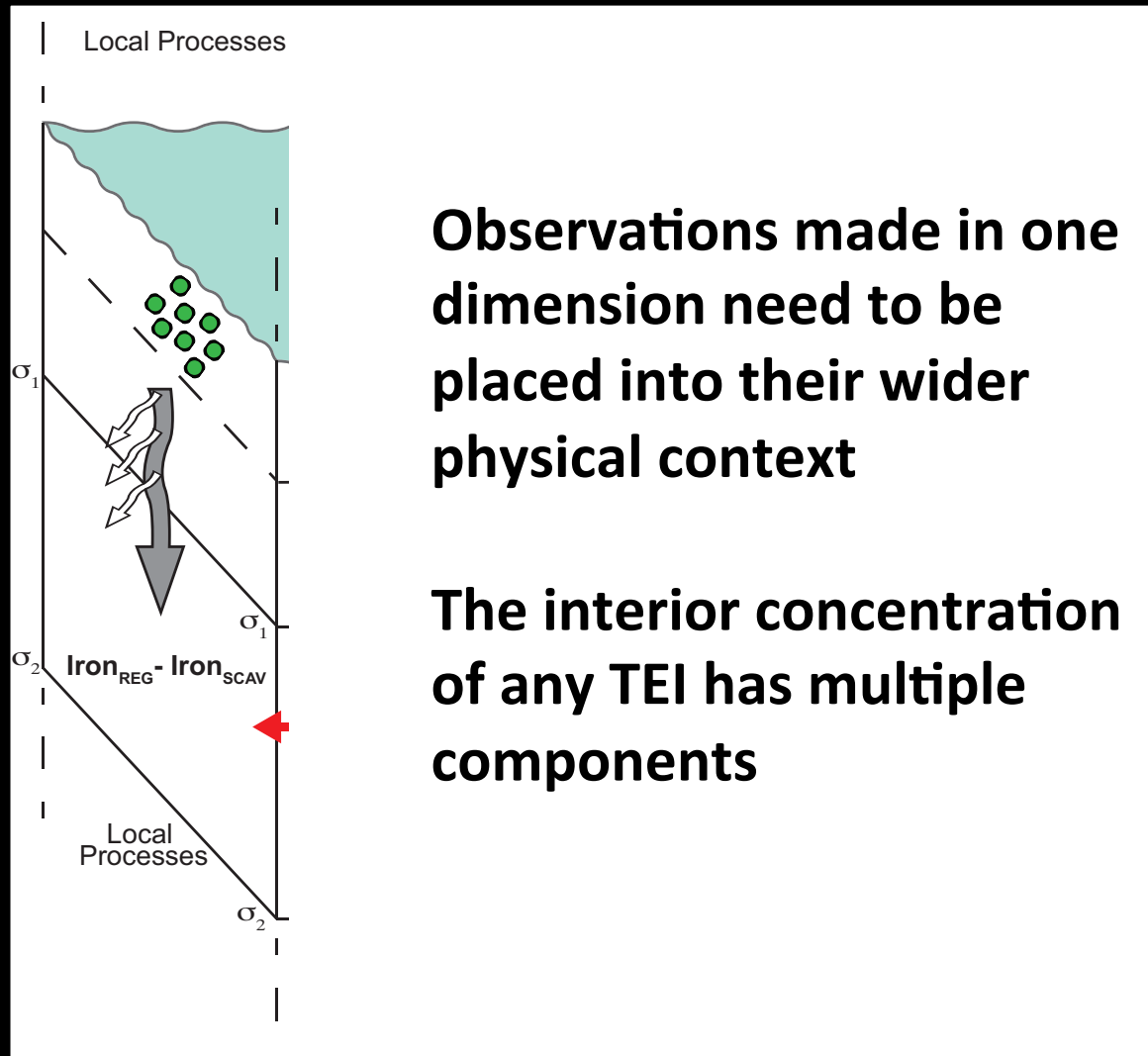
~~Sinking and~~ ✓ regeneration

✓ Scavenging

Subduction and transport

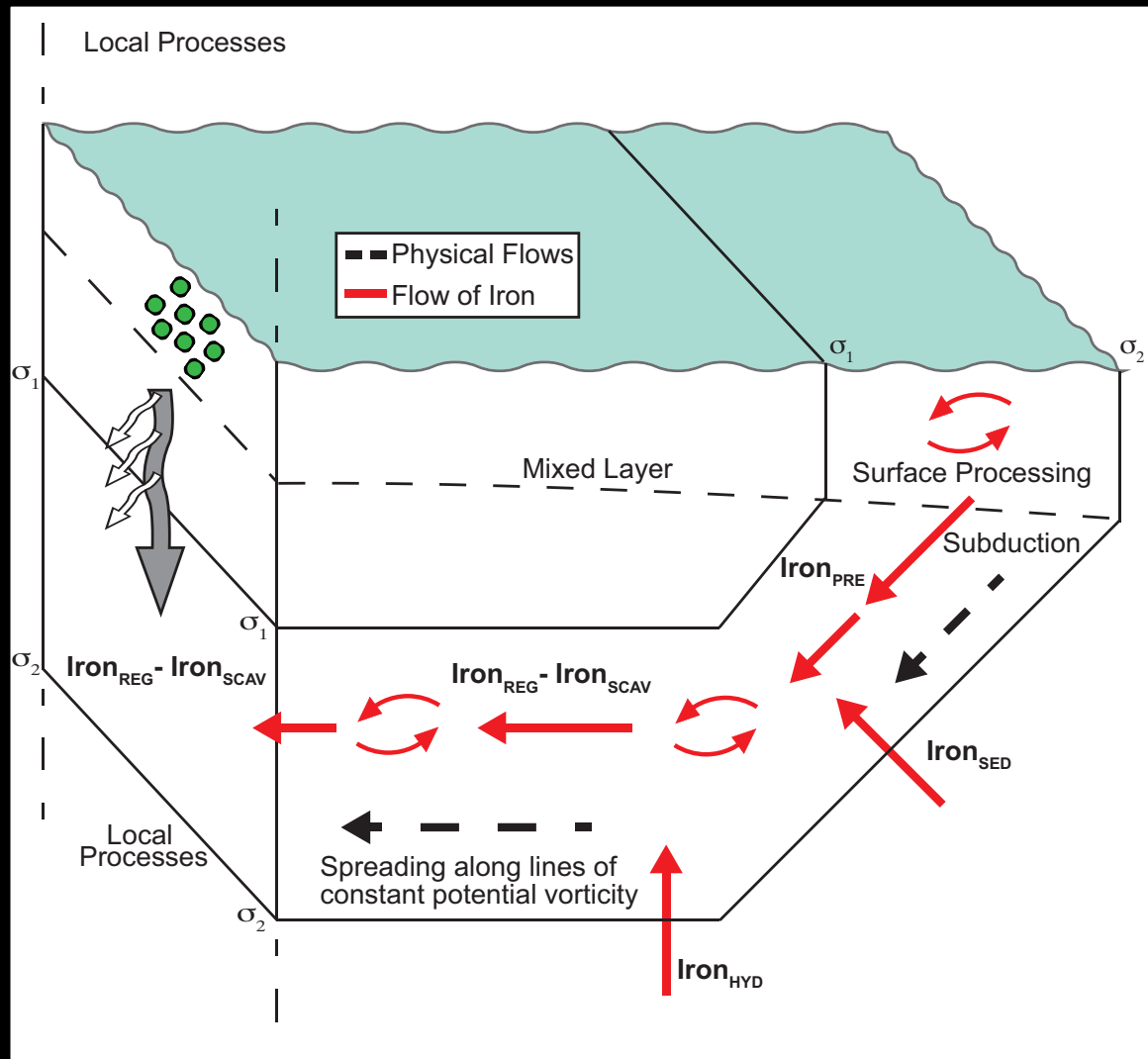


Subduction and Transport



$$D_{\text{Fe}} = D_{\text{Fe}_{\text{PRE}}} + D_{\text{Fe}_{\text{REG}}} + D_{\text{Fe}_{\text{SED}}} + D_{\text{Fe}_{\text{HYD}}} - D_{\text{Fe}_{\text{SCAV}}}$$

