

Depth dependence of OUR – what does it mean? Horizontal vs. vertical C flux

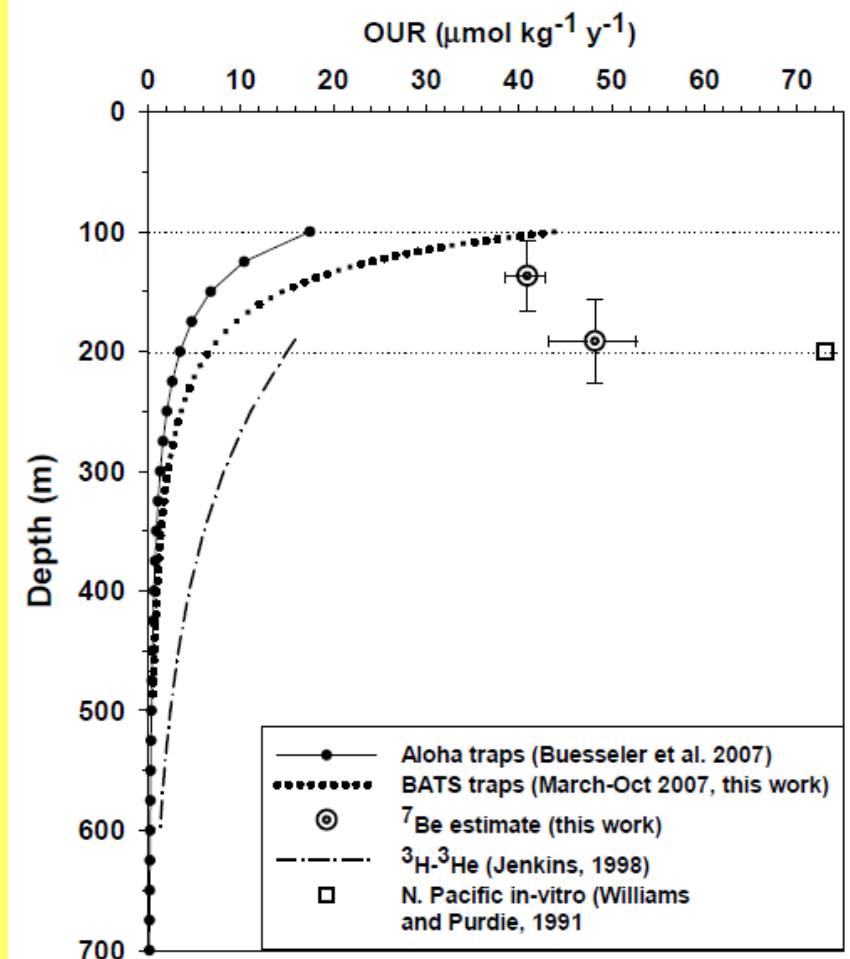
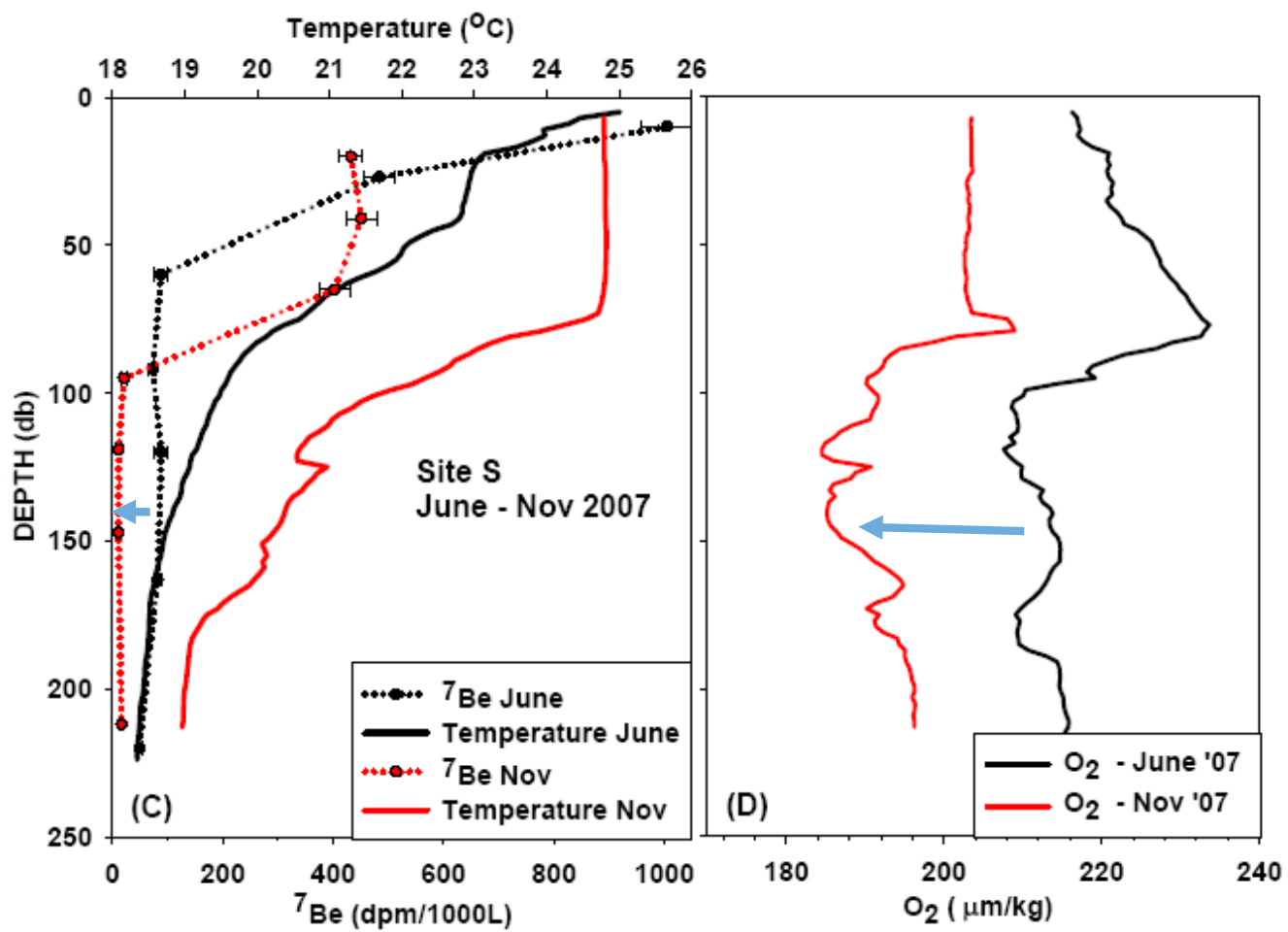


