

# Do we know $K_d$ ? “The slime problem”

Challenge: Phoebe's data shows different  $K_d$  for different mineral phases  
But we know that most particles are covered by organics

Some key questions:

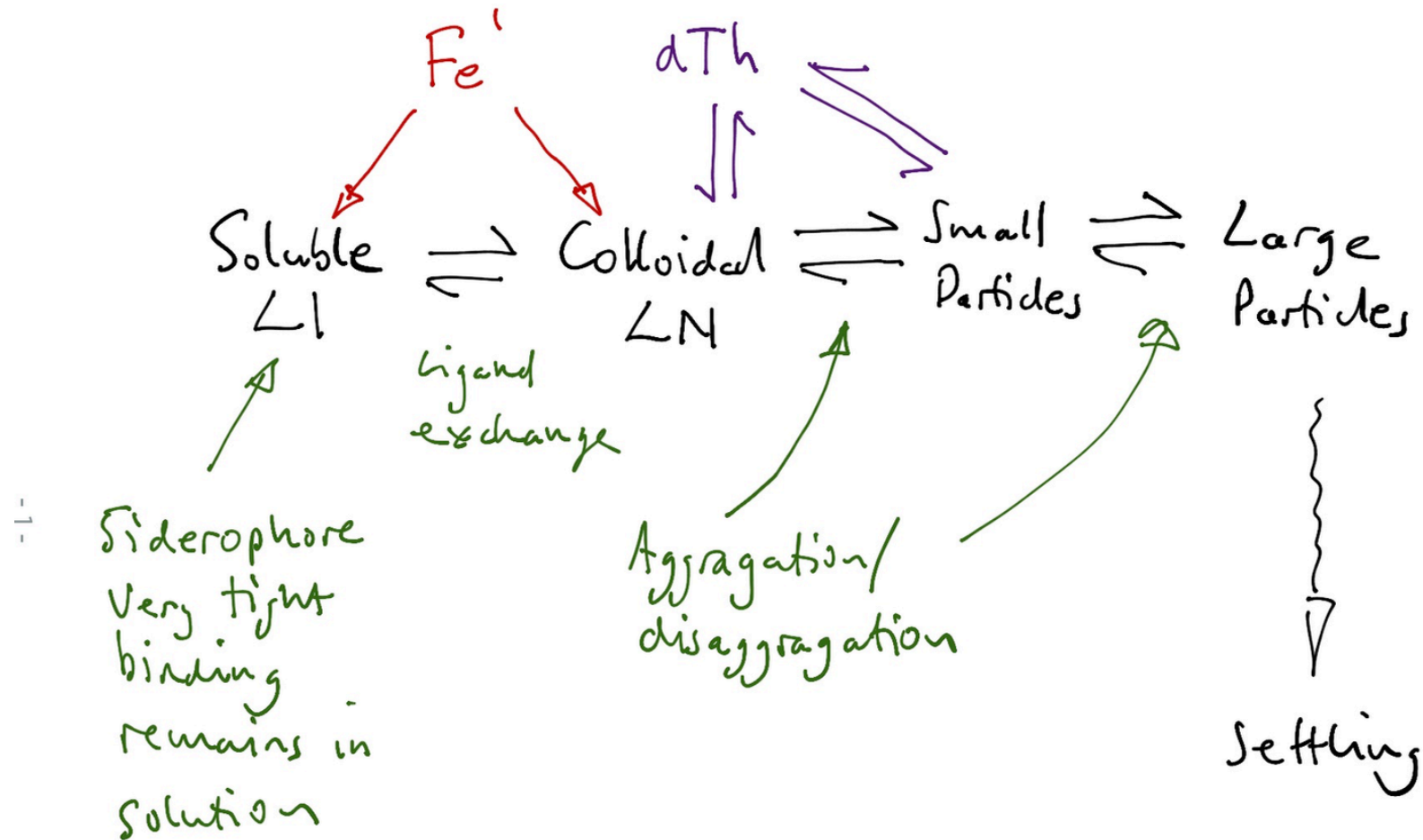
- what is rate of  $K_{on}$  and  $K_{off}$  for metals, and for organics to mineral phases
- how constant are organic coatings with location and particle type?
- is mass the best thing to use to characterise scavenging? Should we use surface area and surface charge instead?
- are lab surface area measurements relevant to the field?
- how much variation is there in  $K_d$  and organic coatings with latitude, productivity,  $O_2$  levels, etc.?
- Mn and Fe oxyhydroxides have similar surface area but very different  $K_d$ ; different organic coating?

Some good news, though:

- just adding Mn-oxyhydroxides with high  $K_d$  is a step forward (and realised in PISCES now)

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Classes are operationally defined

## Some immediate products:

Th isotopes useful for aggregation/disaggregation;  
extend this to colloid-small particle (*Jess/Tom/Paul*)

Provocative review of organic complexation, exchange,  
and aggregation/disaggregation model (*Kristen/Jess/  
Tim/Phoebe/...*)

We are close to being able to model L1 field in ocean

If...

L1 field is modelled, and Th gives us particle dynamics...  
we have quantified major aspects of metal scavenging.

## Nepheloid layers

They are very different from open ocean in terms of particle composition (as well as concentration)

Not all layers are the same (at all!)

Oxyhydroxides are important (and possibly predictable given overlying productivity flux)

Need to target some key locations with range of measurements