Subduction and Transport



$$D_{Fe} = D_{Fe_{PRE}} + D_{Fe_{REG}} + D_{Fe_{SED}} + D_{Fe_{HYD}} - D_{Fe_{SCAV}}$$

Summary of Processes and Assumptions

Surface uptake and cycling

- Affinity for Fe, quotas, physiology

Sinking and regeneration

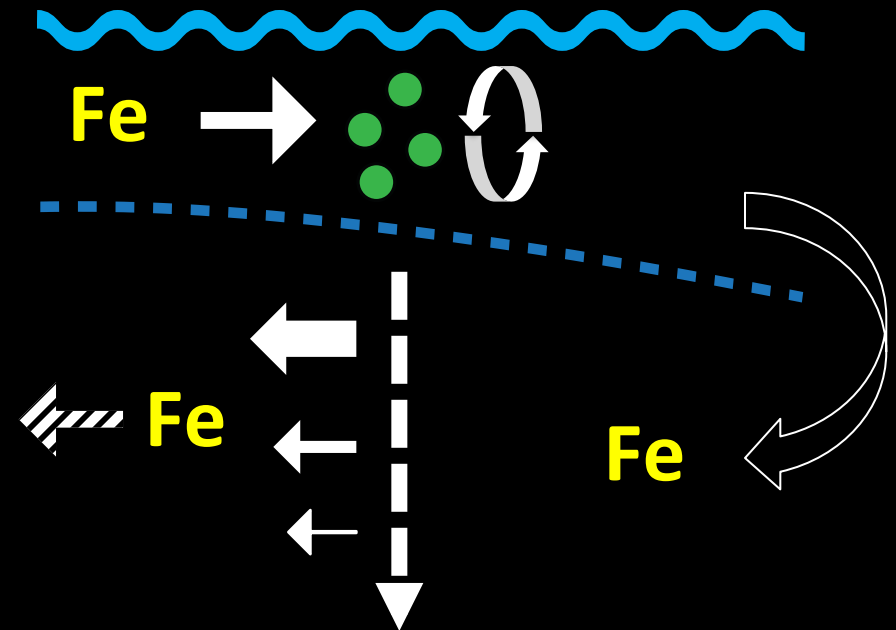
- TEI specific sinking or lability
- Zooplankton
- Bacteria

Scavenging

- Concentration and binding capacity of ligands
- 'Scavenging potential' of different particles