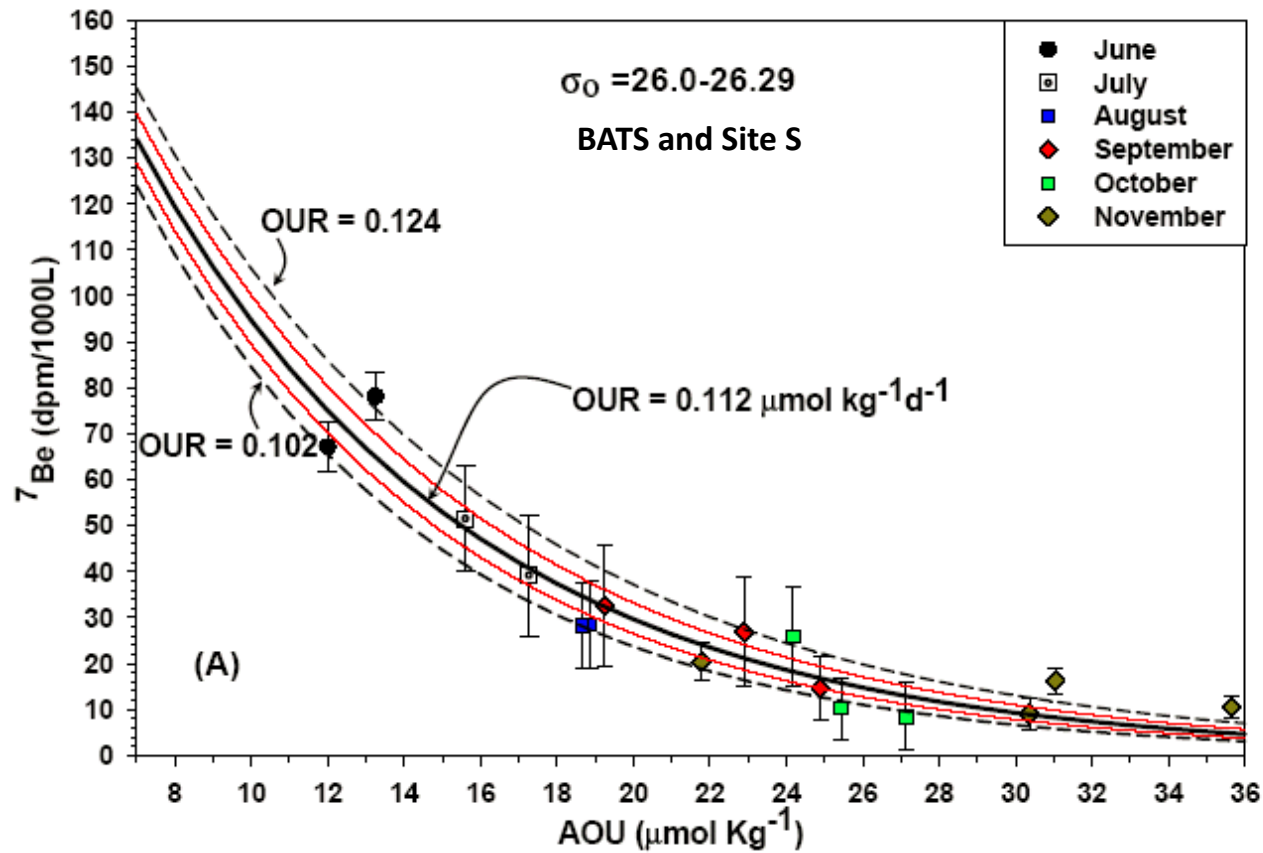
Fig: Oxygen Utilization Rate Profiles obtained by different methods (Kadko, 2009-GBC)