Subduction and transport

- Link physics to new theory / frameworks

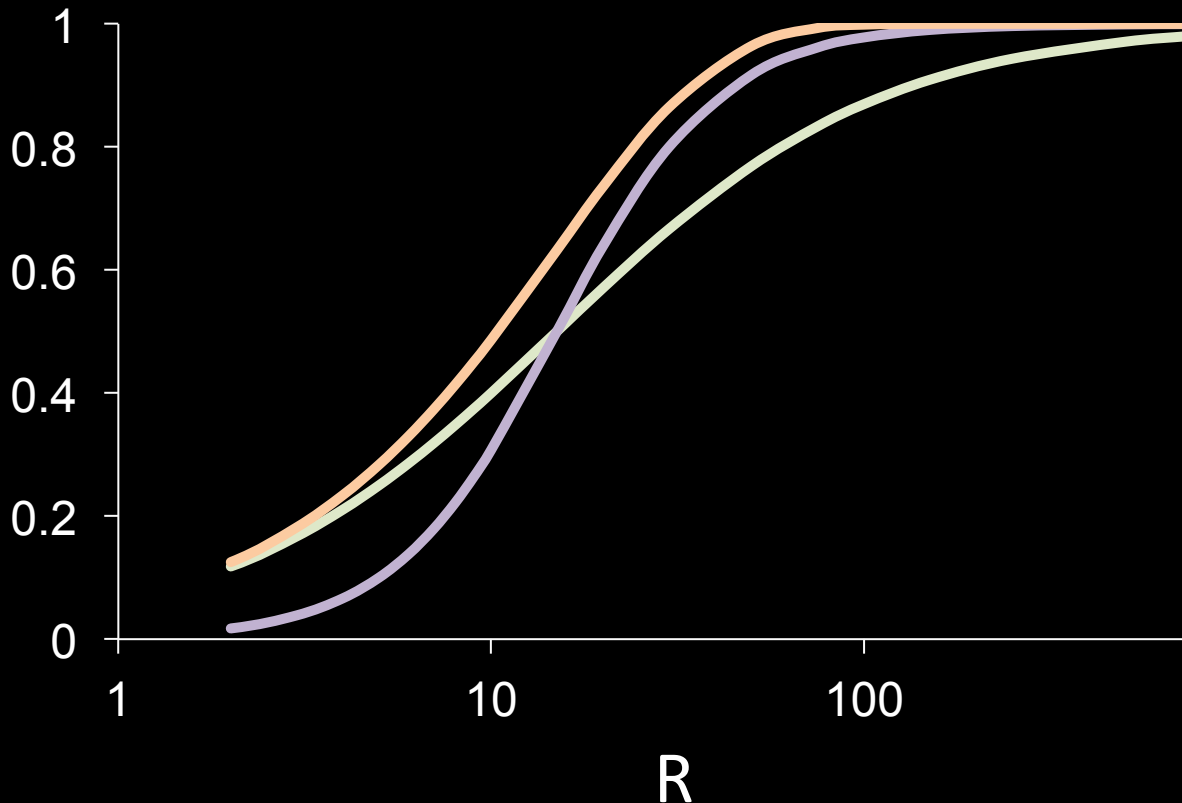


How are new processes added?

- Initially, fundamental mechanistic understanding is lacking
 - More approximate choices are needed!

How are new processes added?

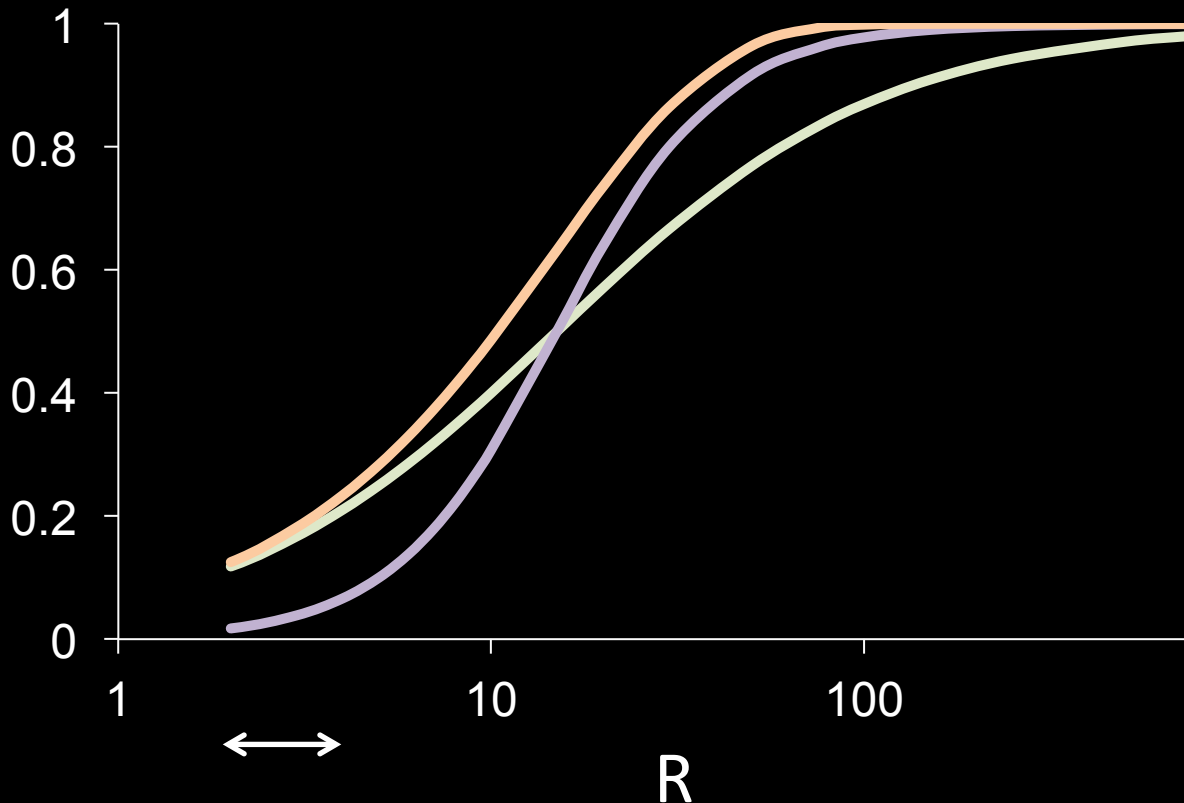
- What are the underlying dependencies and the functional form?



$$\frac{R}{R + K_R}$$
$$\frac{R^2}{R^2 + K_R^2}$$
$$1 - e^{-R/K_R}$$

How are new processes added?

- What are the underlying dependencies and the functional form?



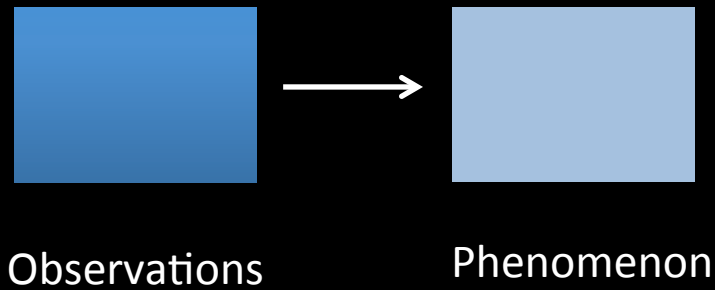
$$\frac{R}{R + K_R}$$
$$\frac{R^2}{R^2 + K_R^2}$$
$$1 - e^{-R/K_R}$$

Thresholds can be imposed (e.g. oxygen based)

Where do we need to go?

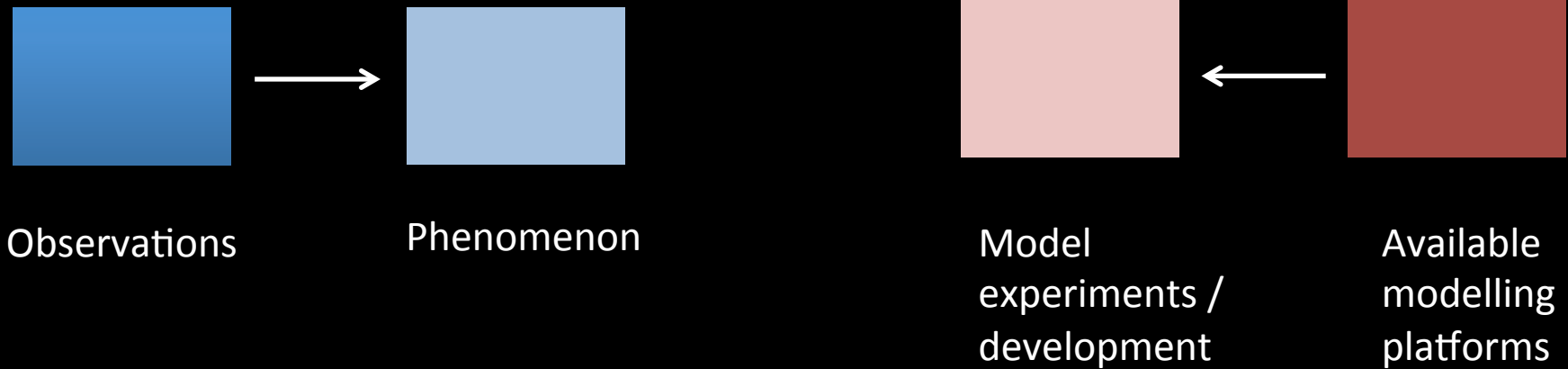
- When built to represent key processes of interest, models can be fantastic hypothesis testing tools
 - Underpinned by their choice of parameterisations and parameters
- Close links between modellers and experimental scientists most fruitful
 - Lots of collaborative opportunity!
 - How can we exploit GEOTRACES section and process studies to identify and constrain the key processes so their wider relevance may be estimated?