The ^7Be tracer ($T_{1/2}=53\text{d}$) technique allows estimation of shallow intense respiration rates

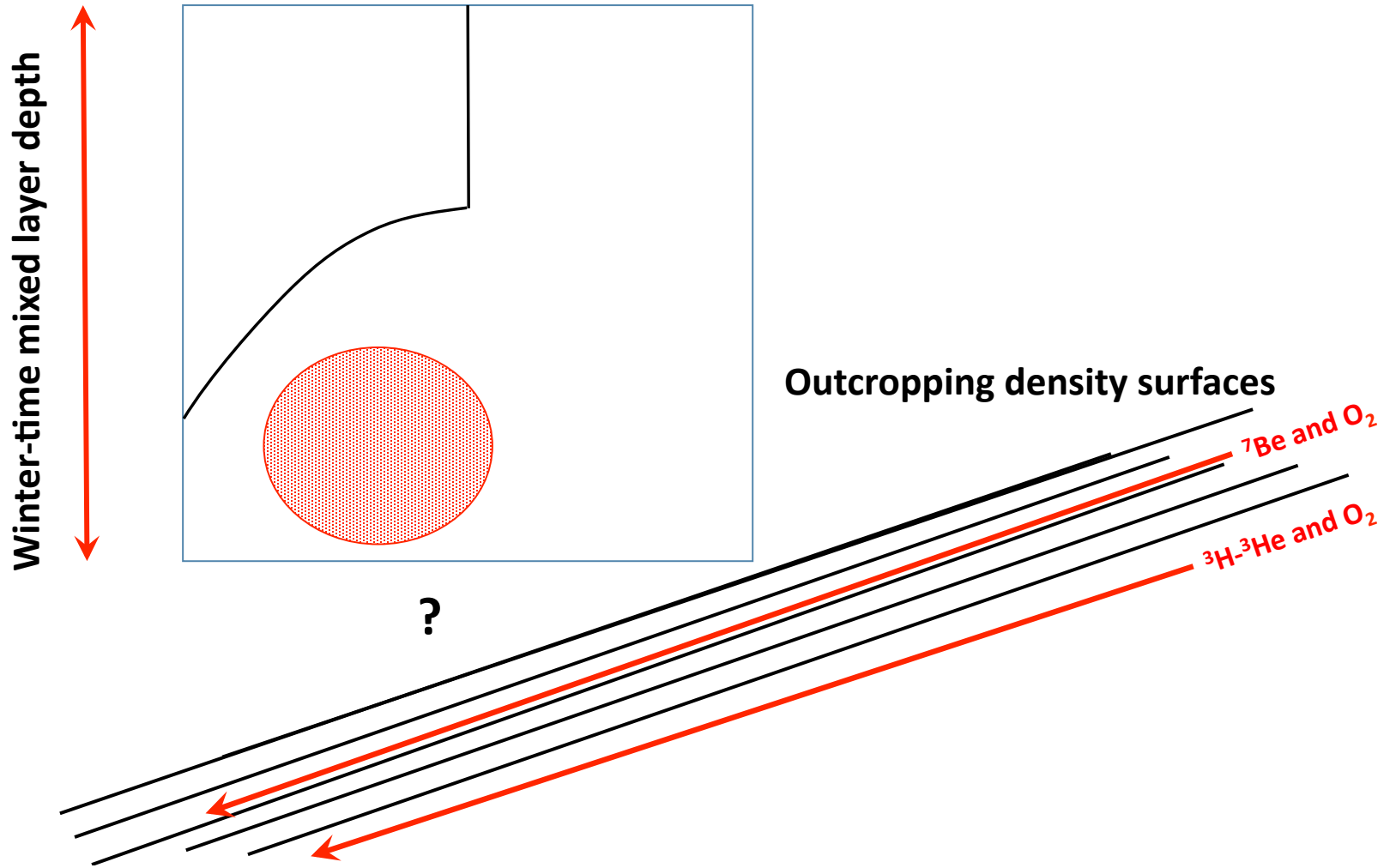
Depth-integrated C respiration rate based on O_2 utilization from ^7Be (100-200m) is $3.2 \text{ mol C m}^{-2} \text{ yr}^{-1}$. This is 65% of the total ($4.9 \text{ mol C m}^{-2} \text{ yr}^{-1}$) derived from the sum of ^7Be and ^3H - ^3He (200m-1000m) methods.