How might we get there?



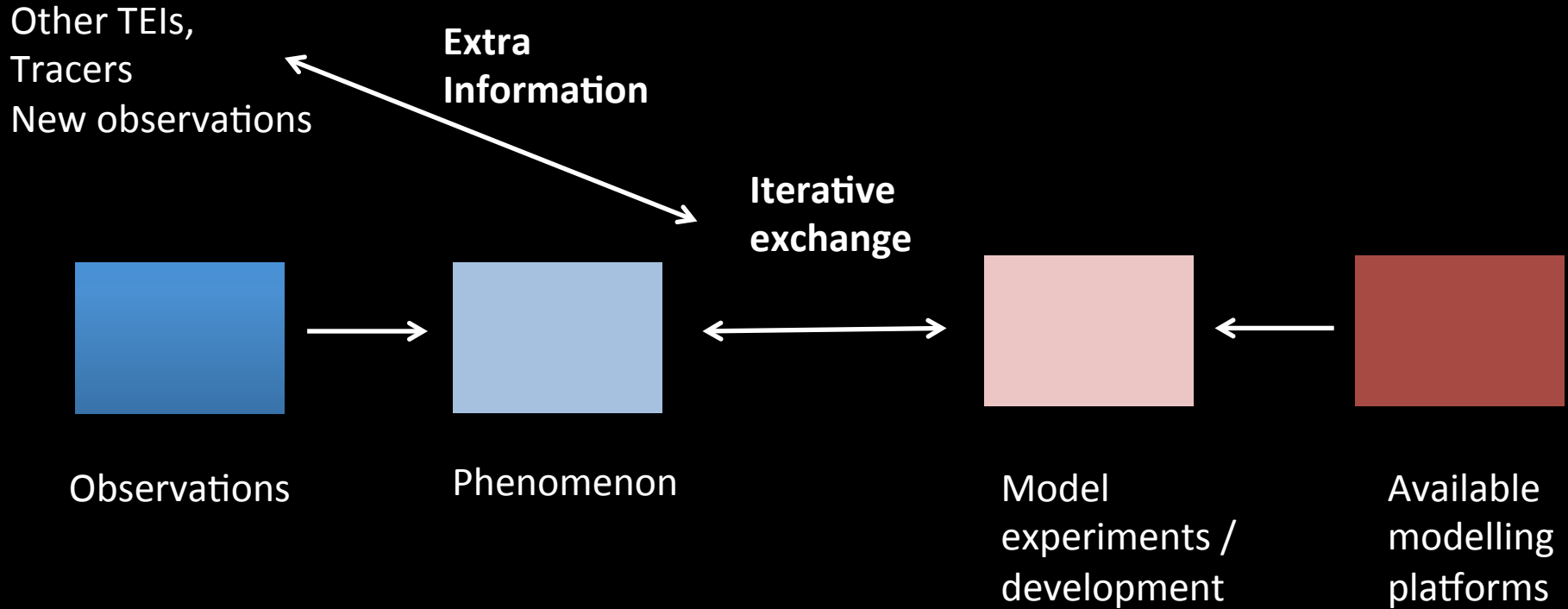
How important is this?

How might we get there?



How important is this?

How might we get there?



How important is this?

How might we get there?