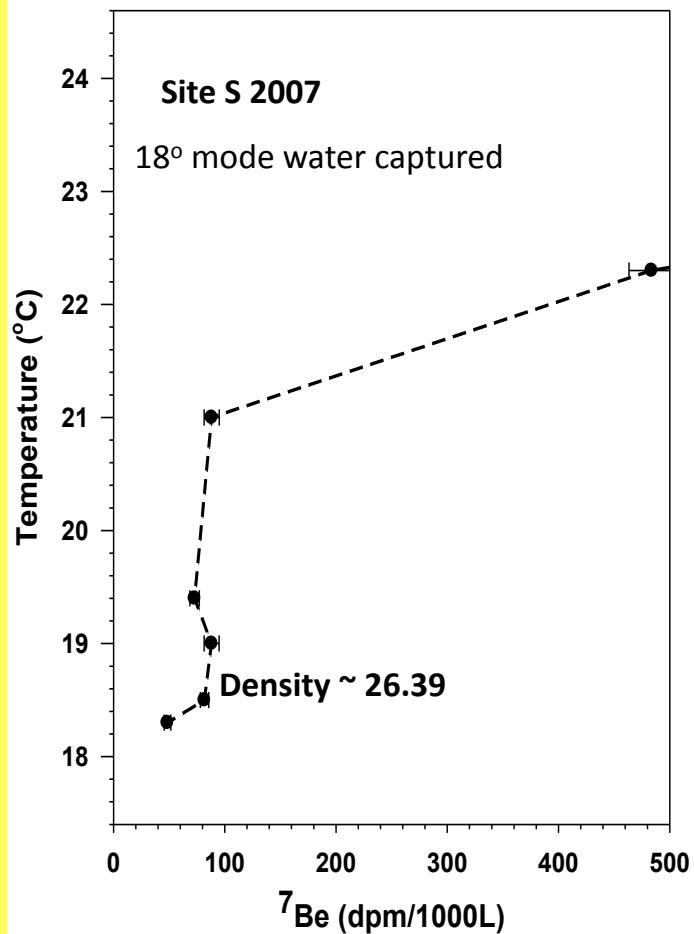


Rapid oxygen utilization in the ocean twilight zone assessed with the cosmogenic isotope ^7Be



Kadko, D. (2009) Global Biogeochem. Cycles





Initial, or upstream conditions?

With ${}^7\text{Be}$ flux = $0.05 \text{ dpm cm}^{-2}\text{d}^{-1}$, a mixed layer of 150 m would have ${}^7\text{Be} \sim 260 \text{ dpm/m}^3$.

At 160m, ${}^7\text{Be} = 82 \text{ dpm/m}^3$. This yields ~ 90 days.

At 160m, $\text{O}_2 = 209.9 \text{ } \mu\text{m/kg}$
 O_2 saturation = $233.8 \text{ } \mu\text{m/kg}$ } $\Delta\text{O}_2 \sim 24 \text{ } \mu\text{m/kg}$

Implies an OUR $\sim 0.25 \text{ } \mu\text{m kg}^{-1} \text{ d}^{-1}$

Kadko and Johnson, 2008 (ASLO)